

Firn modelling in CESM

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Motivation

1. Meltwater can refreeze in snow/firn, thus limiting the amount of freshwater ocean runoff and increasing SMB
2. Glacial inception: firn is thought to turn to ice when it reaches a density of 850 kg/m^3 . The depth at which this happens varies.
3. Ice sheet instability is linked to firn air depletion:

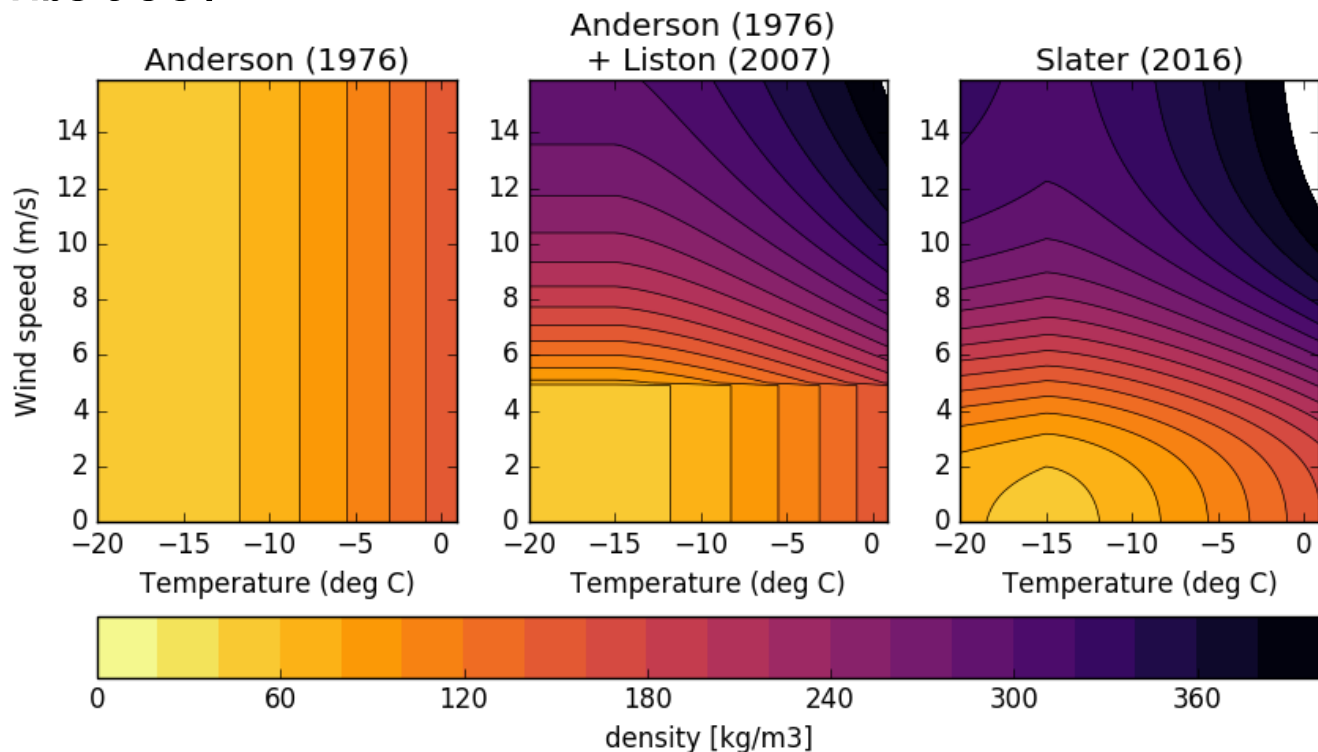
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Firn air depletion as a precursor of Antarctic ice-shelf collapse

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Code changes past year

- Introduced a deeper snowpack, by default 10-metre depth W.E. ~ 25 m actual depth
- Wind-driven fresh snow density
- Tuning of destructive metamorphism, overburden pressure and fresh snow density [Drew Slater & Dave



