

# SIMULATIONS OF PLIOCENE ARCTIC CLIMATE WITH PROGNOSTIC AEROSOL-CLOUD INTERACTIONS

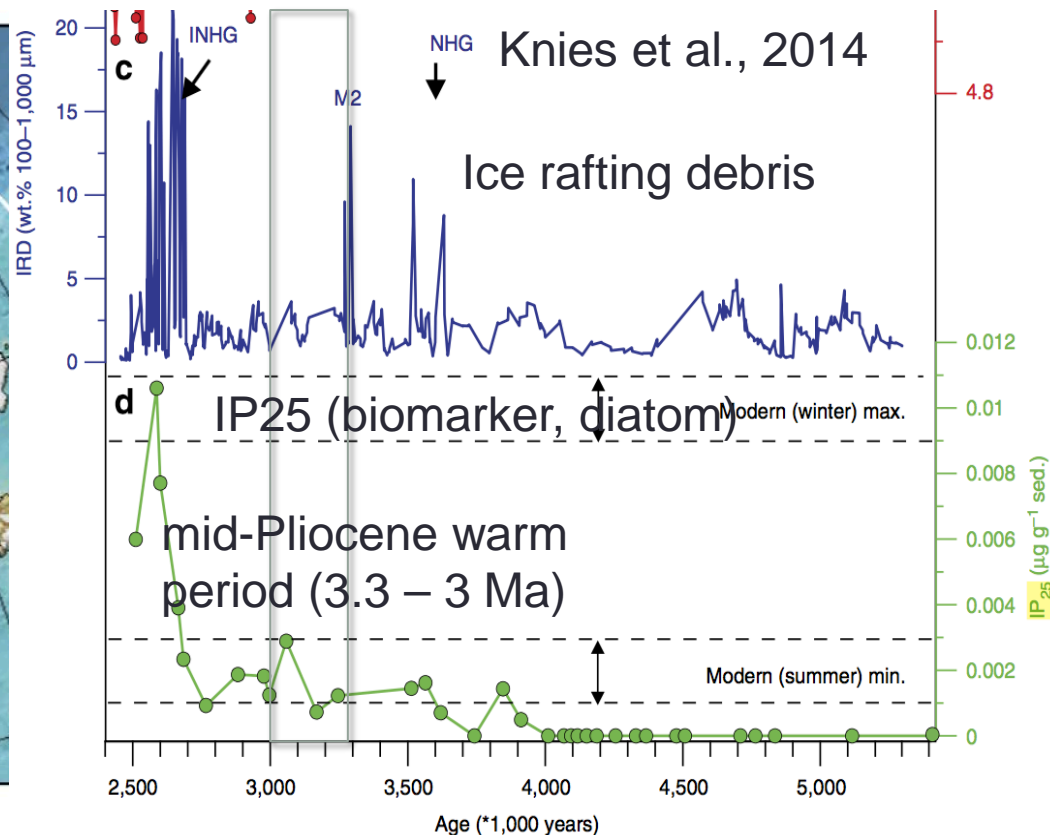
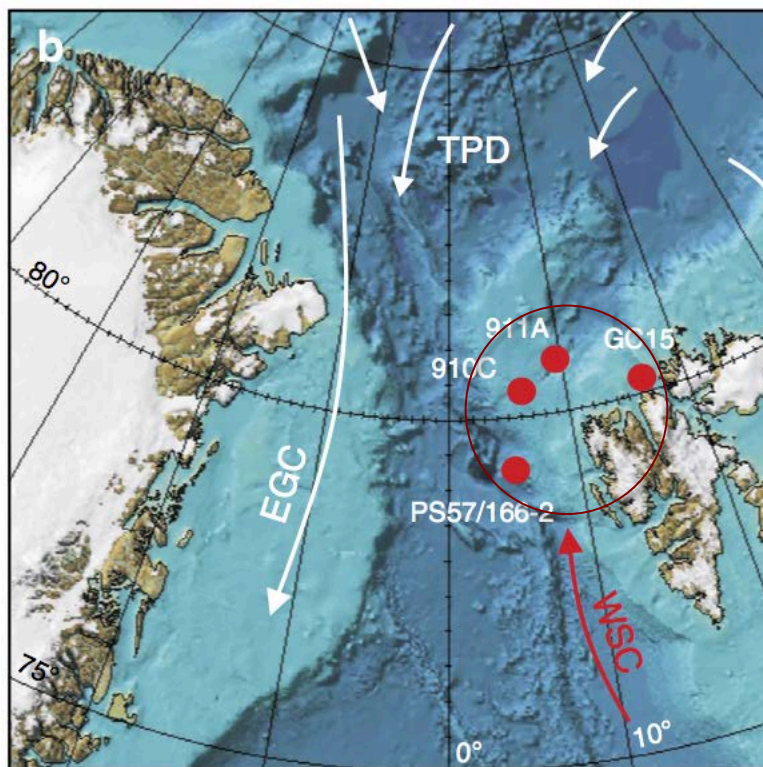
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Brady

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# Low Arctic sea ice and warm high latitude climate

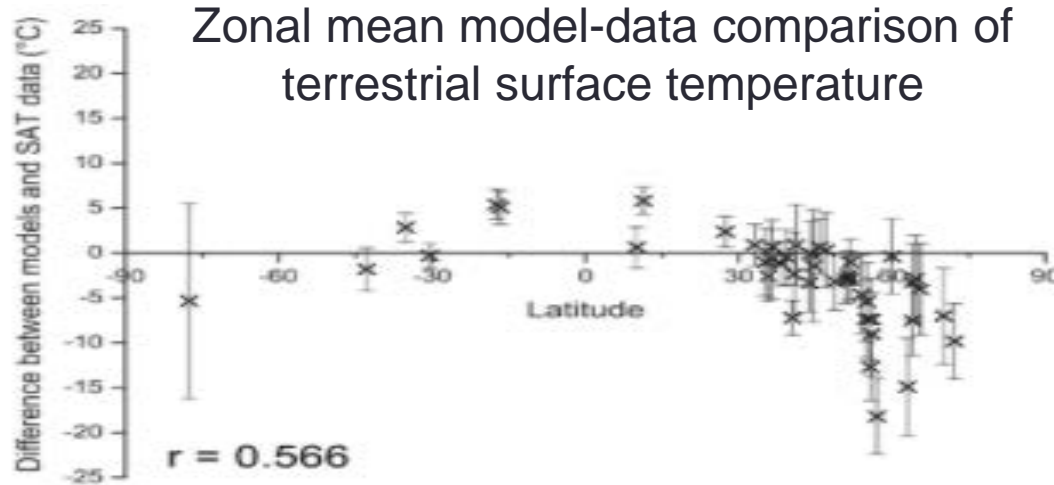
Pliocene: 2.6 – 5.3 Ma. High latitude temperatures: up to 18 °C warmer



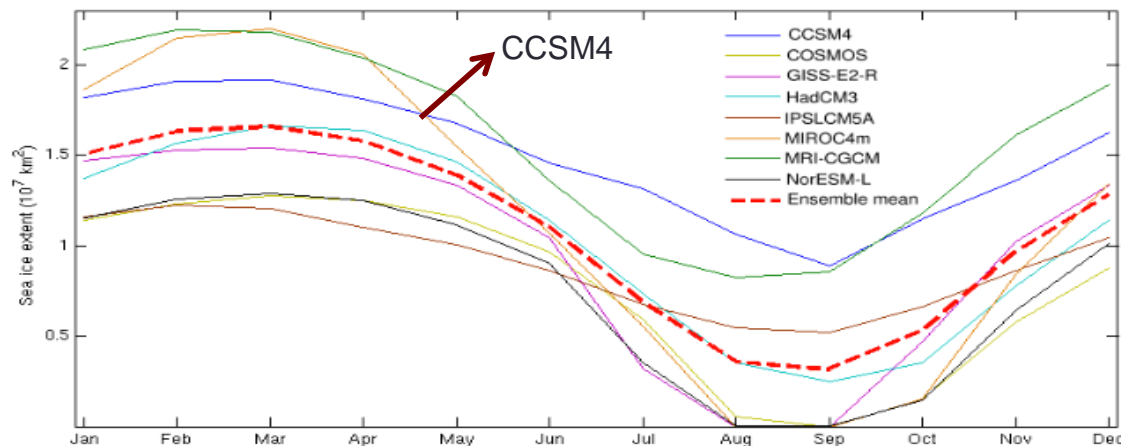
Significant Arctic sea ice reduction  $\rightarrow$  seasonal sea ice free, annual sea ice approaching present-day summer minimum sea ice extent

