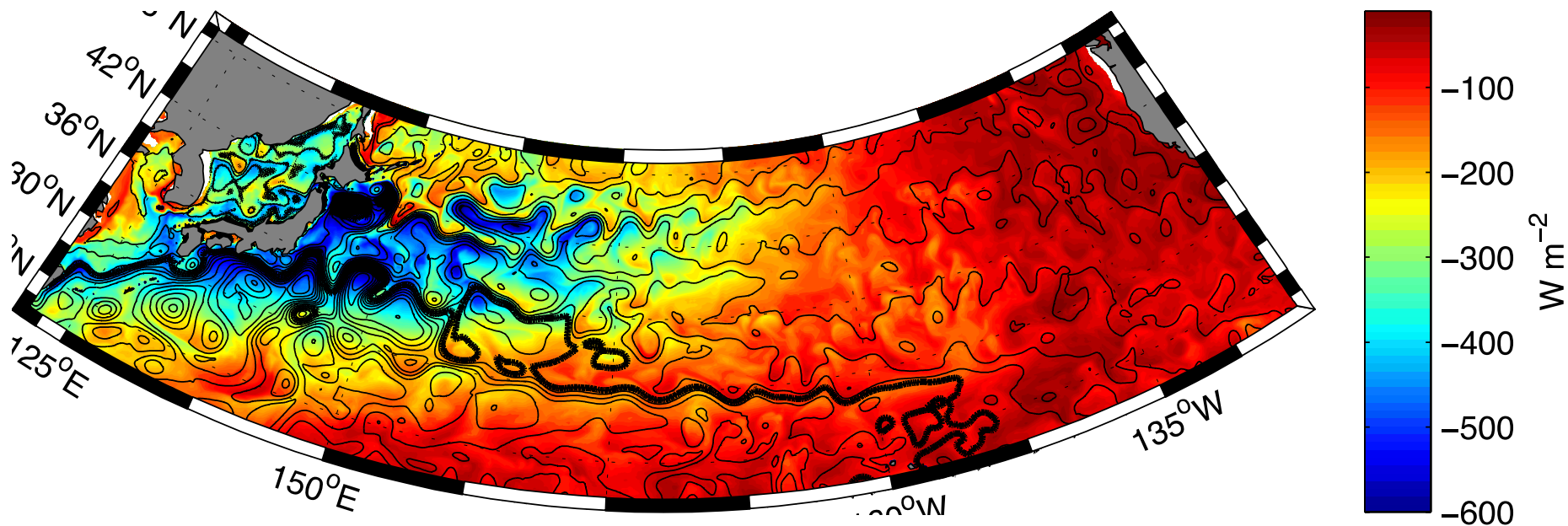


The influence of mesoscale eddies on wintertime air-sea interaction in the Kuroshio Extension region

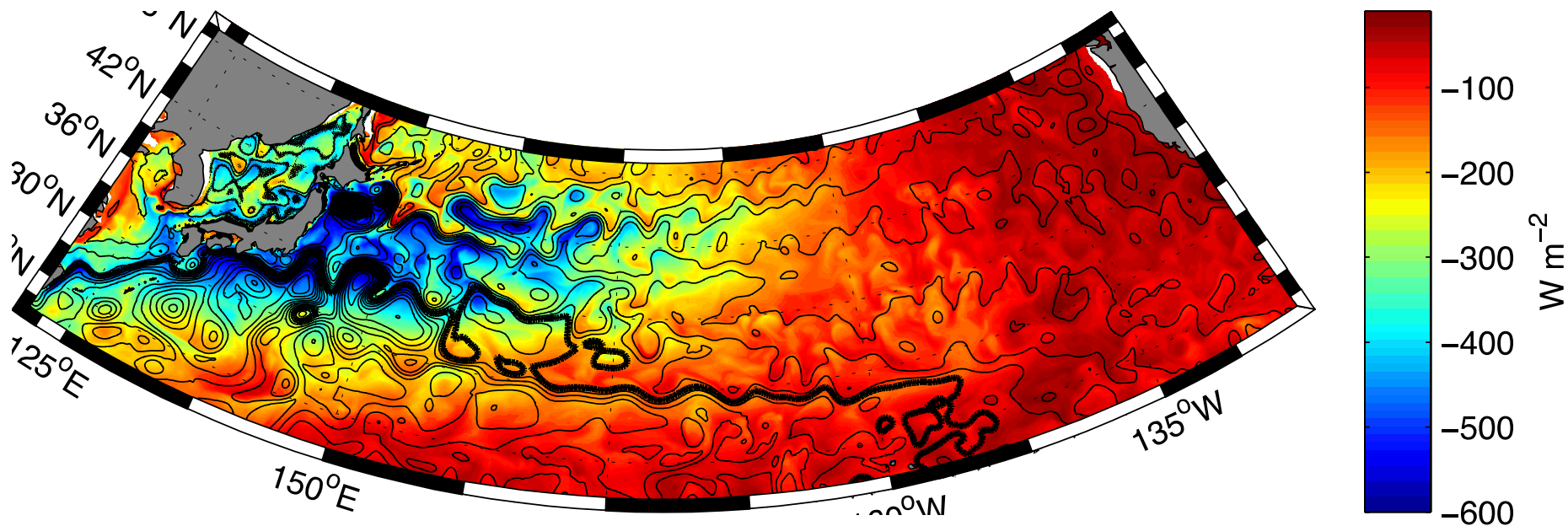


Stuart P. Bishop, Frank O. Bryan, and Justin Small

National Center for Atmospheric Research

CESM Workshop, Breckenridge June 17, 2014

Evidence for Bjerknes compensation in the North Pacific



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Outline

- Introduction to meridional heat transport (MHT) and eddies in the ocean and atmosphere
- Feedbacks of ocean eddies on the atmosphere
- High resolution (HR) fully-coupled climate simulation using CESM
- Ocean and atmospheric response in HR CESM
- Evidence for “Bjerknes Compensation” in the North Pacific

Meridional Heat Transport

- Earth receives insolation from the sun:

• Surplus of heat in the tropics

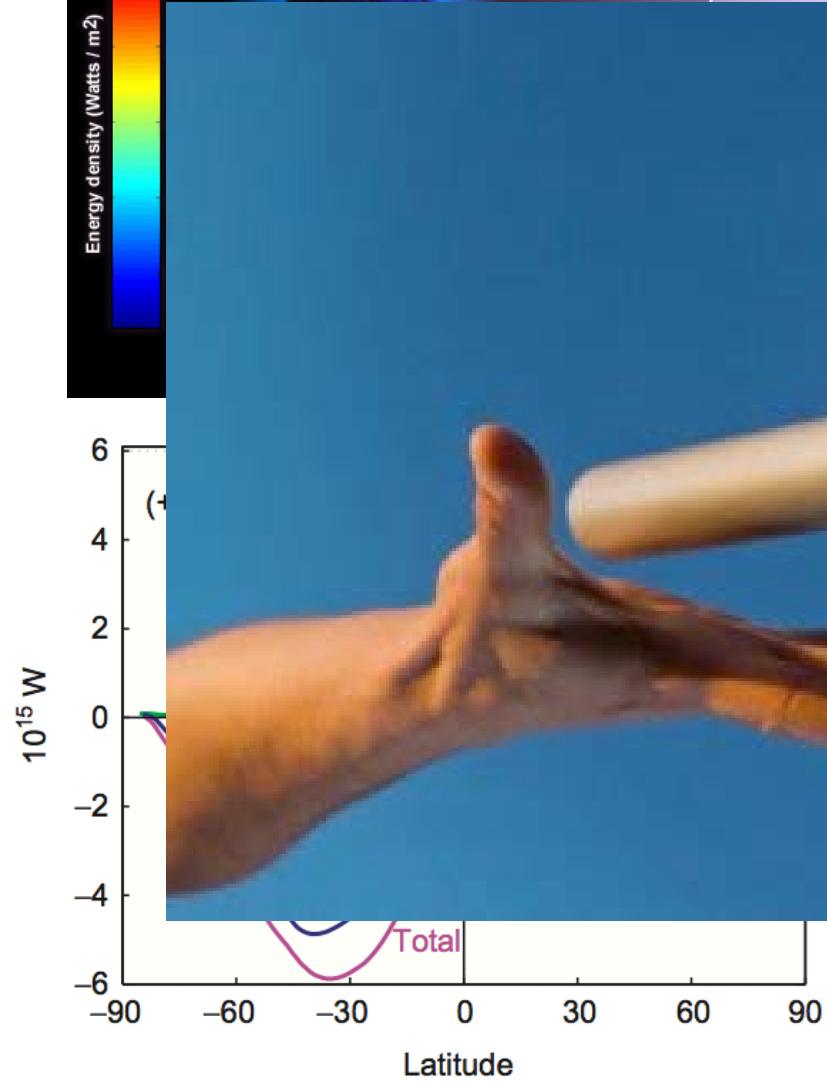
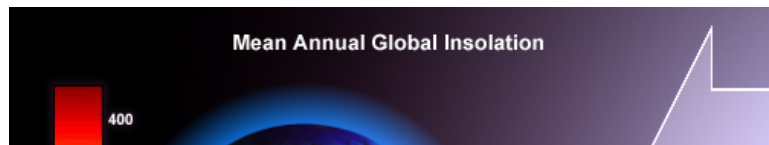
ex

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at
sphere

- How does the ocean do this?



From *Ocean Circulation and Climate: A 21st Perspective*

Bjerknes Compensation

- Bjerknes (1964), Atlantic Air-Sea Interaction
- When net TOA radiation is zero, if there is a reduction in MHT by the ocean this reduction must be compensated by the atmosphere

$$\int_0^L \int_{-h}^0 \rho_0 C_p^o \overline{v_o T_o} dz dx = - \int_0^L \int_0^H \rho_a C_p^a \overline{v_a T_a} dz dx$$

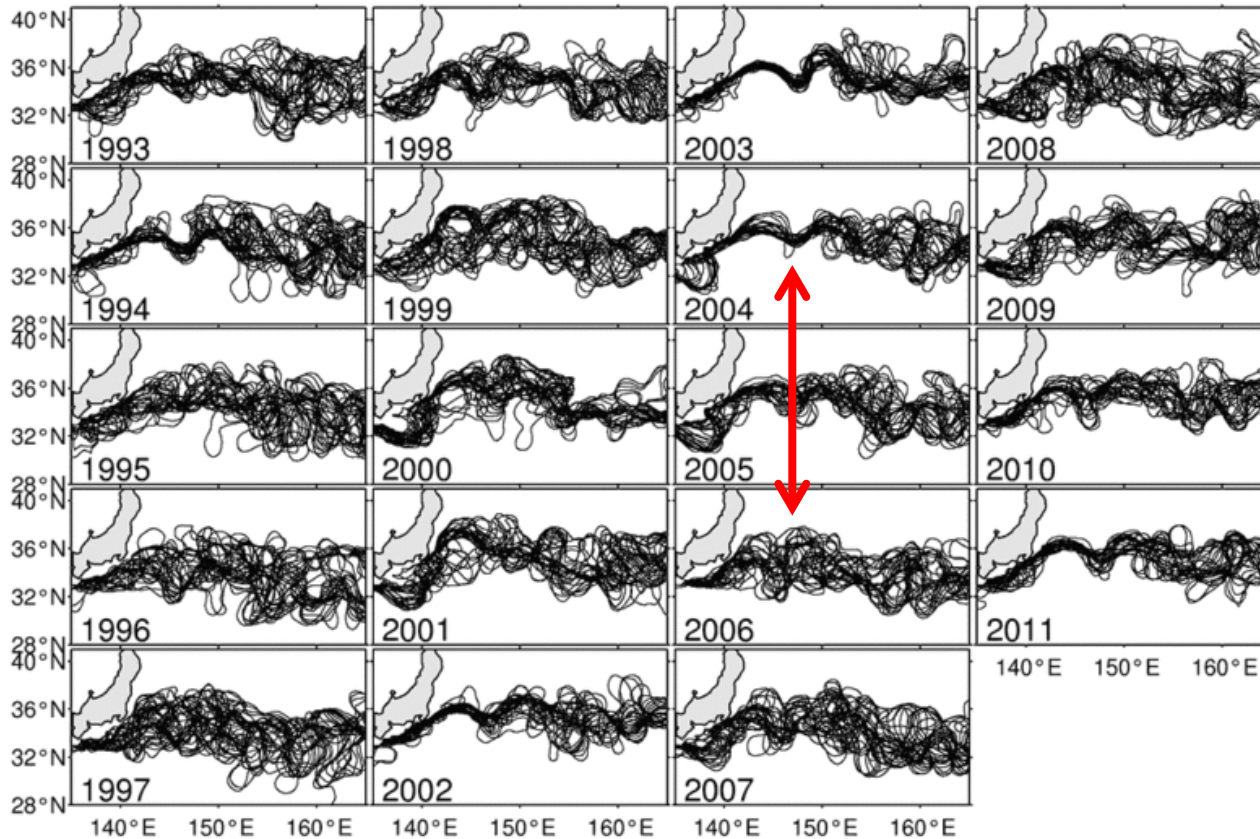
- North Atlantic was favored as a region for this due to the MOC and decadal time scales linked to the AMO
- Farneti and Vallis (2013) quoting Bjerknes (1964) “the atmosphere is likely to drive the ocean at interannual time scales, while for decadal-multidecadal periods it is ocean dynamics that prevail.”