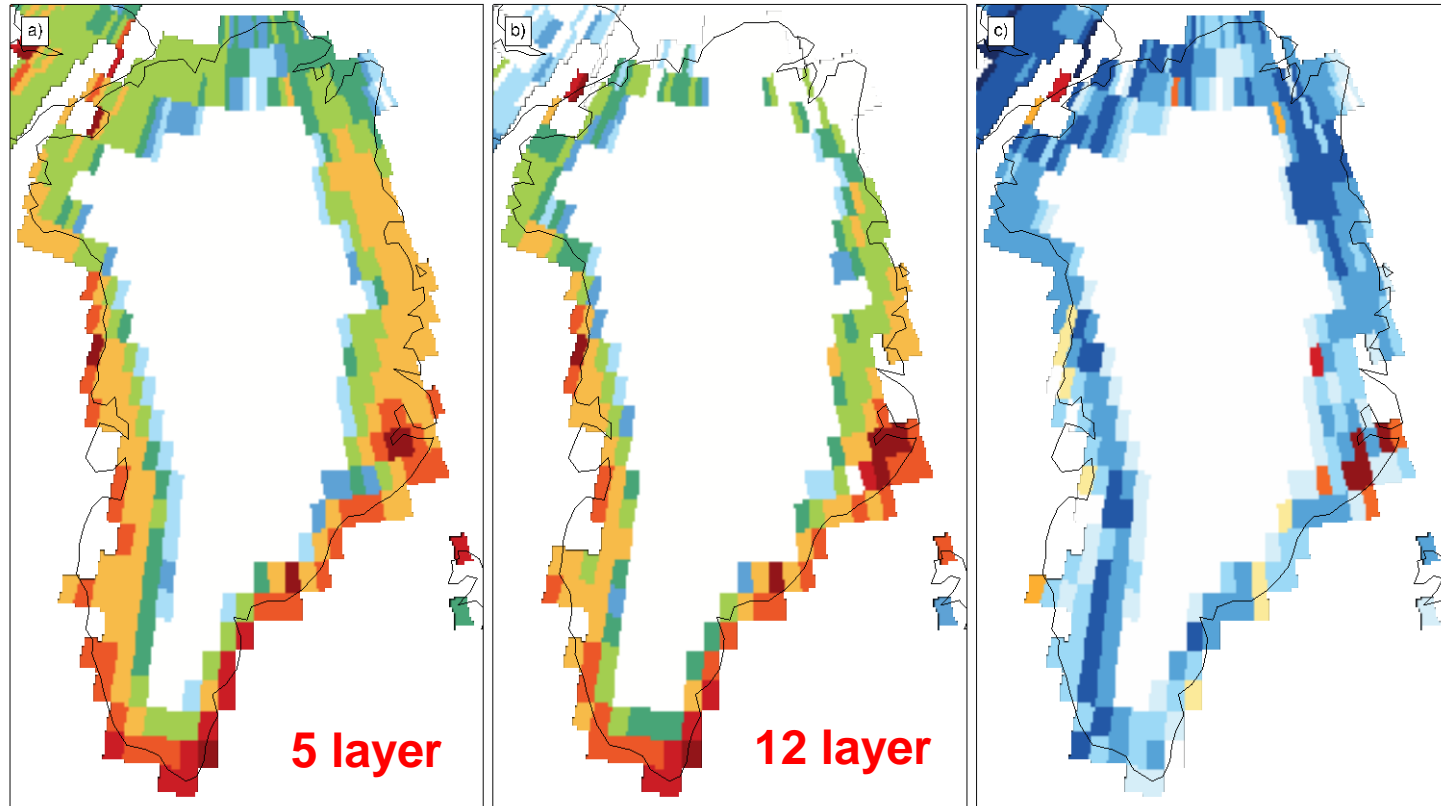
10 meter snowpack

Fraction of meltwater to runoff

tr_eraimod_5_cap

clm117_eraimod_default

Right - Left



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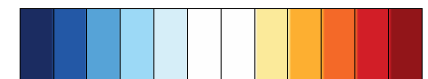
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1 5 10 20 40 60 80 100



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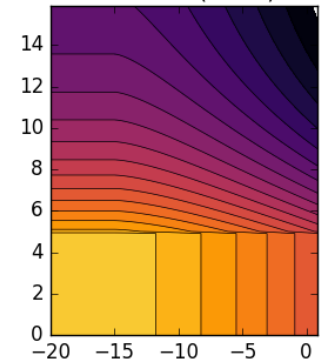
-50 -20 -10 -5 -1 0 1 5 10 20 50

Offline CLM

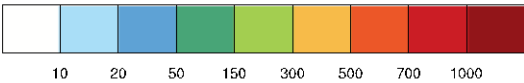
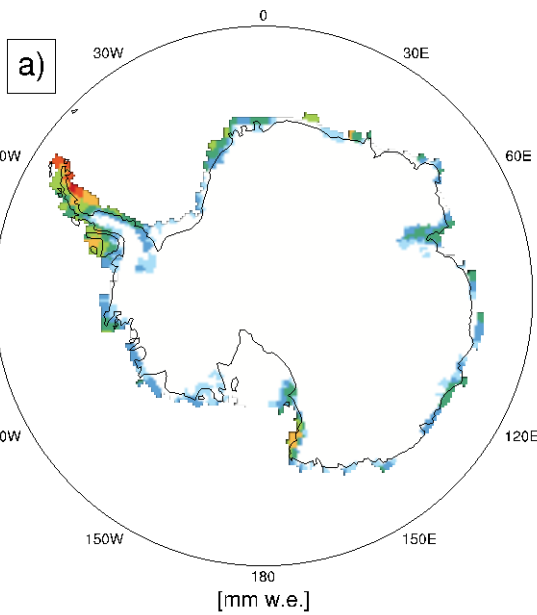
ERA-Interim forcing
1979-2014

Spinup: 96 years

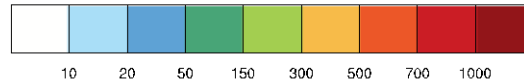
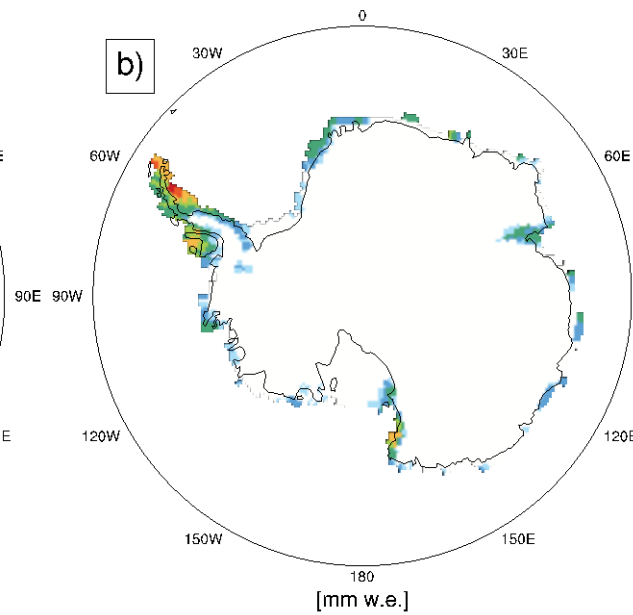
Wind-driven compaction