# Challenges for model simulations

- Too cold high latitudes and too much Arctic sea ice



## Arctic sea ice extent of PlioMIP I



# Hypotheses to resolve the underestimate of Pliocene high latitude warming

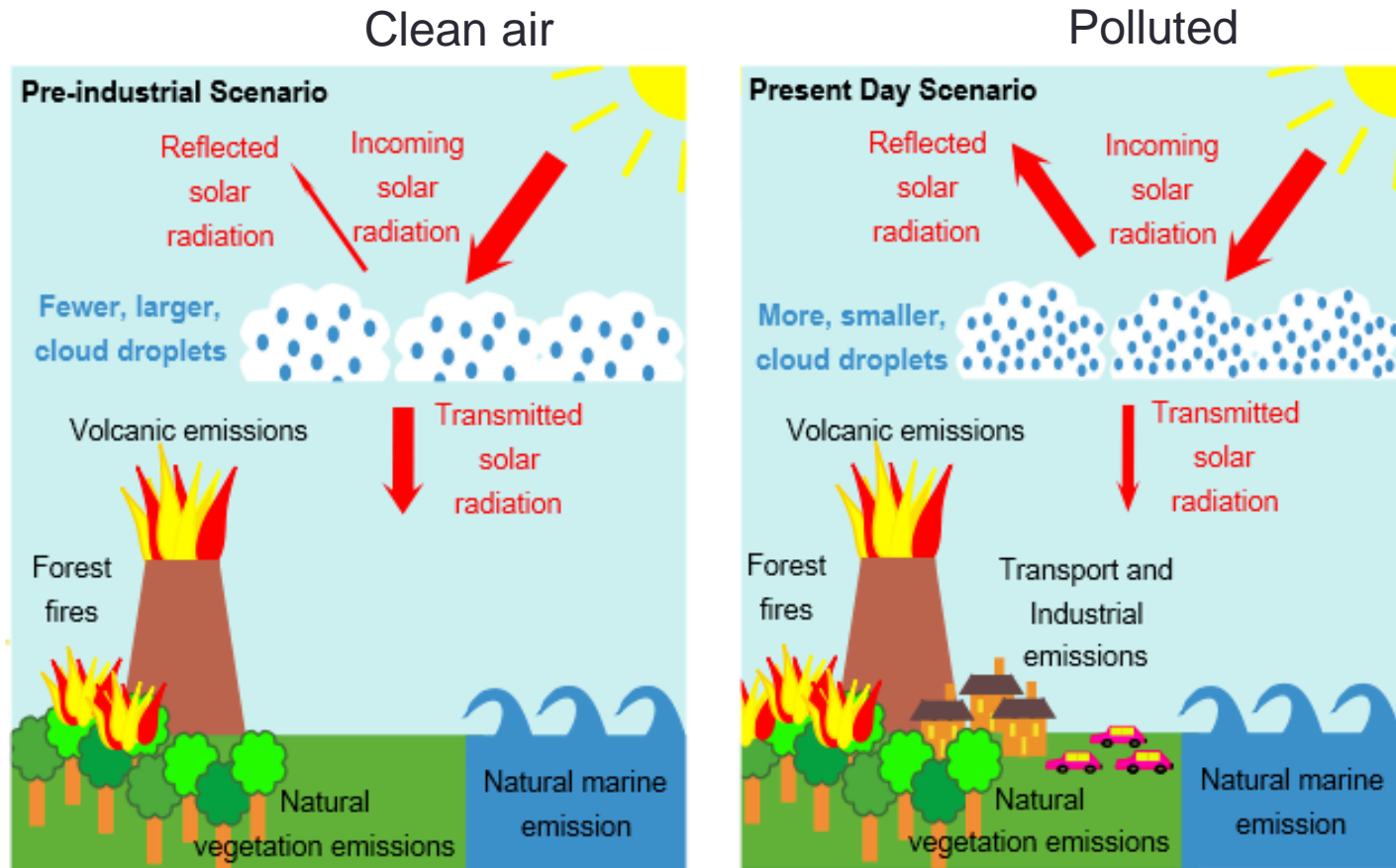
- 1. Variability of **CO<sub>2</sub>** and **orbital forcing** (Prescott et al., 2014; Haywood et al., 2016; Feng et al., 2016, *in prep*)
- 2. **Arctic gateway changes** (Otto-Bliesner et al., 2016, *in Prep*) and **regional geographic changes** (Hill et al., 2015)
- 3. **Other climate forcings and feedbacks**
  - **Sea ice albedo** (HadCM3, Howell et al., 2016)
  - **Changing atmosphere chemistry in the past**
    - **Direct radiative effect** (Unger and Xu, 2014)
    - **Indirect effect – changes in cloud albedo and lifetime**
      - PETM (55 Ma), changing cloud condensation nuclei (Kiehl and Shields, 2013)

# Outline

- Aerosol indirect effect
- Model simulations:
  - **Pliocene Arctic: an equilibrium state of clean atmosphere?**
- Changes in energy budget, sea ice and clouds

# Indirect effect

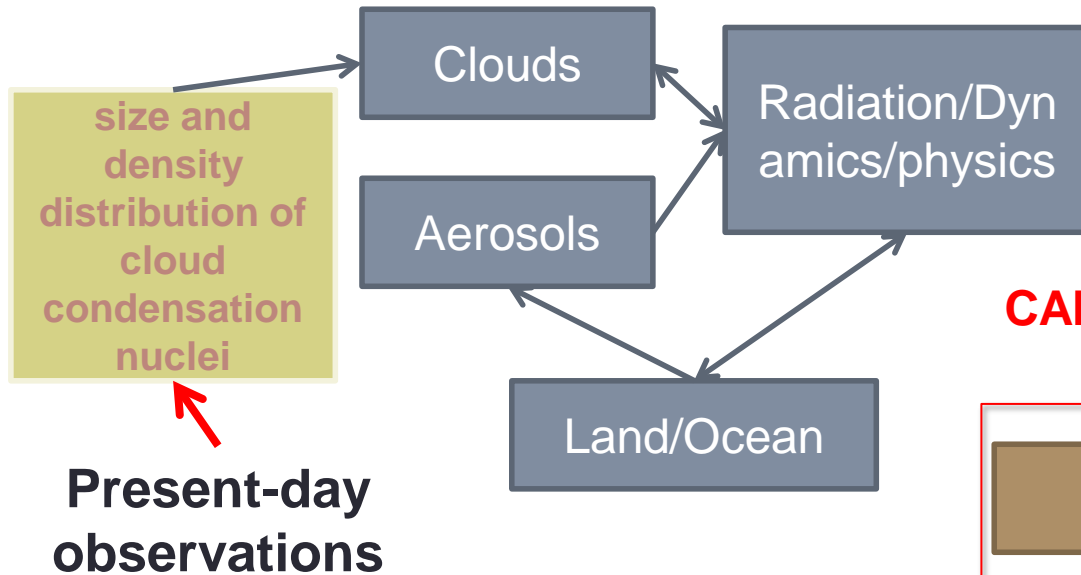
- Cloud albedo
- Cloud lifetime



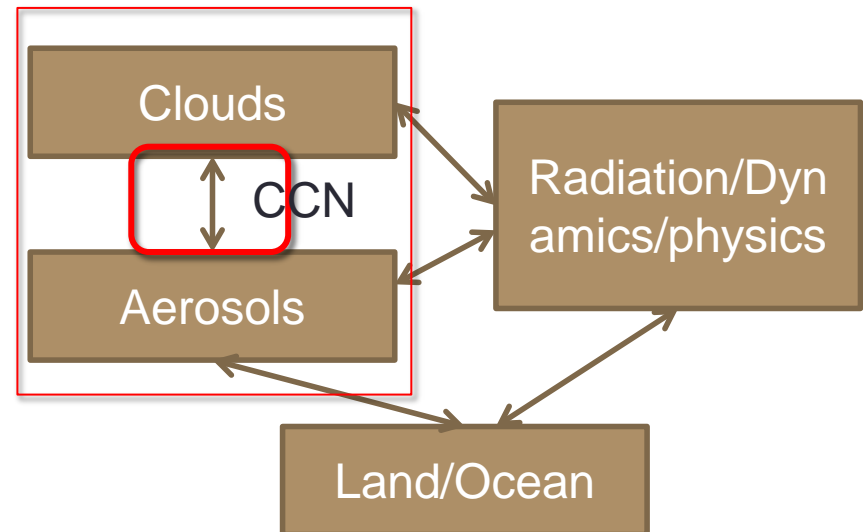
IPCC AR5: total radiative forcing due to changes in aerosols since 1750 is  $0.7 \text{ W/m}^2$  (compare to  $1.68 \text{ W/m}^2 \text{ CO}_2$  forcing)