Srokosz et al. 2013, BAMS

Kuroshio Extension

Decadal Variability

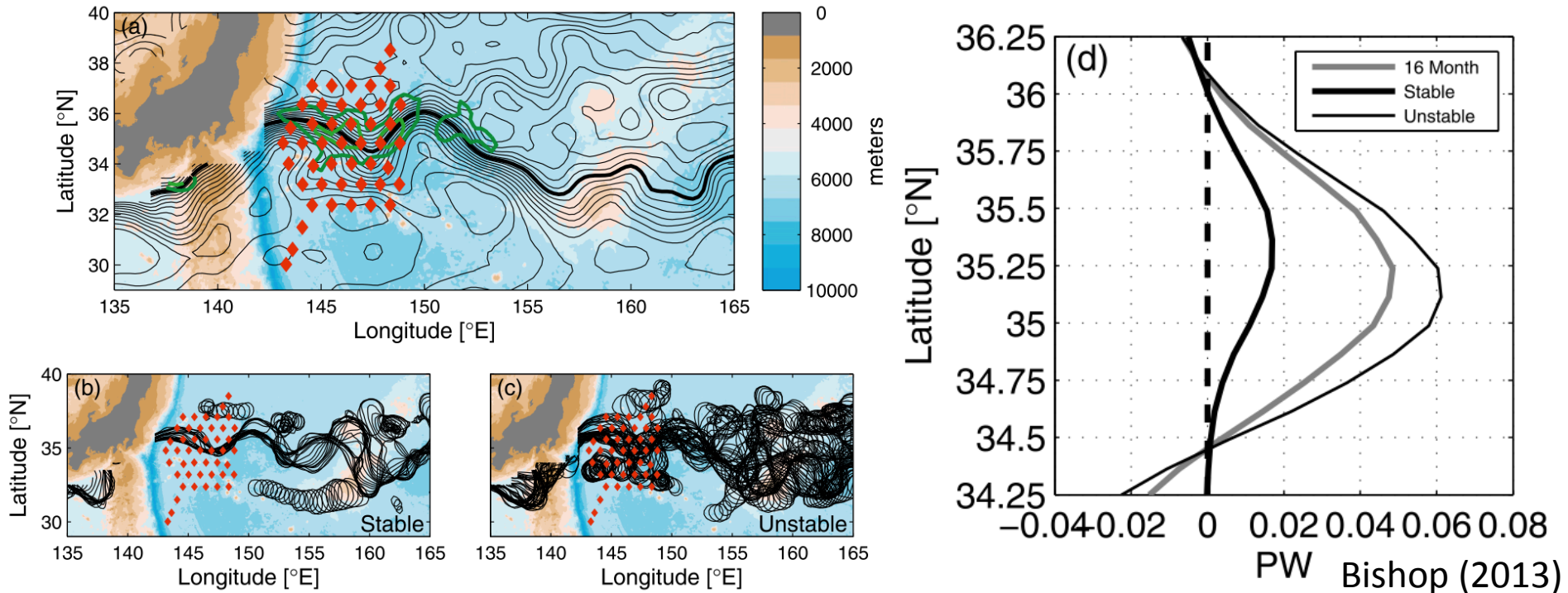


Updated from Qiu and Chen 2005

<http://www.aviso.altimetry.fr/en/news/idm/2012/mar-2012-following-kuroshio-paths.html>

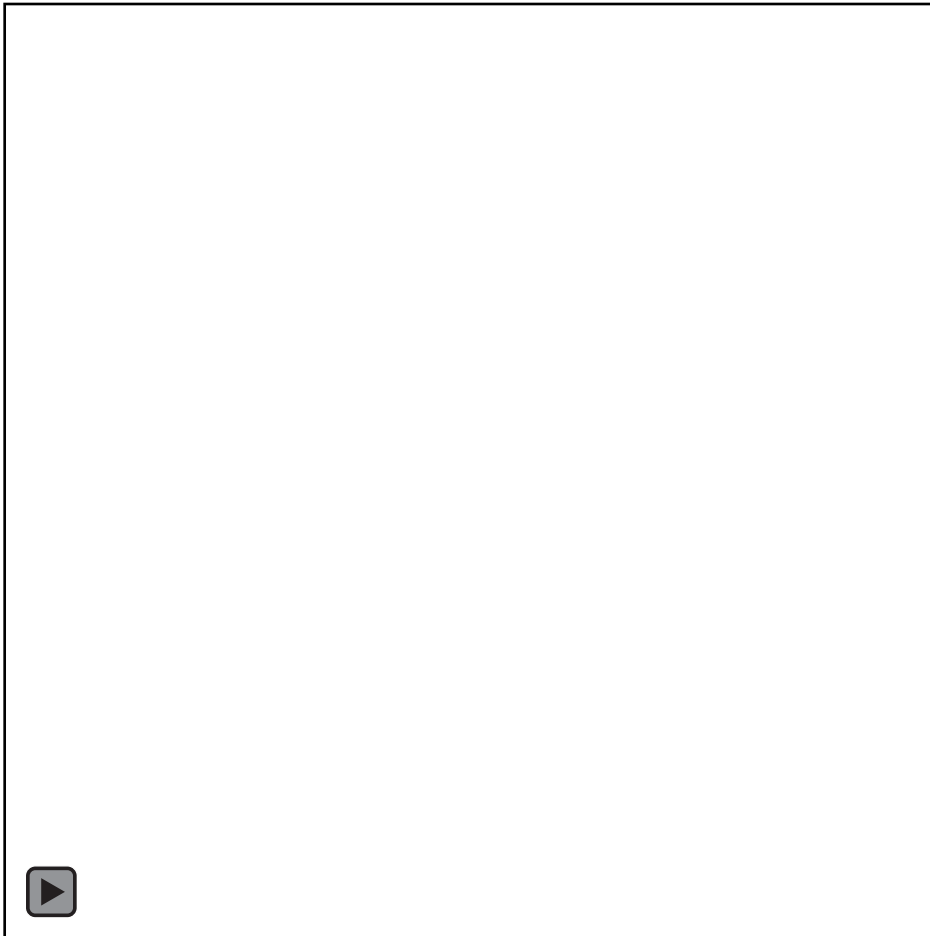
- “Stable” years
 - 1993-1994, 2002-2005, 2011
- “Unstable” years
 - 1995-2001, 2006-2010
- Stable characteristics:
 - Strong transport (strong meridional SST gradient)
 - Weak small amplitude eddies
 - Small meridional eddy heat transport (Presumably)
- Unstable characteristics:
 - Weaker transport (weaker meridional SST gradient)
 - Strong high amplitude eddies
 - Large meridional eddy heat transport