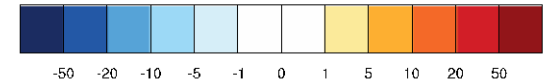
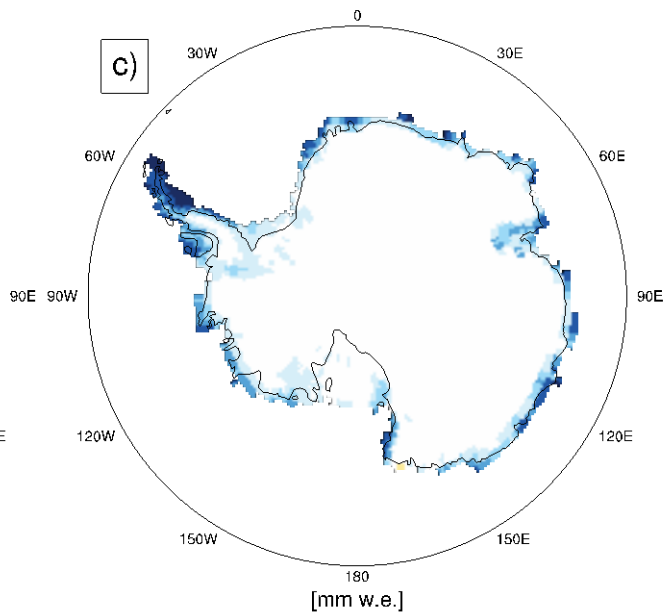
clm117_eraimod_default



clm117_eraimod_wnd



Right - Left



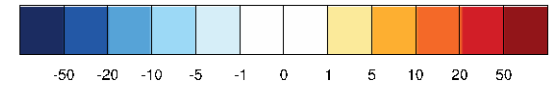
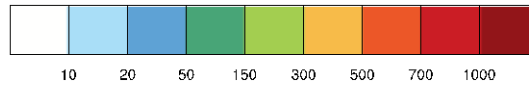
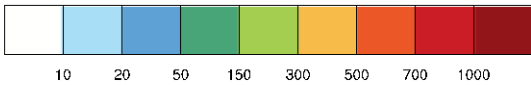
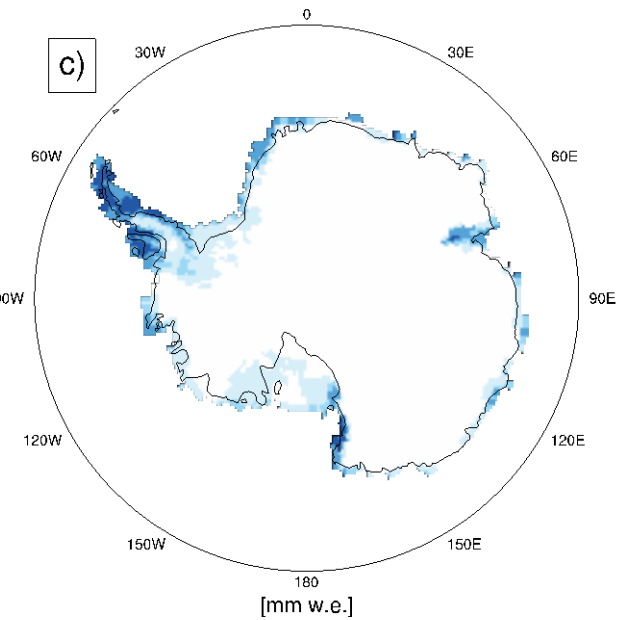
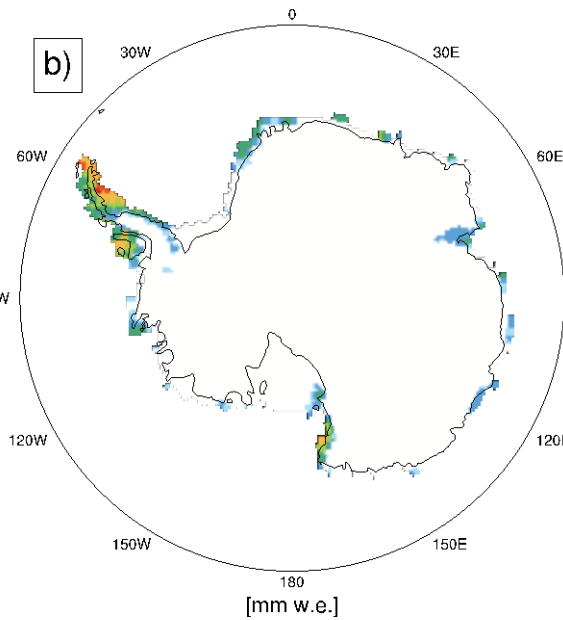
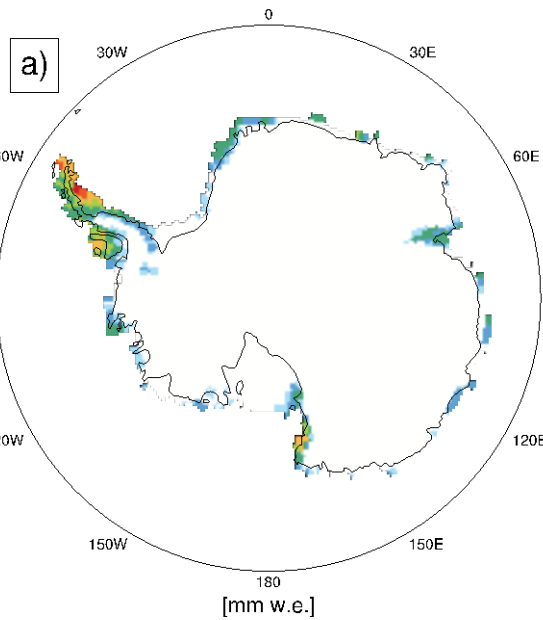
Melt