# A clean Pliocene atmosphere

**CAM4: only direct radiative effect of aerosol**



**CAM5 and later: aerosol-cloud interaction**



# Experiment Design

CCSM4 (CAM4)  
PlioMIP I  
(Rosenbloom et al., 2013)

Slab ocean  
(CAM5.3-  
MAM3) with  
fixed ocean  
heat flux

	Slb-clean	Slb-polluted	HCO <sub>2</sub> -clean	LCO <sub>2</sub> -clean	HCO <sub>2</sub> -polluted	LCO <sub>2</sub> -polluted
CO <sub>2</sub>	405 ppm	405 ppm	560 ppm	280 ppm	560 ppm	280 ppm
Emission	Clean: 18 50 emission	Polluted: 1850: natural emission, 2000: industrial emission	clean	clean	polluted	polluted

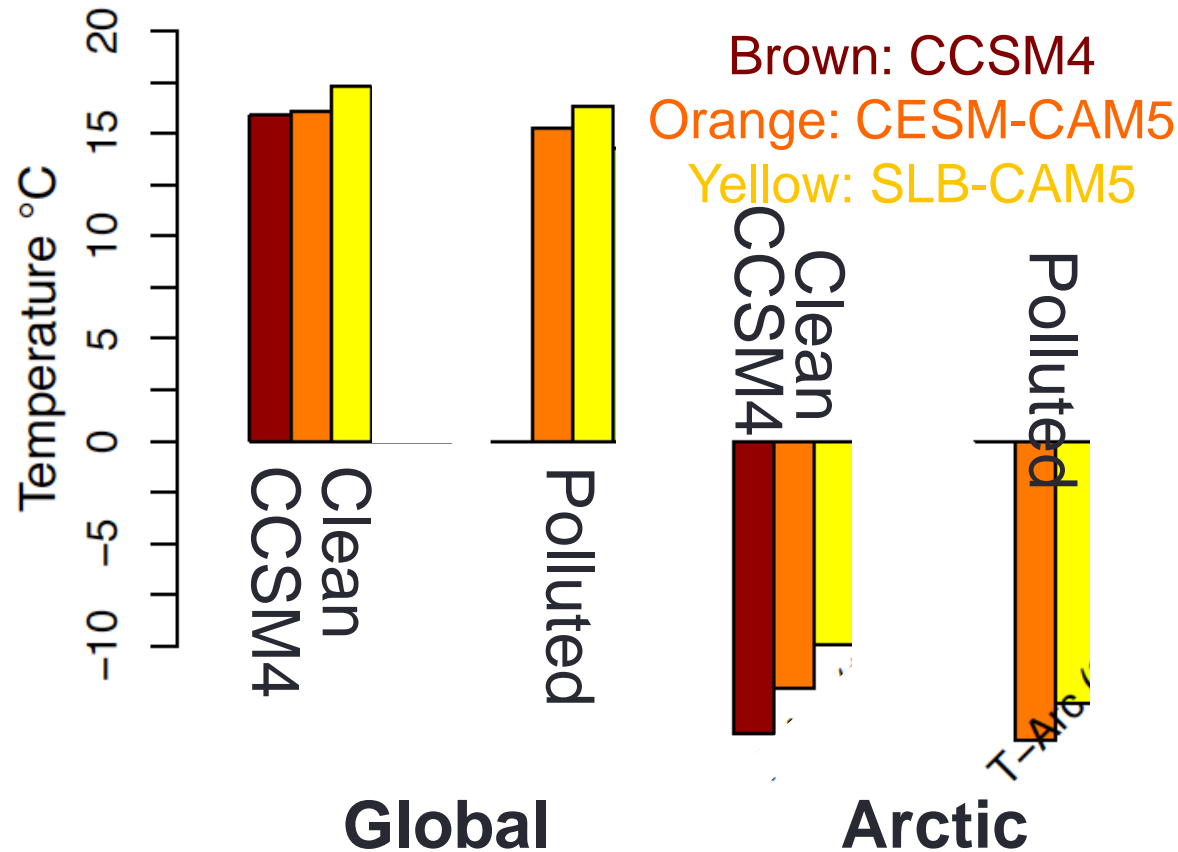
CESM1.2  
(CAM5.3-  
MAM3)