Kuroshio Extension System Study (KESS)



- Observational field study from 2004-2006
- Captured a transition from stable to unstable state
- Meridional eddy heat transport was 3X larger during the unstable state

High-Resolution Fully-Coupled CESM Simulation



Movie courtesy of Tim Scheitlin (NCAR Visualization Lab)

Model

- Community Earth System Model:
 - 0.1° Ocean (POP2)
 - 0.25° Atmosphere (CAM5)
- Fully coupled (atmosphere-ocean)
 - 6 hourly
- 100 years
- Monthly archived data
- Advantages of model over observations:
 - Subsurface oceanic data
 - Long time series for statistical robustness

ASD run status and acknowledgements

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3 and an eddy-resolving ocean model.

4 Justin Small, Julio Bacmeister, David Bailey, Allison Baker, Stuart Bishop, Frank Bryan, Julie
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8 ABSTRACT

9 High-resolution global climate modeling holds the promise of capturing planetary-scale climate
10 modes and small-scale (regional and sometimes extreme) features simultaneously, including
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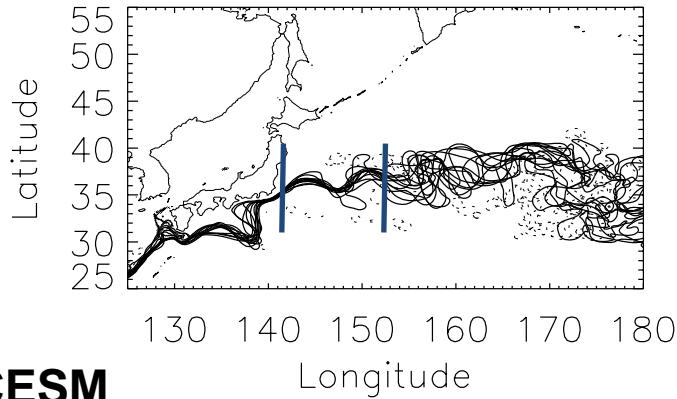
Sensitivity of modeled dynamics to atmosphere and ocean resolution in the CESM