Destructive metamorphism

clm117_eraimod_wnd

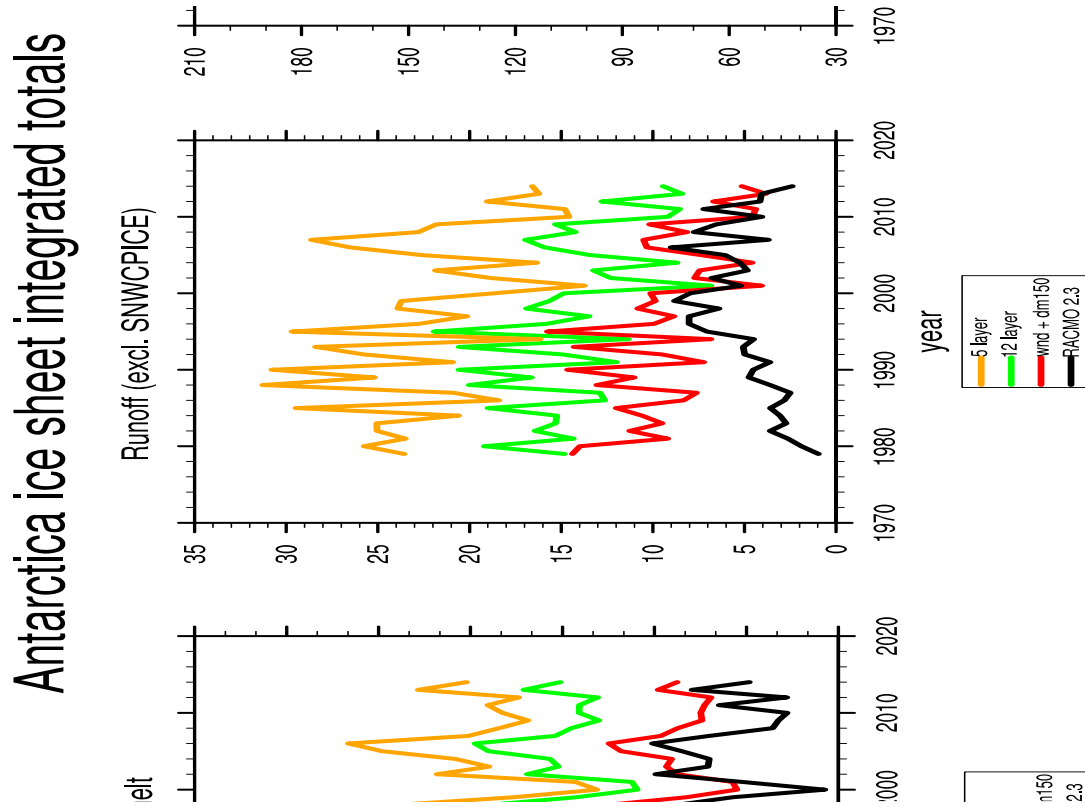
clm117_eraimod_wnd_dm150

Right - Left



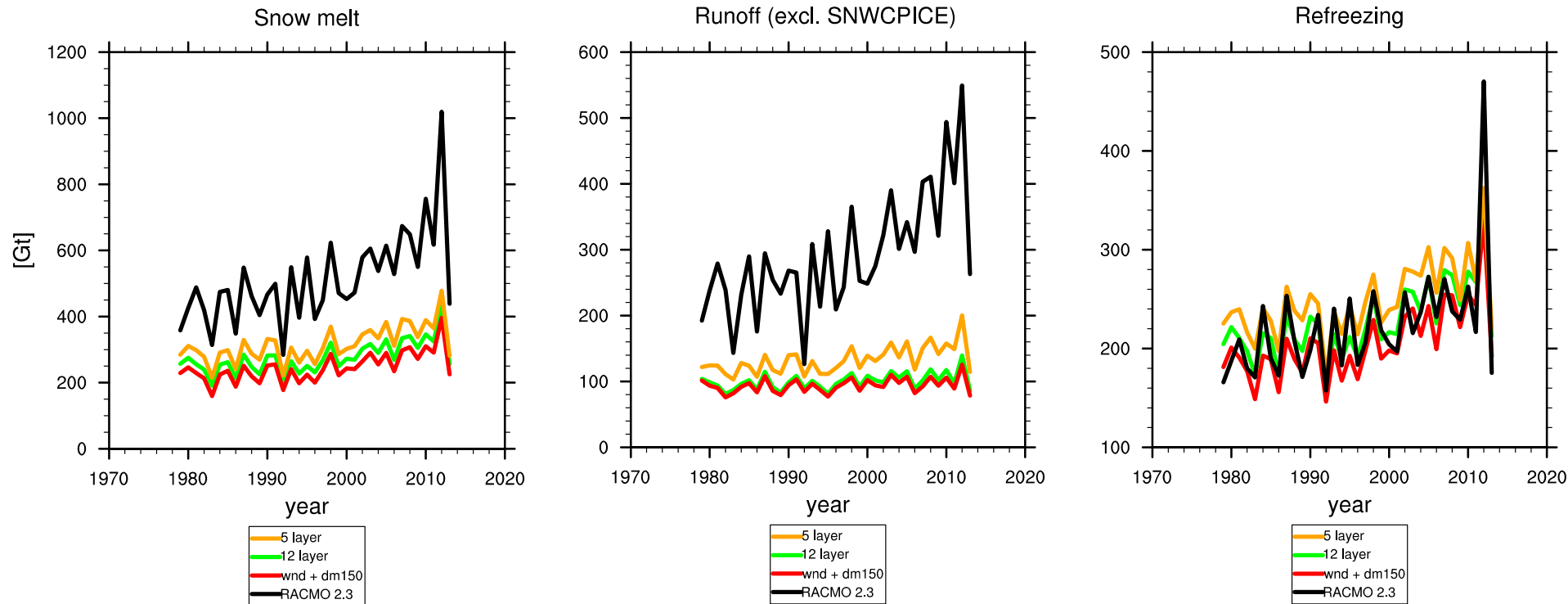
Melt

AIS integrated totals



- Increases in density help resolve the Antarctic (subsurface) melt bias

GIS integrated totals



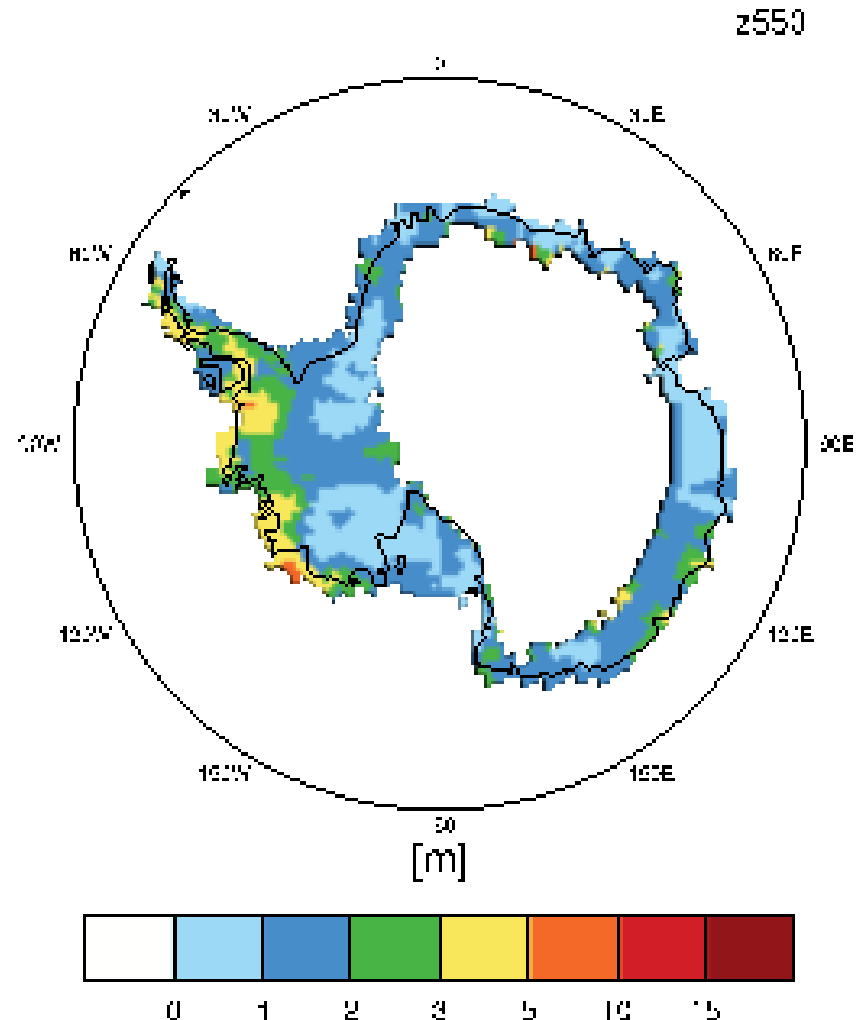
- Melt on Greenland is underestimated, but arguably not made worse by the changes

Something is not right...

z550: depth at which density of 550 kg/m³ is reached

On Antarctica, typically between 7-25 meters, with highest values in cold interior and high precip areas

Culprit: overburden compaction scheme



Compaction by overburden

$$\frac{d\rho}{dt} = \frac{-\sigma}{\eta}$$

CLM:

$$\eta = \eta_0 * \exp(a(T_{freeze} - T) + b\rho)$$

CLM 4.5 Tech Note

Crocus:

$$\eta = f_1 f_2 \eta_0 \frac{\rho}{c} * \exp(a(T_{freeze} - T) + b\rho)$$

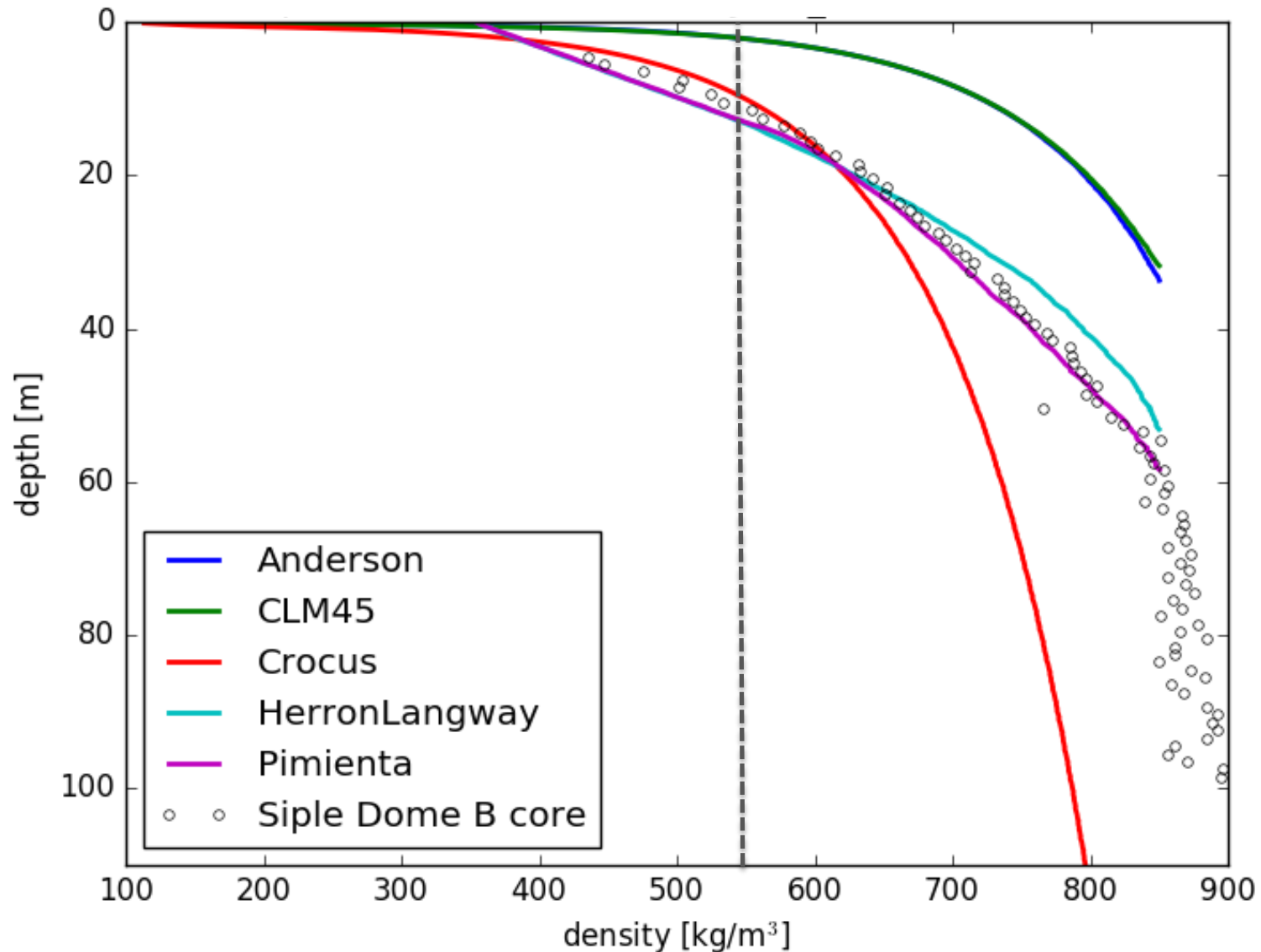
Vionnet et al, 2012

Testing in a dry firn model (DFM)

- Single column model, 10cm vertical resolution
- Explicit / implicit heat equation solver
- Different compaction formulations (steady-state, overburden)
- Forced by 6-hourly RACMO2.3 meteorology (T, acc, U)