0.9 x 1.25 ° atmosphere, 320x384 grid cells ocean

	B-clean	B-polluted
Emission	clean	polluted
CO <sub>2</sub>	405 ppm	405 ppm

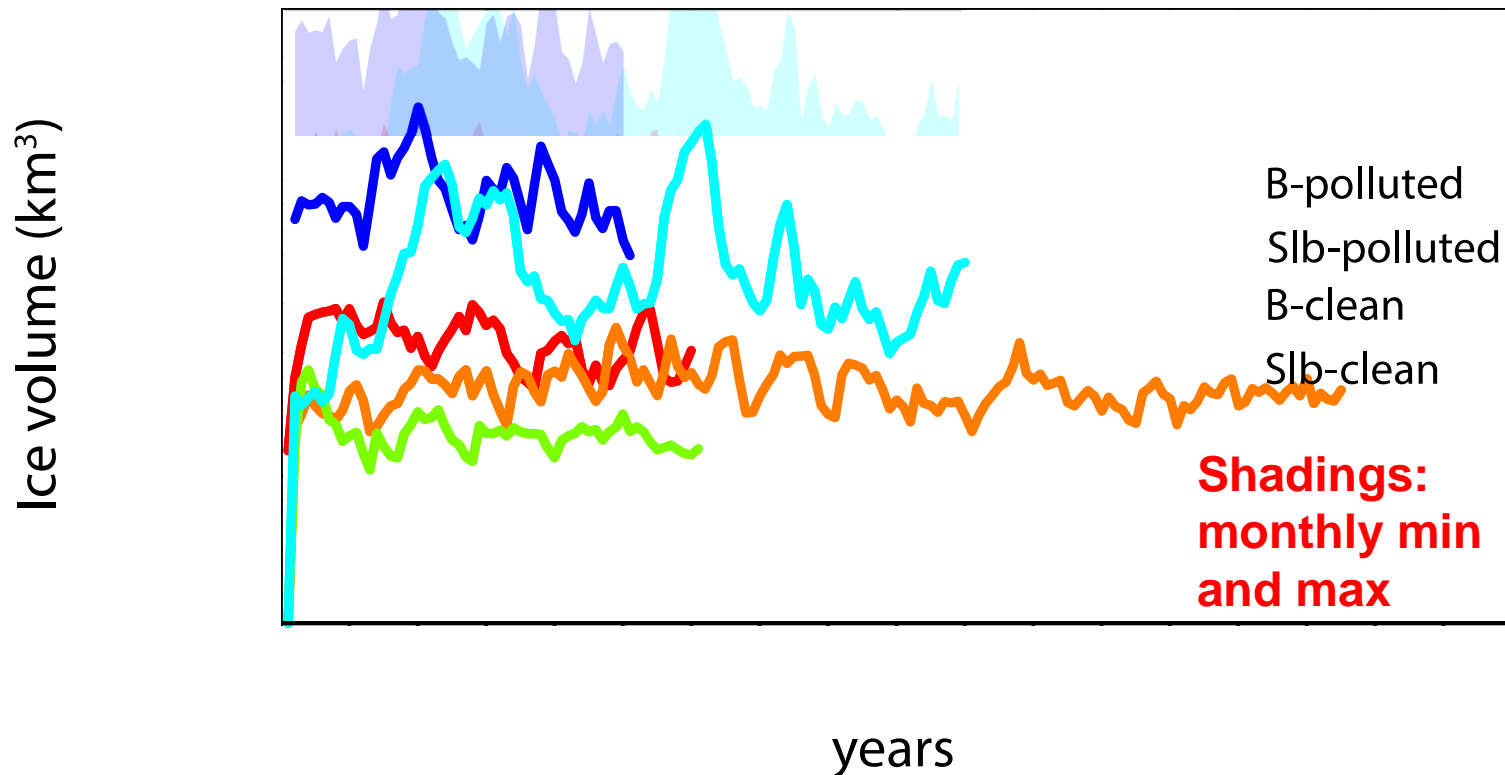


# Global & Arctic temperature responses



1.  $T_s$  CCSM4 is comparable to B-polluted.
2.  $\sim 0.8^\circ\text{C}$  warming of  $T_s$ -global, and  $2.6^\circ\text{C}$  warming of  $T_s$ -Arctic ( $70^\circ - 90^\circ\text{N}$  average).

# Arctic sea ice volume



- 46% annual sea ice reduction in B-1850, 23% reduction in B-2000. 86% reduction in monthly minimum in B-1850, 53% in B-2000.
- Large improvement in simulating Pliocene low ice state

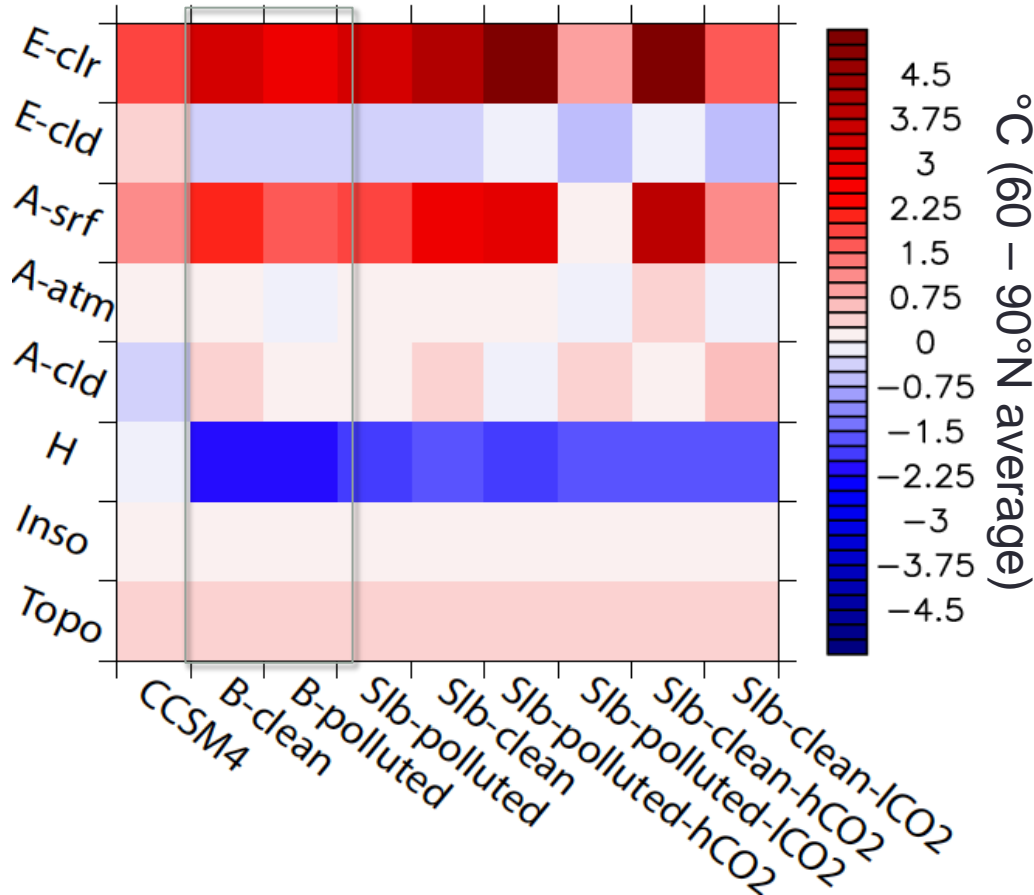
# Energy balance (refer to CCSM4 pre-industrial)

$$\Delta T = \Delta T(\Delta\alpha, \Delta\varepsilon, \Delta H, \xi) \text{ (Lunt et al., 2013; Hill et al., 2014)}$$

$$\Delta\alpha = \Delta\alpha_{\text{cld}} + \Delta\alpha_{\text{srf}} + \Delta\alpha_{\text{atm}} \text{ (APRP, Taylor, 2007) } (\alpha: \text{planetary albedo})$$

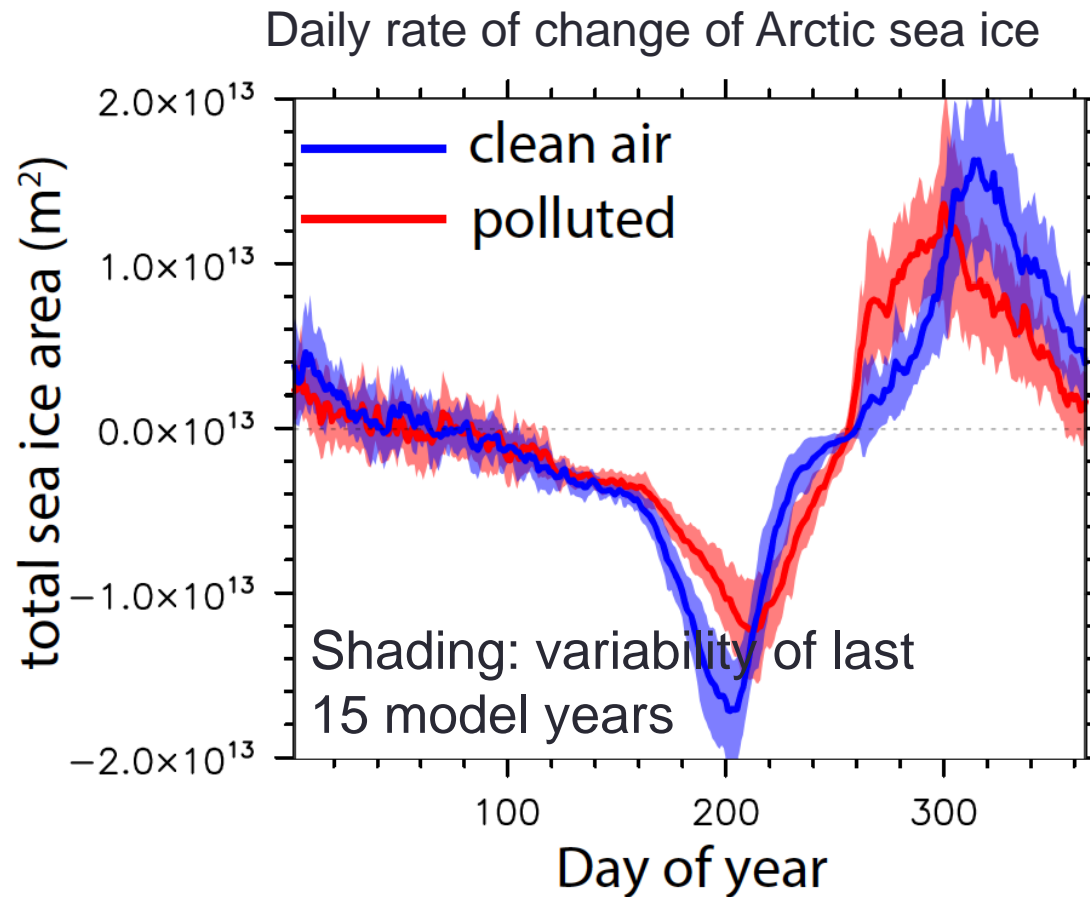
$$\Delta T - \varepsilon = \Delta T - \varepsilon_{\text{cldy}} + \Delta T - \varepsilon_{\text{clr}} \text{ } (\varepsilon: \text{Emissivity})$$

## $\Delta T$ Energy balance decomposition



- Warming:  $\Delta\alpha_{\text{srf}}$  and  $\Delta T - \varepsilon_{\text{clr}}$
- Cooling: total heat transport
- Mild warming effect from  $\Delta\alpha_{\text{cld}}$  (clean conditions) and cooling from  $\Delta T - \varepsilon_{\text{cldy}}$ .