Tomas, Small, Tseng, Bailey and ASD group

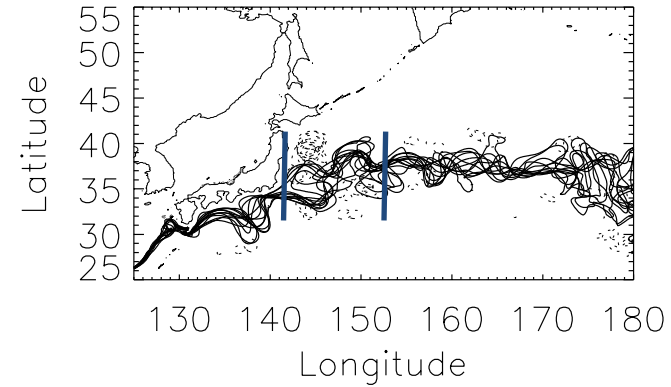
DATA AVAILABLE ON EARTH SYSTEM GRID

Decadal Variability in CESM

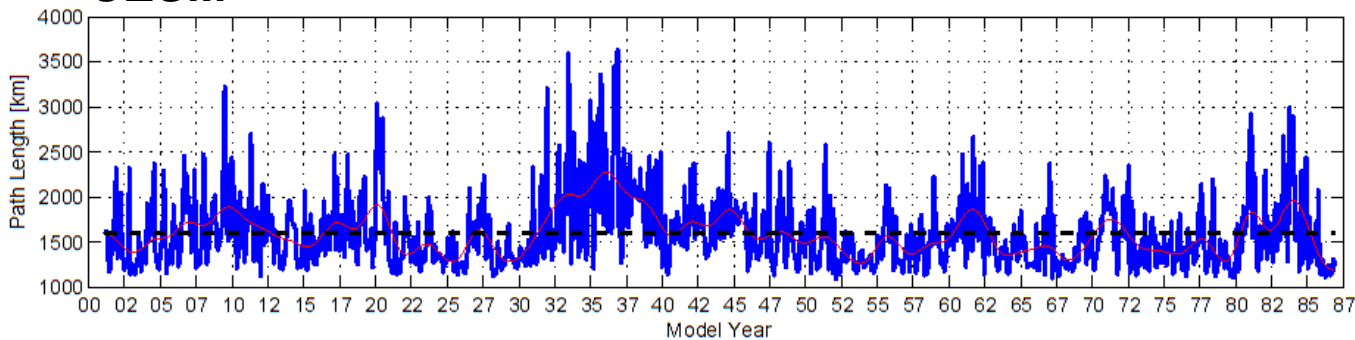
Year 0015



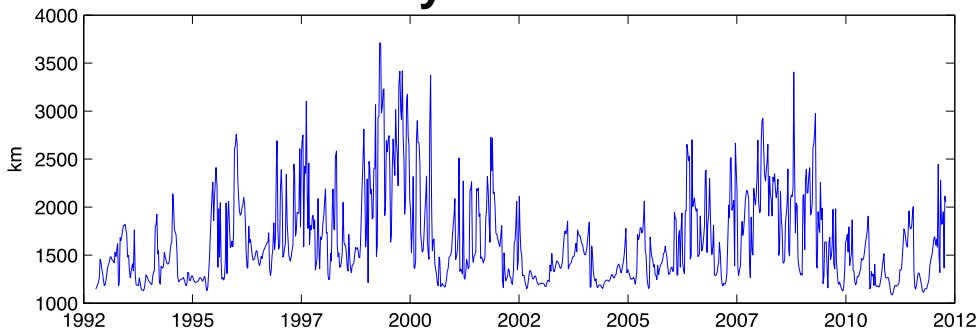
Year 0037



CESM

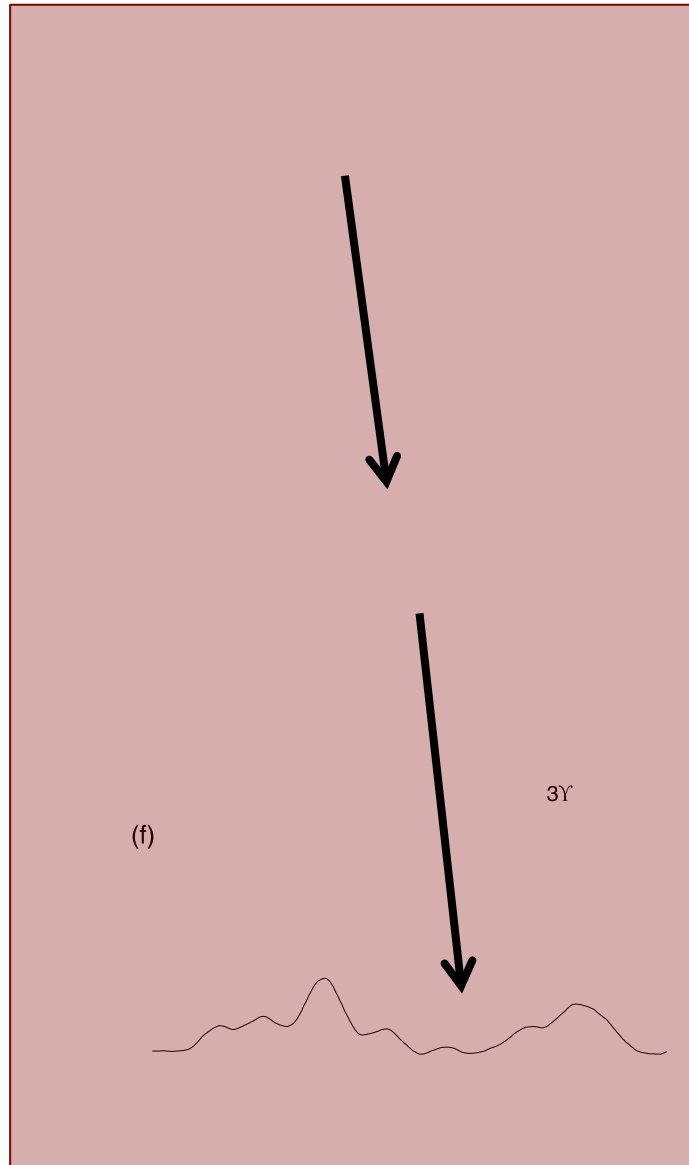
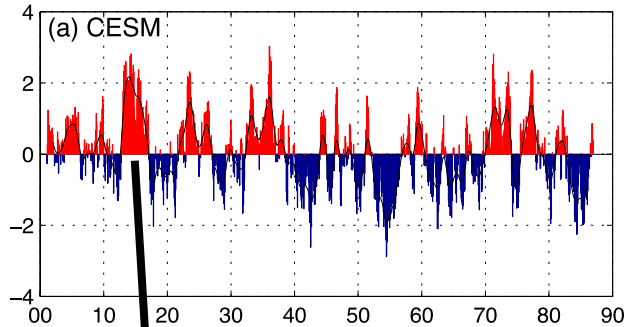


AVISO Altimetry Observations



- Path length between 141° - 153° E
 - Longer path length = more eddies
 - Shorter path length = less eddies
- Model exhibits variability similar to observations