<https://github.com/kampe004/compaction>

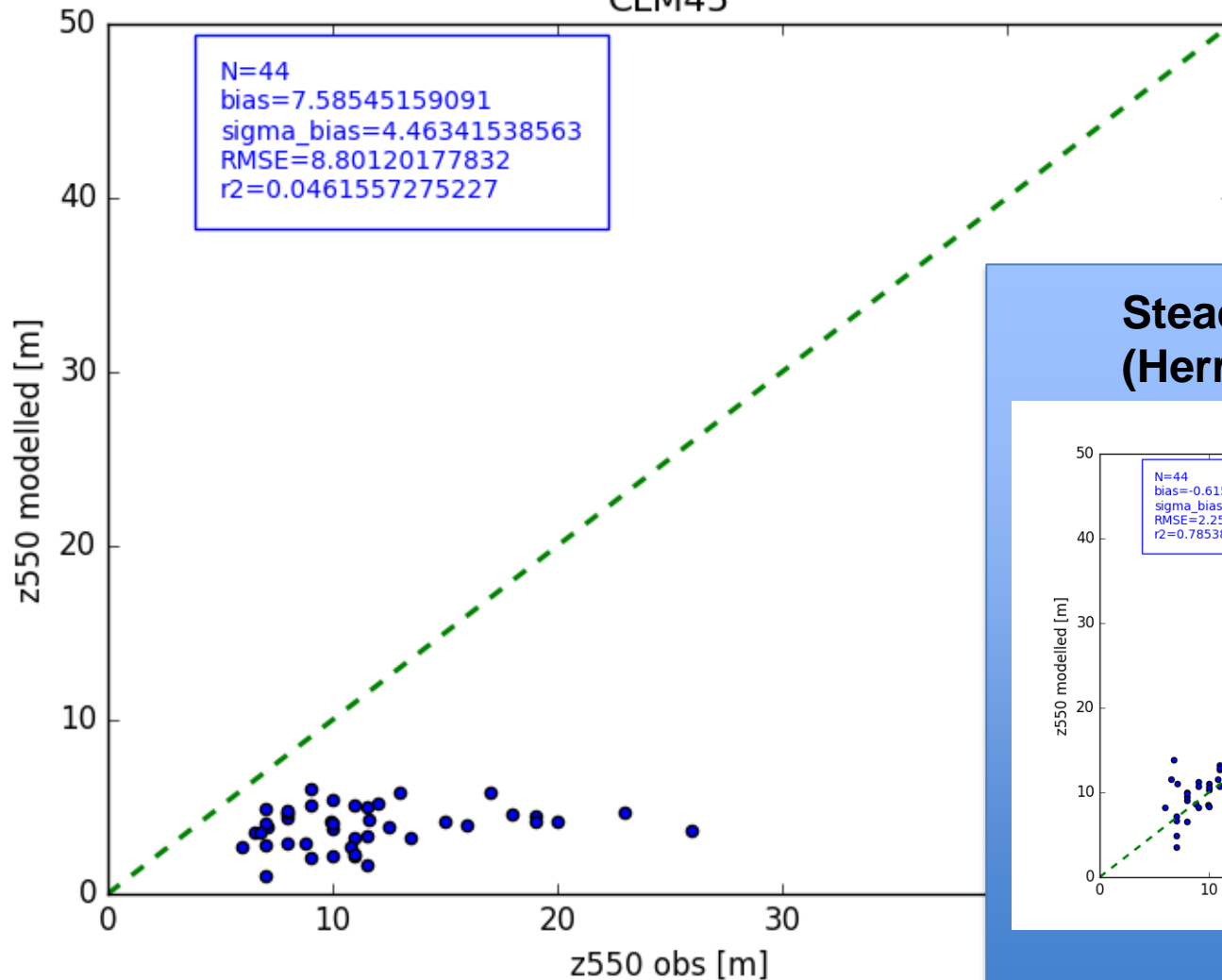
Siple Dome, Antarctica



Modelling z550

Antarctic firn cores: obs vs model

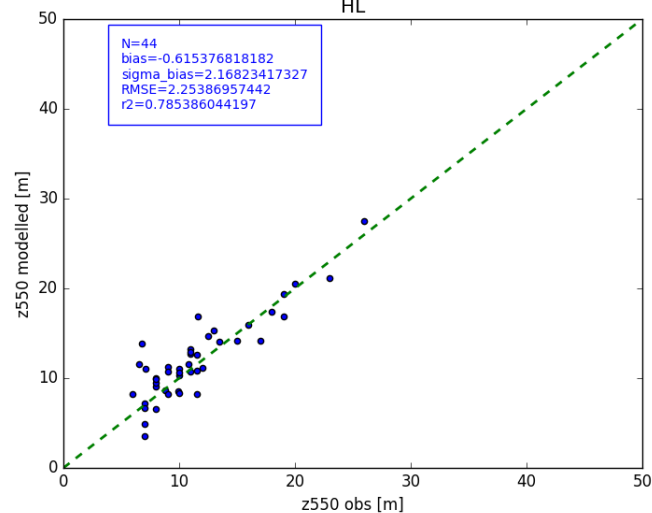
CLM45



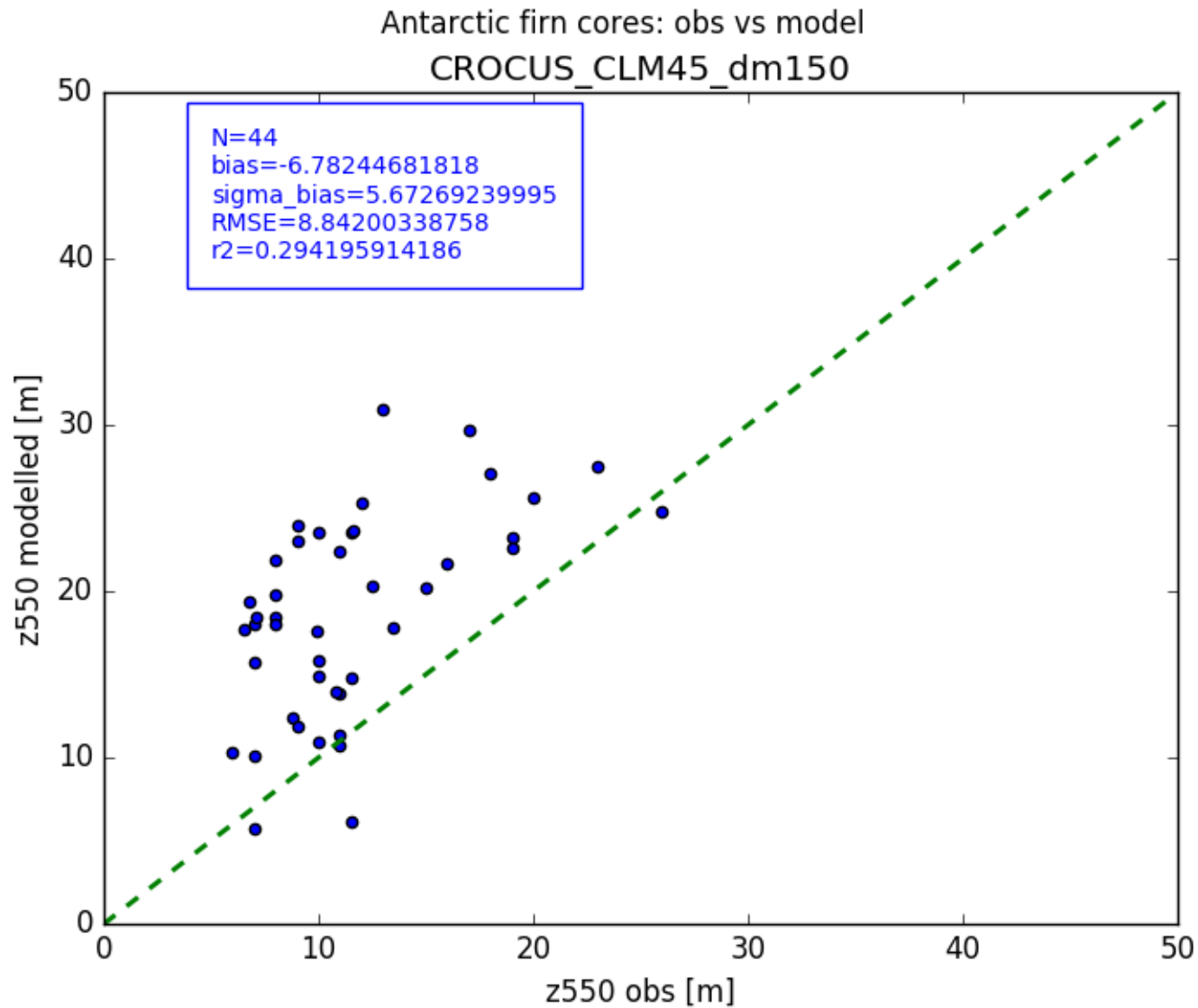
Steady state model (Herron & Langway)