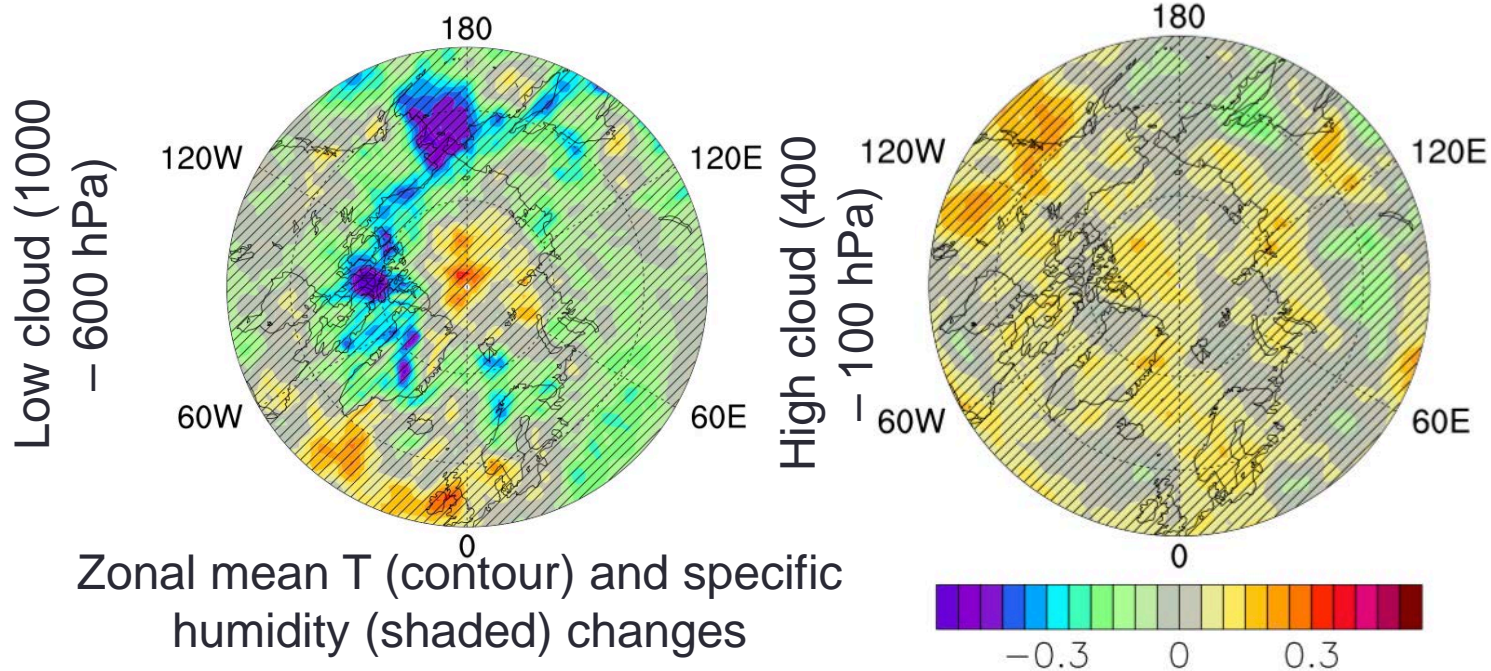
# Sea ice melting behavior



**Clean air: earlier and stronger melting, slower accumulation**

# Cloud contributions

July changes in cloud fraction (clean minus polluted)



Height  
(hPa)



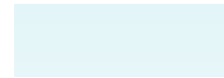
fading of shortwave cloud forcing (cooling effect) and increasing longwave cloud forcing (warming effect)

kg/kg

# Synthesis



Polluted regime:  
More CCN →  
more summer low clouds surrounding  
the Arctic →  
slower summer ice retreat → high  
surface albedo and less water vapor

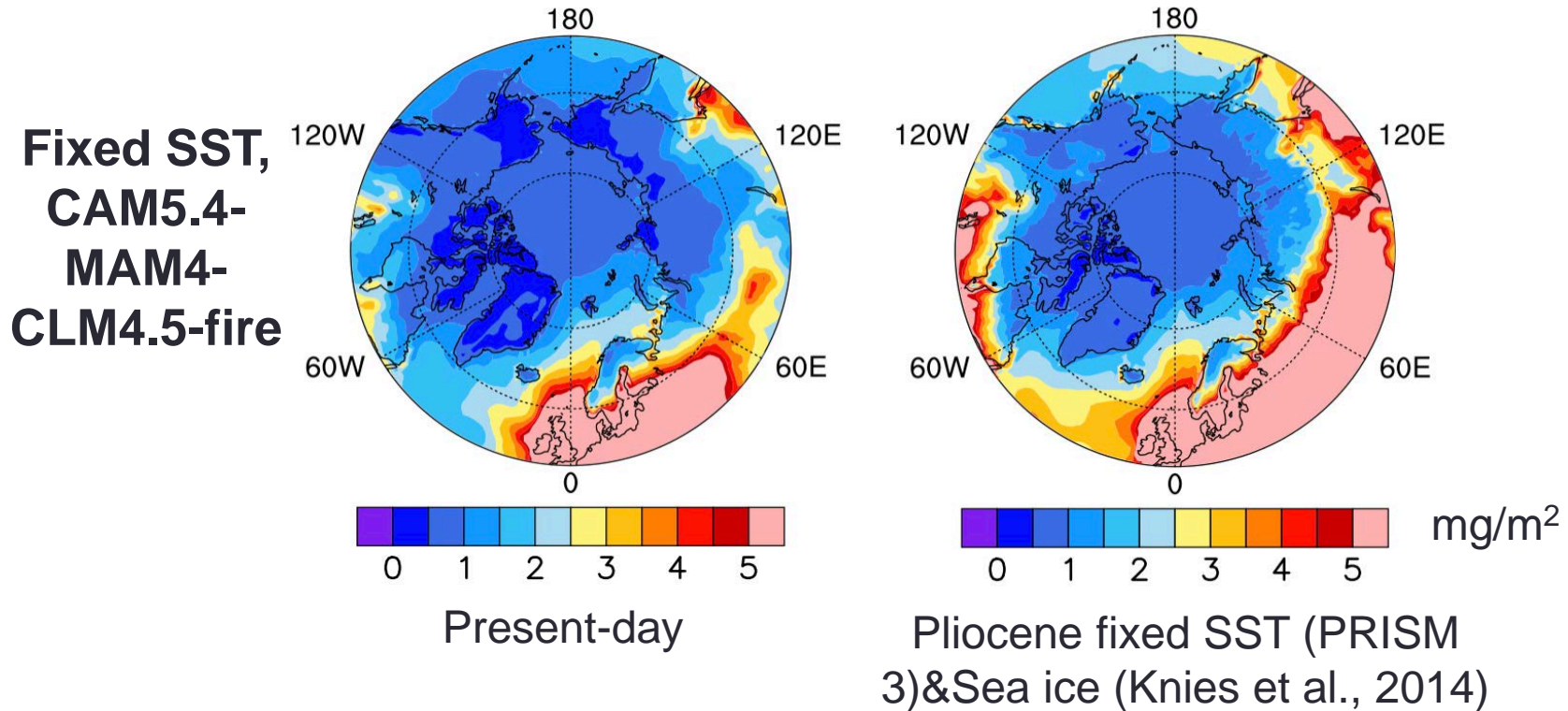


clean air regime:  
Less CCN →  
less summer low clouds and more  
high clouds surrounding the Arctic →  
rapid retreat → low surface albedo,  
more water vapor