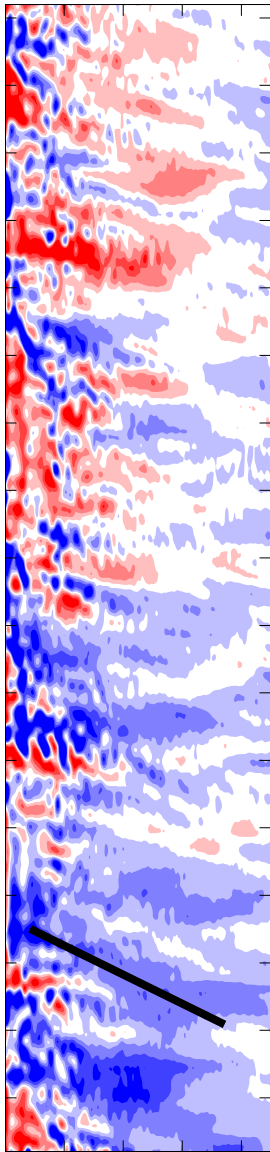
Kuroshio Variability and PDO



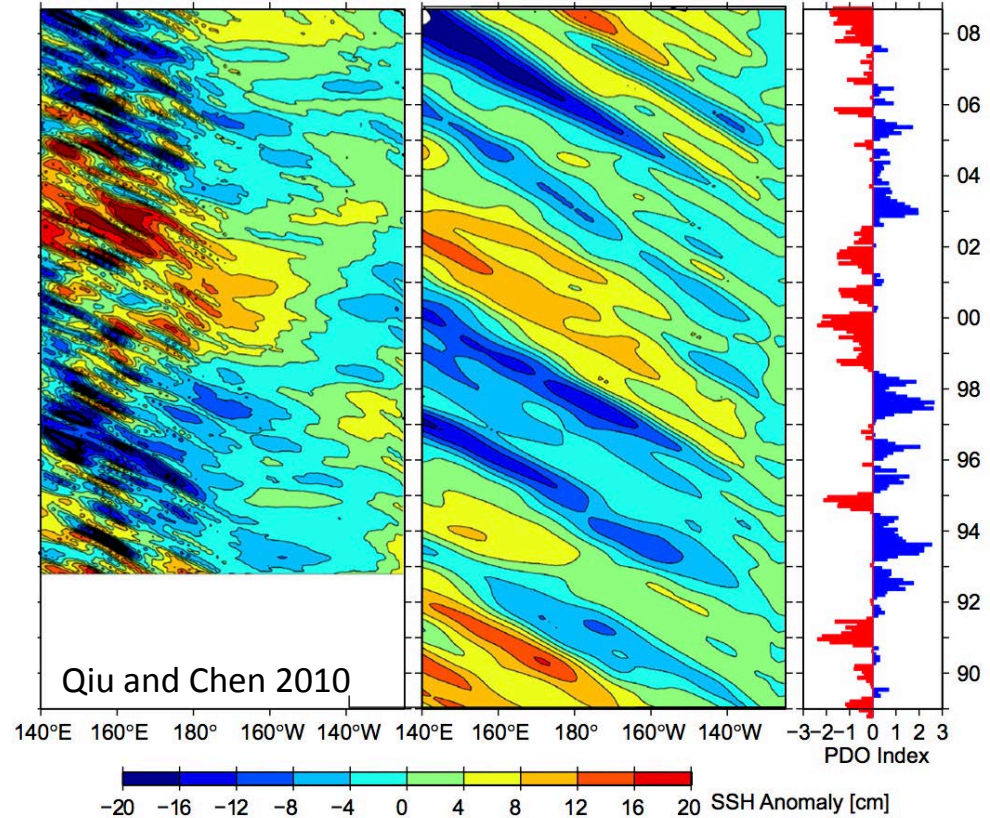
- Phase of PDO leads jet latitude by $\sim 3-4$ years
- Seen in both observations and model
- Jet latitude precedes meander variability
 - Northerly jet has a more “stable” meander state
 - Southern jet has a more “unstable” meander state

PDO & SSH Anomalies

CESM

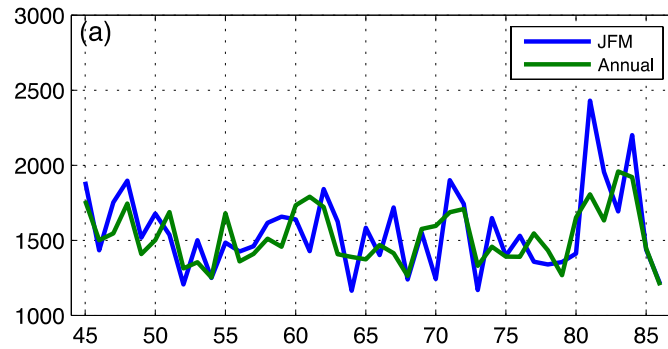


Observations+Rossby Wave Model



- Phase of PDO drives SSH anomalies that propagate west from the central Pacific
- SSHa's drive meridional shifts of the jet
 - Positive SSHa shifts the jet **North**
 - Negative SSHa shifts the jet **South**

“Waviness” Index (KEI)

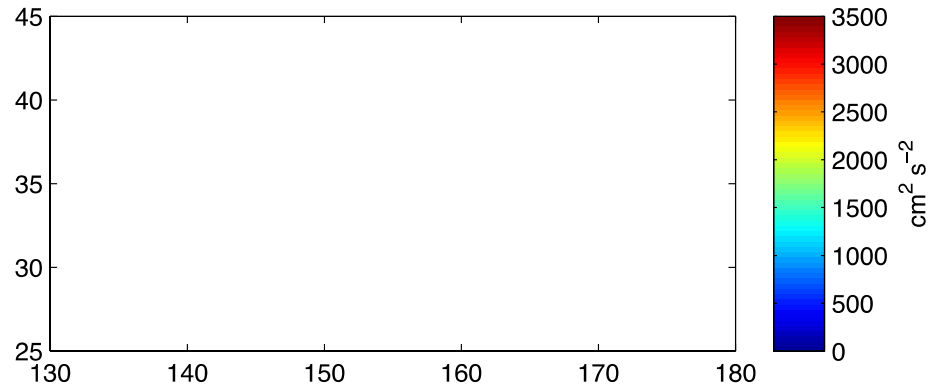


Bishop et al. in prep

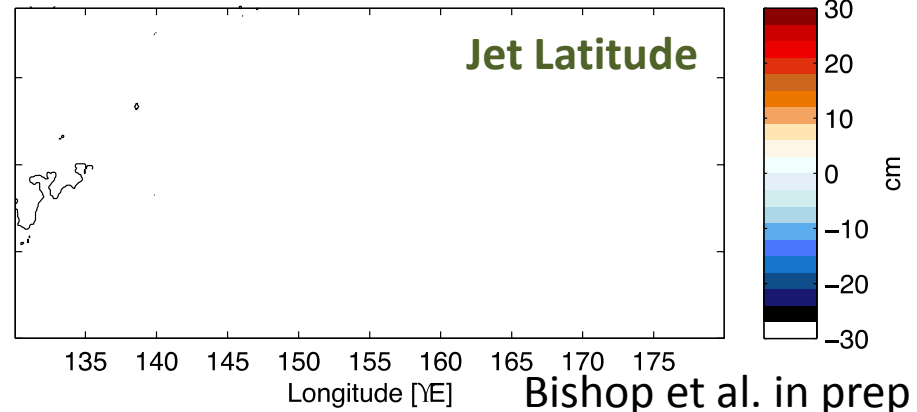
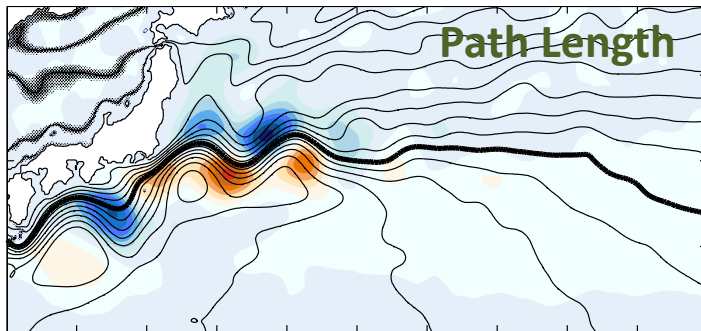
- Index: negative normalized path length
 - +ve “Stable”
 - -ve “Unstable”
- Index: JFM mean (blue), annual mean (green)
- Most years if JFM “stable” annual mean is “stable” too (68% chance)
- Composite threshold = 0.25 (1/4 standard deviation)

