Announced title: Elevation classes for SMB calculations

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NEW TITLE

Coupled climate and Greenland ice sheet evolution up to A.D. 2300 simulated with the MPI model

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Motivation

- ISMIP6
- Ice sheet sensitivity to climate variability
- Propagation of SMB biases in a coupled model
- Processes that can be investigated with a coupled model

Model



ISMIP6 coupled experiments

Table 1: Summary of the ISMIP6 experiments.

Experiment Title	CMIP6 Label (experiment id)	Experiment Description	Start Year	End Year	Minimum # Years Per Simulation	Major Purposes						
DECK Experiments												
AMIP	amip ism-amip-std	observed SSTs and SICs prescribed 1979 2014 36 evaluation, va		evaluation, variability								
pre-industrial or present day control	piControl piControl-withism ism-piControl-self ism-pdControl-std	pre-industrial (pi) conditions imposed with CO ₂ concentration fixed, or present day (pd) condition imposed	n/a	n/a	<mark>(500</mark>)	evaluation, unforced variability						
abrupt quadrupling of CO2	abrupt-4xCO2	atmospheric CO ₂ concentration abruptly quadrupled and then held constant	n/a	n/a	150	climate sensitivity, feedbacks, fast responses						
1% per year CO ₂ increase	1pctCO2 1pctCO2-withism ism-1pctCO2-self ism-1pctCO2-std	atmospheric CO_2 concentration prescribed to increase at 1% yr ¹ and then held constant to quadruple levels	n/a	n/a	<mark>(150</mark>)	climate sensitivity, feedbacks, idealized benchmark						
CMIP6 historical Simulations												
CMIP6 historical	historical historical-withism ism-historical-self ism-historical-std	simulation of the recent past with CO2 concentration prescribed	1850	2014	165	evaluation						
		CMIP6 ScenarioMIP S	imulatio	15								
ScenarioMIP	ssp5-8.5 ssp5-8.5-withism ism-ssp5-8.5-self ism-ssp5-8.5-std	future scenario with high radiative forcing by the end of the century	2015	2300	286	climate sensitivity						
CMIP6 PMIP Simulations												
PMIP last interglacial	lig127k ism-lig127k -std	simulation of the last interglacial	n/a	n/a	3000	climate sensitivity, feedbacks, long responses						

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Nowicki et al., GMDD

Holocene results



- From 9 to 3.5 ky BP, volume decreases in response to high insolation
- Without human GHGs, the modeled present-day GrIS gains mass

Present-day topography

Simulated 1960-2005 thickness and bias w.r.t. observations (Bamber et al. 2001)



- Total **area** and **volume** are **overestimated**
- NW margin too far from coast

Contours:

- Simulated and observed topography
- Modeled and observed (white) margins

SMB compares relatively well with high-resolution regional climate model

4000



- kg m⁻² yr⁻¹ Differences with RCM are due to atmospheric and topographic biases (* isolates atmospheric bias)
 - NW thickness biases are caused by insufficient accumulation

Three scenarios



Vizcaino et al., GRL, 2015

- 3 scenarios: RCP3.0, RCP4.5, RCP8.5 and their extensions (ECP's)
- In **RCP2.6**, T peaks in 21st century, then declines
- 1, 3 and 9 K warming by
 2300

Volume and SMB evolution



Gt a ⁻¹		RCP/ECP 2.6		RCP/ECP4.5		RCP/ECP8.5_4x	
	1980- 1999	2080- 2099	2280 - 2299	2080- 2099	2280- 2299	2080- 2099	2280- 2290
SMB	288	88	149	-5	-145	-313	-2132
Accu	707	+7%		+9%		+18%	
RU	419	+60%		+84%		+273%	

- 37, 44 and 68
 (2100) and 68, 135
 and 539 (2300) mm
 global-mean SLR
- 10-30% less mass loss if SMB
 - elevation feedback
 - is not included (*)
- SMB becomes negative by 2100 in RCP4.5 and RCP8.5

Variability in the simulated 1990-2013 GrIS mass trend



- Net **mass loss** in the three ensemble members
- Large spread: 0.06, 0.08 &
 0.15 mm a⁻¹ SLR
 - Mass loss is **underestimated** (Shepherd et al. 2012: 0.39±0.13 mm a⁻¹ SLR) due to lack of ocean forcing

Small sensitivity of future volume loss to internal climate variability



- Approx. linear increase of sigma with the mean
- Much less variation within ensemble than for different scenarios

Future SMB

1960-2005



- ELA increases in all simulations, particularly in the northern half
- SMB increases over N & SE accumulation areas due to increased snowfall

No "pinning-point"



Eemian GrIS as simulated with RACMO-ANICE (Helsen et al., 2013)

Sources of elevation change

2080-99 minus 1980-99, under RCP8.5:



- Gross pattern largely given by cumulative SMB
- Increased surface slope from cumulative SMB enhances ice flow
 - This partially counterbalances elevation reduction from enhanced melt

Elevation-SMB feedback

SMB anomaly associated with elevation change 2080-99 minus 1980-99 under **RCP8.5**



Vizcaino et al., GRL, 2015

- Elevation change affects
 melt, precipitation, and rain
 fraction
- Feedback estimated from simulations where SMB is calculated at fixed 1980-1999 ISM topography
- Feedback reduces SMB over most of the ice sheet
 - Melt increases at reduced elevations
 - Accumulation is reduced at increased elevations
- Large spatial gradients, as in *Helsen et al., [2012]*

Ice discharge evolution



Simulated ice discharge in Gt a⁻¹ per 10⁴ km² and GrIS margin (green contour) :



- Decreases in all simulations due to marginal retreat and thinning caused by increased runoff (as in Lipscomb et al., J. Clim., 2013)
- However, ice discharge distribution is not realistically simulated (e.g., NW) & model does not reproduce current glacier acceleration

Conclusions

- **First** coupled simulation from GrIS model and AOGCM under RCP/ECP forcing
- Results suggest that GrIS would be growing in absence of anthropogenic climate change
- Surface mass balance, elevation and ice flow are **coupled**
- SMB-elevation feedback enhances mass loss by 10% (2100) and 30% (2300)
 - Limited validity of fixed-topo SMB projections beyond 2100
- Ensemble reveals
 - wide spread in simulated present-day mass trends
 - ➤Caution in evaluation of models with observations
 - small sensitivity to future climate variability compared with sensitivity to GHG scenario