# What happens next?

More realistic representation of Pliocene aerosol emissions  
**Fire & biogenic emission (really preliminary)**

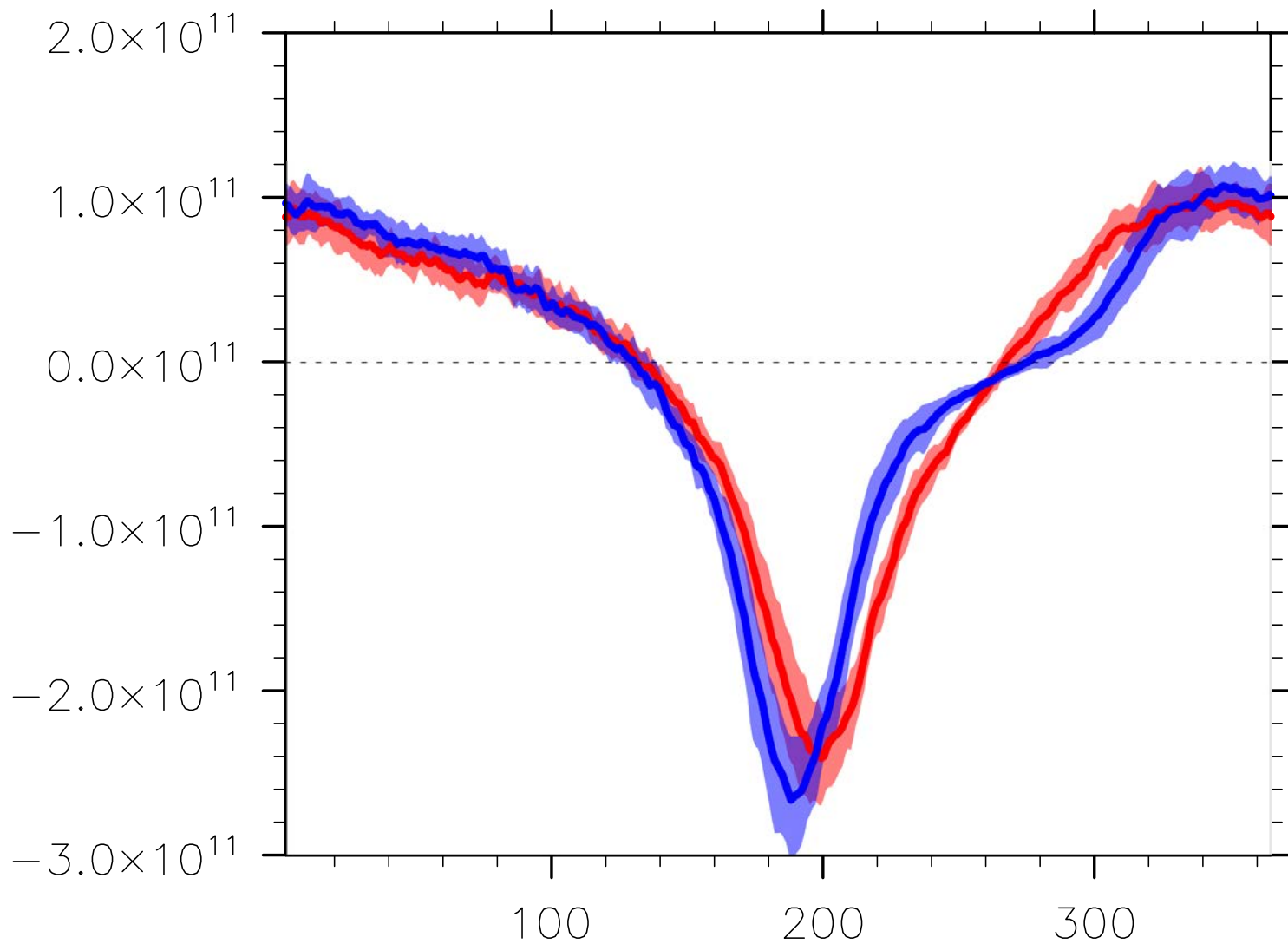
Annual wet + dry black carbon deposition on the surface



Thanks. Questions?

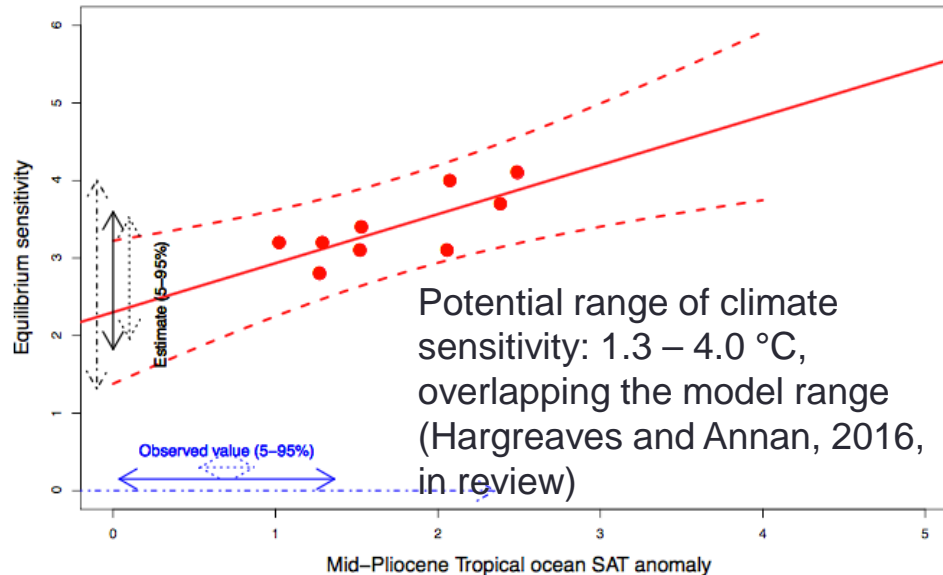


Daily rate of change of ice volume



# Pliocene Earth System Sensitivity vs Climate sensitivity

Moderate CS: comparable to proxy estimates



Underestimates ESS (7.1 – 9.6 °C/doubling of CO<sub>2</sub>): unlikely a CO<sub>2</sub> issue, changes in other components

(a) Experiment 2 Climate Models/Mean	(b) Pliocene ΔT (°C)	(c) CS (°C)	(d) ESS (°C) = mPWP ΔT · 1.88
CCSM4	<u>1.86</u>	3.2	<u>3.51</u>
COSMOS	3.60	4.1	6.77
GISS-E2-R	2.12	2.7	3.98
HADCM3	3.27	3.1	6.16
IPSLCM5A	2.18	3.4	4.10
MIROC4m	3.46	4.05	6.51
MRI-CGCM 2.3	1.84	3.2	3.45
NorESM-L	3.27	3.1	6.14
Ensemble Mean	2.66	3.36	5.01