Antarctic firn cores: obs vs model

HL

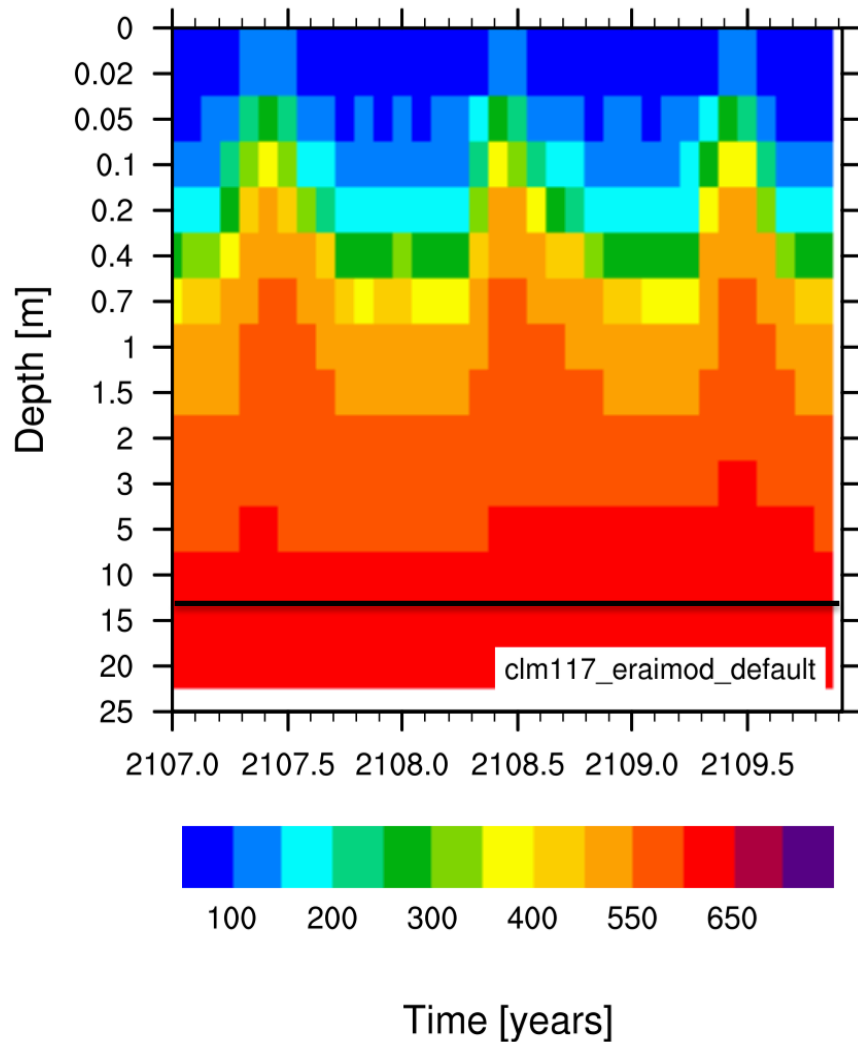


Modelling z550 (2)

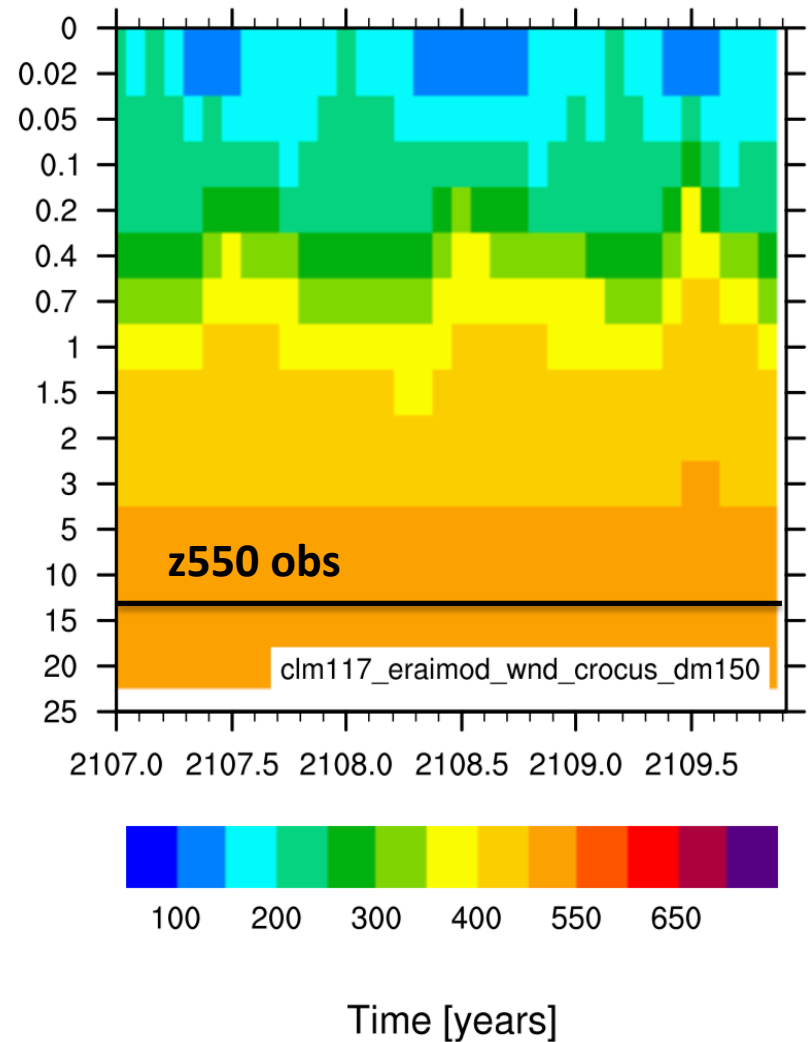


Summit, Greenland

CLM (Anderson 1979)



CROCUS (Brun 1989)



Future plans

- Introduce new overburden compaction parameters + wind drift parametrization
- Water on ice → effect on energy fluxes
- Increase vertical resolution and layering based on physical properties (creates potential for impermeable layers and isotope tracking)

THANK YOU