Stable vs. Unstable SSH

Composite (± 0.25) of SSH and EKE



Regression of SSH onto normalized path length and jet latitude

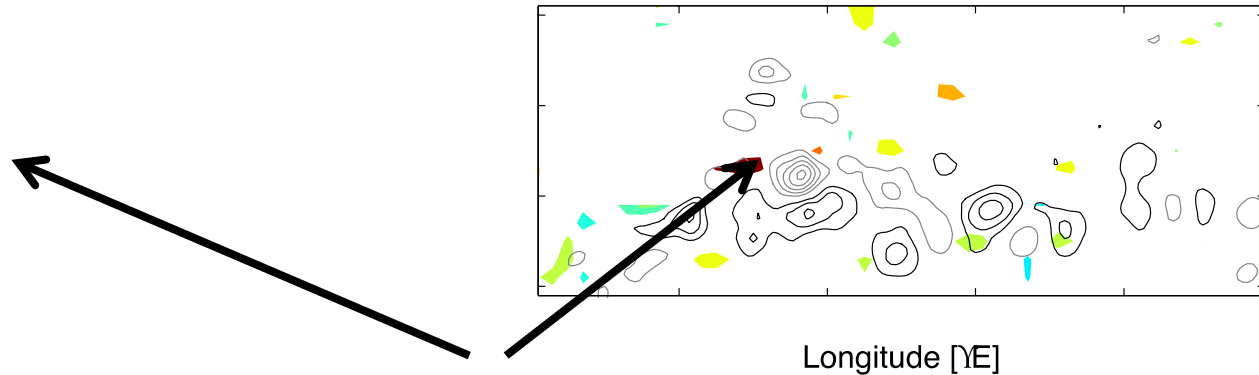


- Stable vs. Unstable:
 - Tighter front (ΔSSH of 84 vs. 60 cm between 34° - 36°N at 145°E)
 - Tight swath of EKE along the jet
 - SSH is higher (lower) to the south (north) – higher transport
 - Stronger Southern Recirculation Gyre

Wintertime Surface Heat Flux

CESM

OAFlux



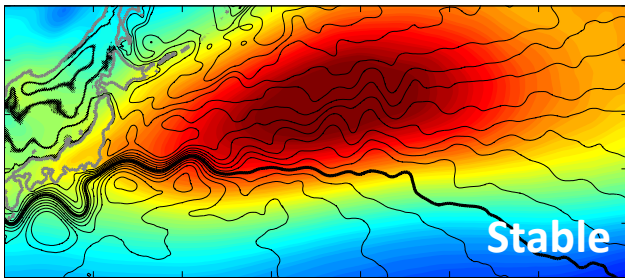
Excess fuel for storms in stable states

Bishop et al. in prep

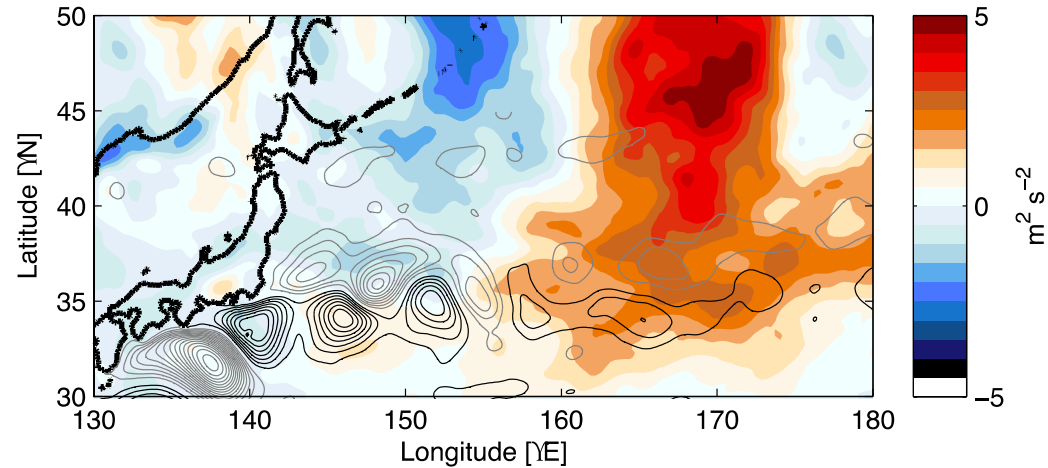
- JFM composite difference shows enhanced SHFs during stable meandering states
 - $\sim 100 \text{ W m}^{-2}$ north and near the coast
 - Excess fuel for storms?