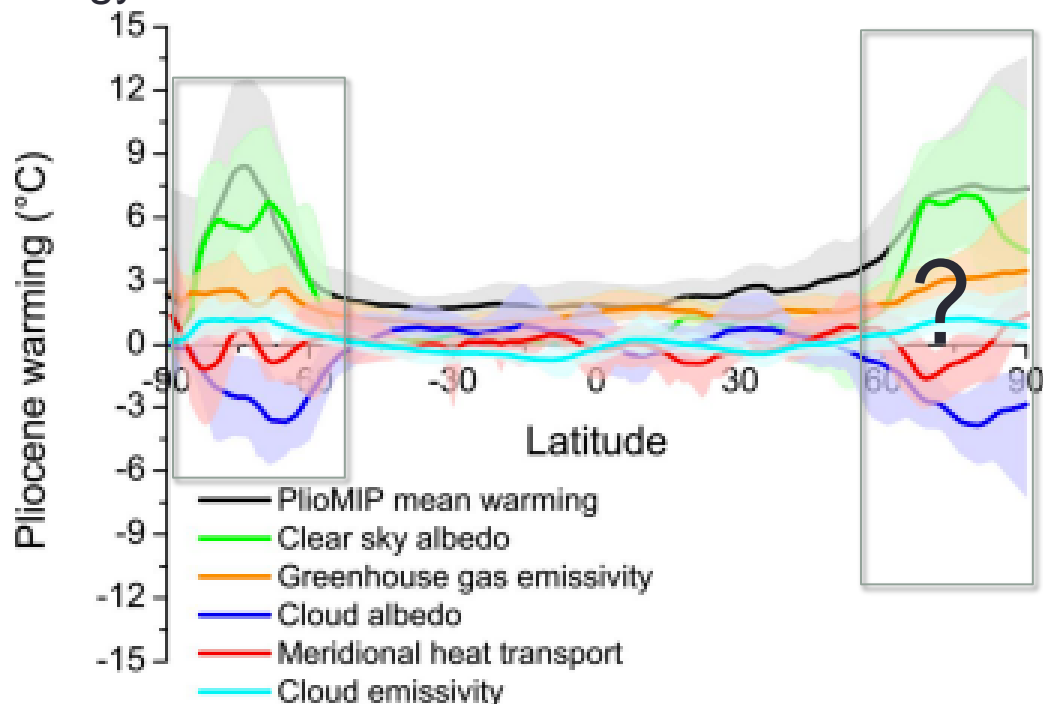
# Energy balance

$\Delta T = \Delta T(\Delta\alpha, \Delta\varepsilon, \Delta H) + \Delta T(\text{syn})$ ,  $\Delta\alpha = \Delta\alpha(\text{srf, cld, atm})$ ,  $\Delta\varepsilon = \Delta\varepsilon(\text{clr, cldy})$  (Lunt et al., 2013; Hill et al., 2014).

$\alpha\text{-cld: Rad (total) - Rad (clear)}$

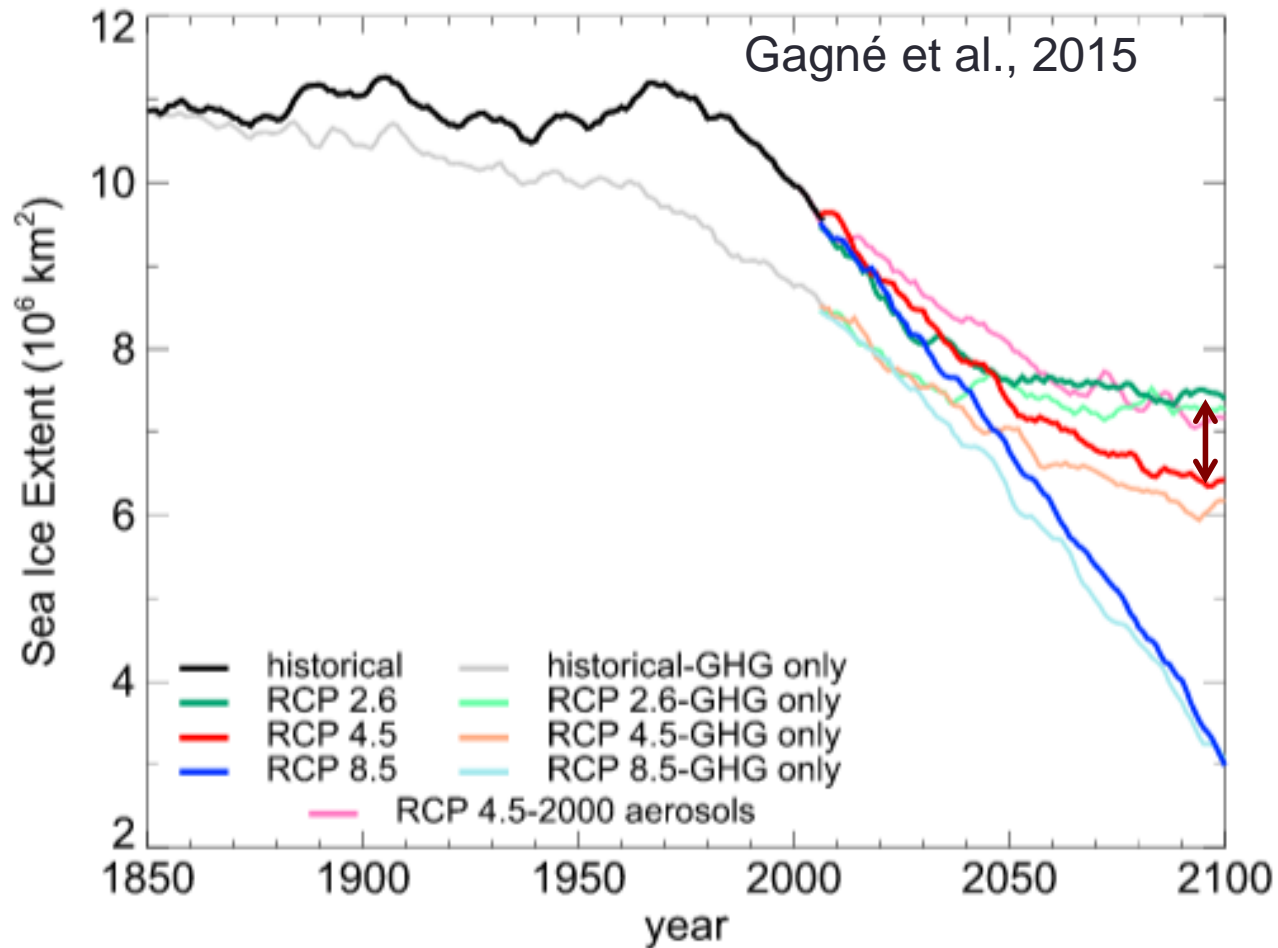
$\Delta\alpha\text{-cld} = \text{Rad-A (total)} - \text{Rad-A (clear)} - (\text{Rad-B (total)} - \text{Rad-B (clear)})$   
*Rad-B (clear) - Rad-A (clear) - due to surface albedo changes instead of clouds*

Energy balance for ensemble members of PlioMIP I



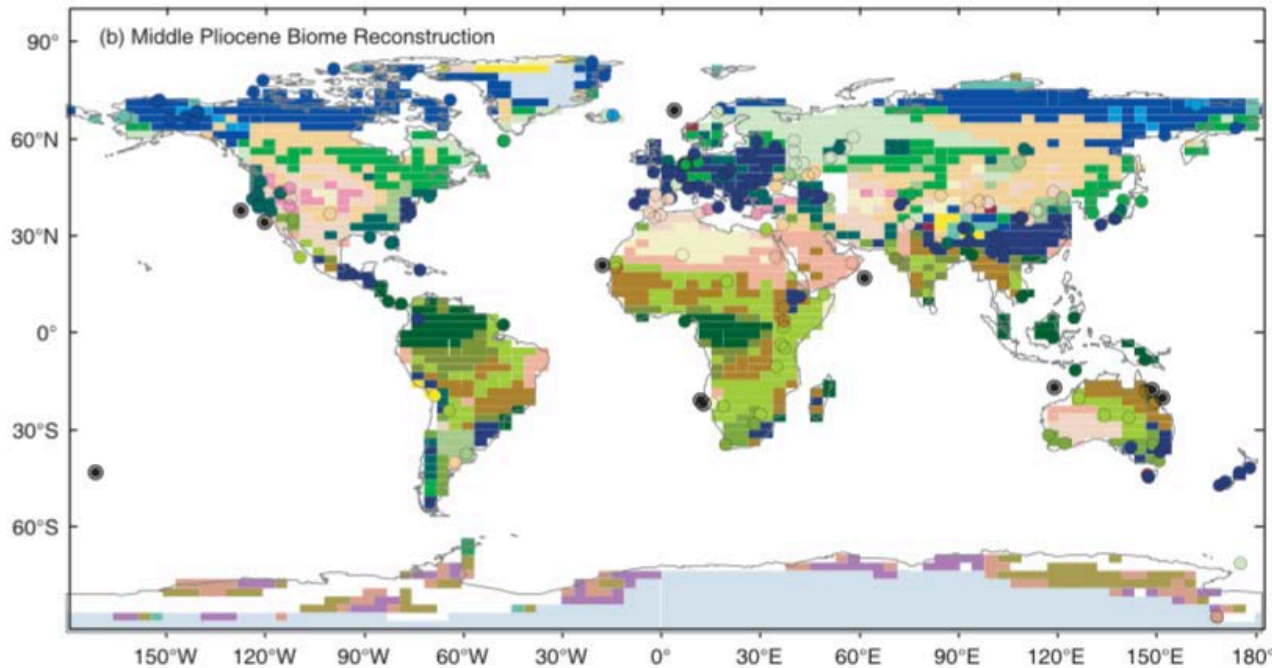
When surface albedo lowers,  $\text{Rad-B (clear)} - \text{Rad-A (clear)} < 0$ ,  $\Delta\text{Rad} < 0$ , which could mislead the previous analysis.