Atmospheric Response: Storm Track & Eddy Heat Flux

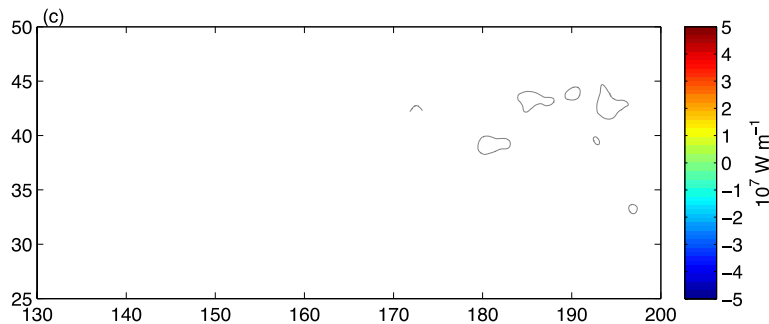
Composite of vertically-integrated EHF



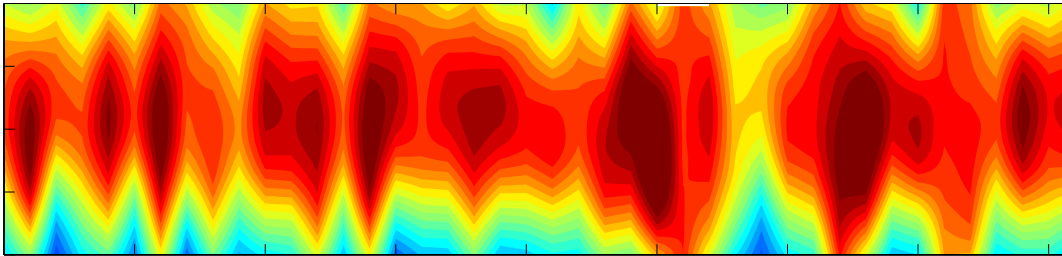
Regression of JFM $V'V'$ onto KEI



- Storm track is enhanced along and north downstream.
- Meridional eddy heat fluxes (MEHFs) are enhanced during stable meandering states.
- MEHFs are enhanced geographically where storm track is enhanced.



Meridional Eddy Heat Transport (Atmosphere)



stream Axis [km]

(c)

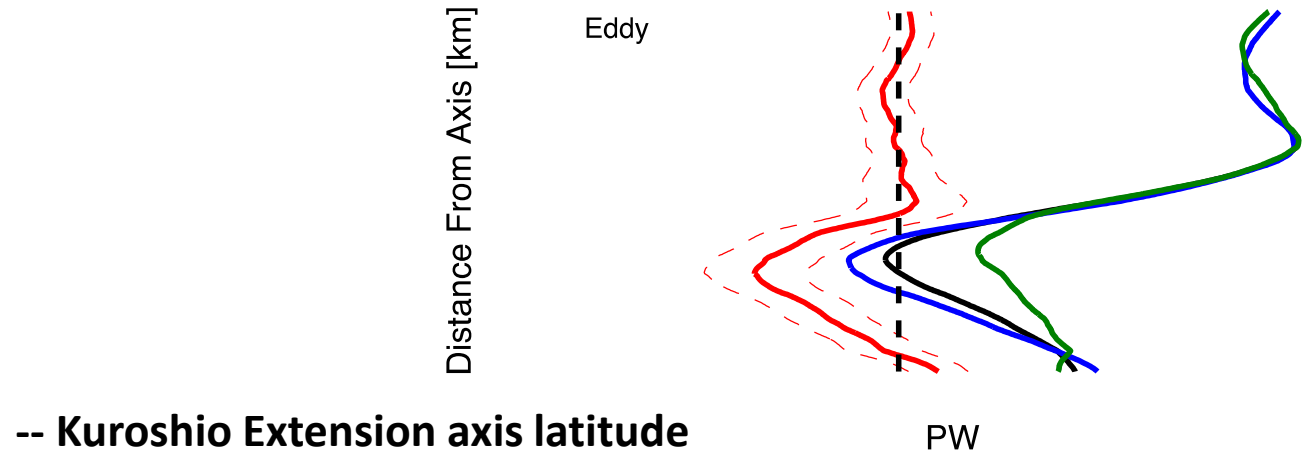
Difference

$$\frac{1}{g} \int_{p_0}^{p_s} (C_p \overline{v'T'} + L \overline{v'q'}) dp$$

- MEHT:
 - combination of temperature and moisture fluxes
 - Vertically integrated from 50 hPa → surface
 - Zonally integrated across Pacific basin

- MEHT was calculated relative to Jet Stream axis (maximum 850 hPa zonal mean wind, Ari Solomon personal communication)
- MEHT is enhanced by ~0.1 PW at its maximum during +ve KEI

Meridional Eddy Heat Transport (Ocean)



- MEHT has a range of ± 0.3 PW
- Stable states have the persistence of negative MEHT just south of the Kuroshio Extension jet core
- Aoki et al. (2013) notes persistence of negative MEHT south of the Gulf Stream and Kuroshio Extension (only 5 years of OFES simulation)
- Attributed to southward migration of warm water cores from cold-core rings interacting with meanders
- Unstable state has a larger MEHT by 0.07 ± 0.02 PW

Local Bjerknes Compensation

stream Axis [km]

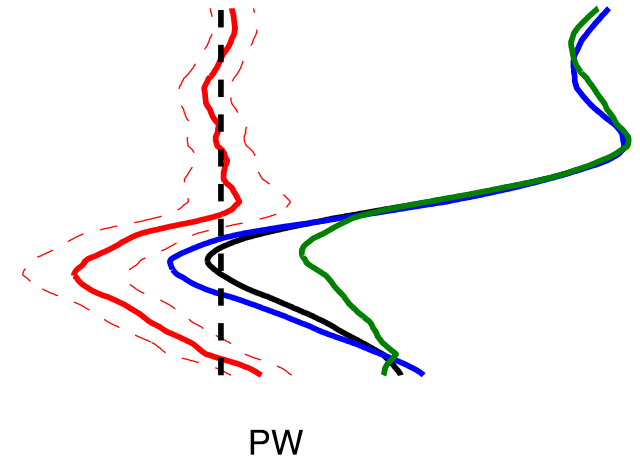
Atmosphere

(c)

Ocean

Eddy

Distance From Axis [km]



- The atmosphere and ocean have a degree of compensation (other studies did not find a perfect compensation in the N. Atl.):
 - Compensation is not coincident in space or time.
 - Atmosphere compensates during winter
 - Future work will look at this compensation throughout the year
- Local compensation was not expected since the atmosphere could compensate for the ocean anywhere on the globe!

Conclusions & Future Work

- Wintertime air-sea fluxes affect atmospheric circulation in high-resolution climate simulations with important implications for climate
- Fully-coupled simulation resembles observations with interannual to decadal variability in the Kuroshio Extension
- The atmosphere and ocean partially compensate as in “Bjerknes Compensation”
- Future Work:
 - Determine what the origin is of negative MEHT south of the jet during stable states.
 - Perform a full ocean heat budget to determine what processes are responsible for driving the atmospheric response in stable vs. unstable states (advection, heat content tendency, etc...)

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