# Clean world in future



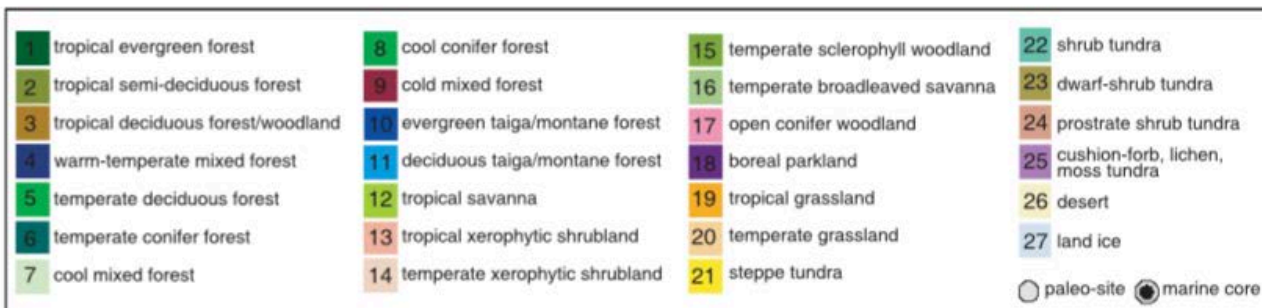
# Arctic climate during the Pliocene

Pliocene: 5.3 – 2.6 Ma

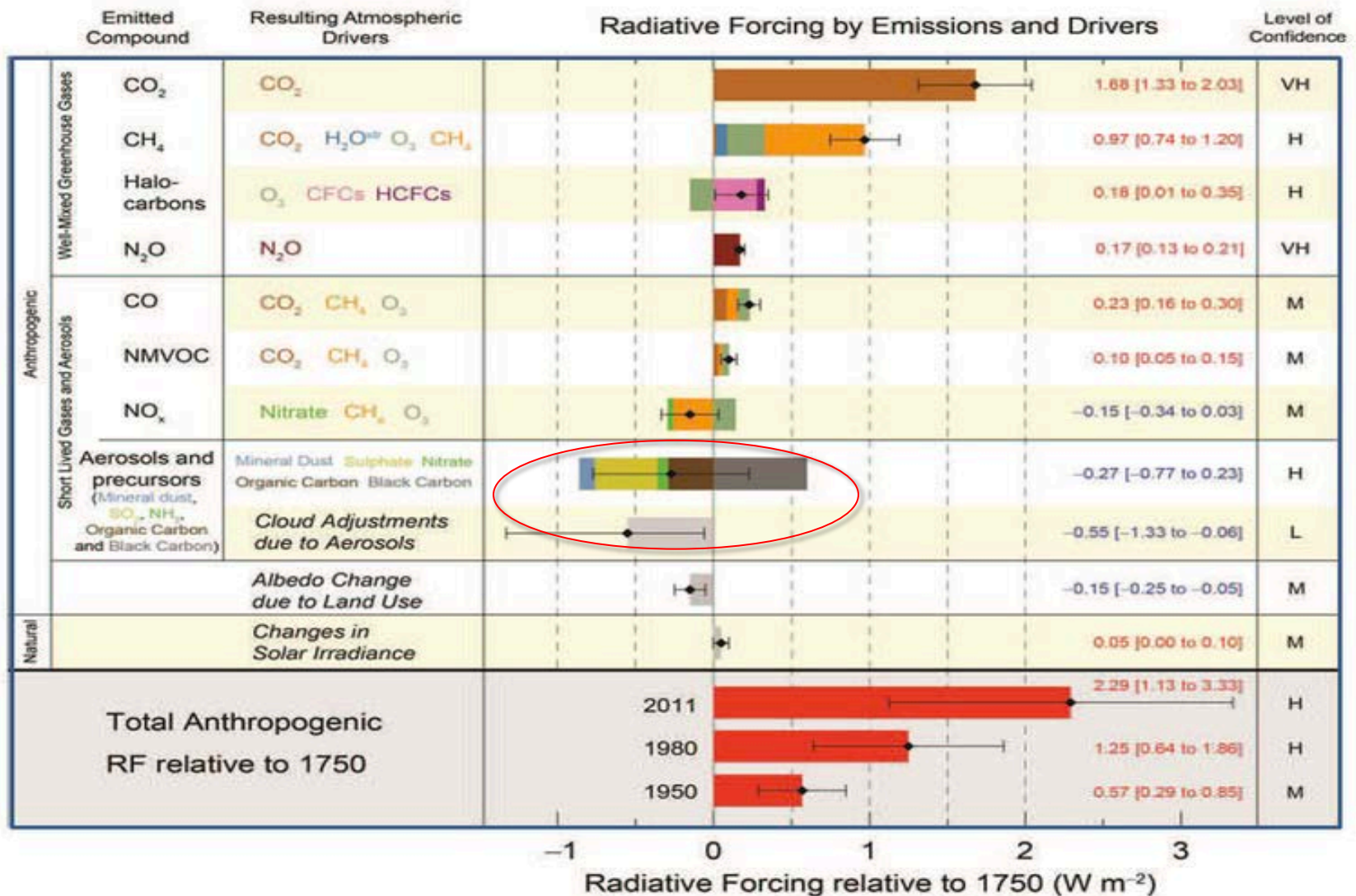


**Ts-Arctic: up to  
~18°C warmer  
than present-day  
at individual sites**

**CO<sub>2</sub>: close to  
present**

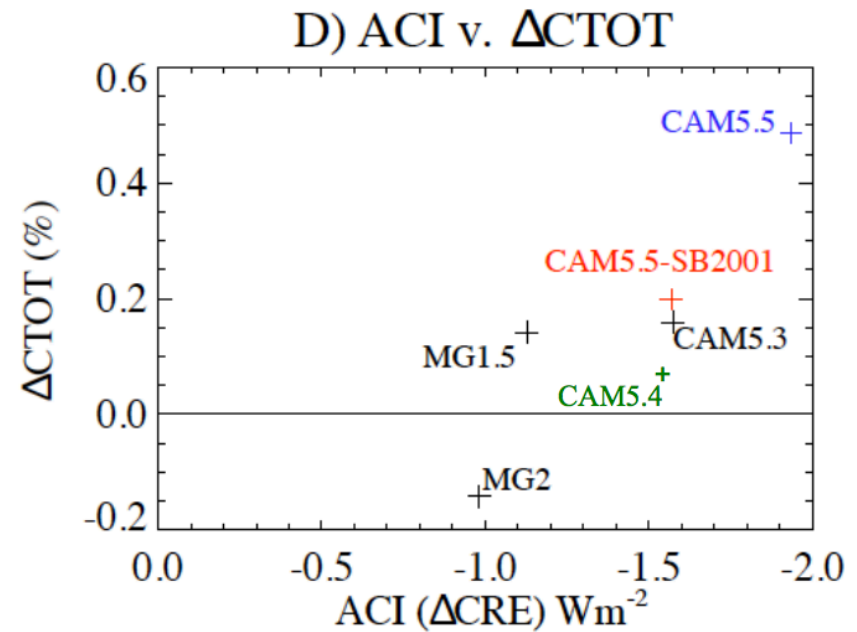
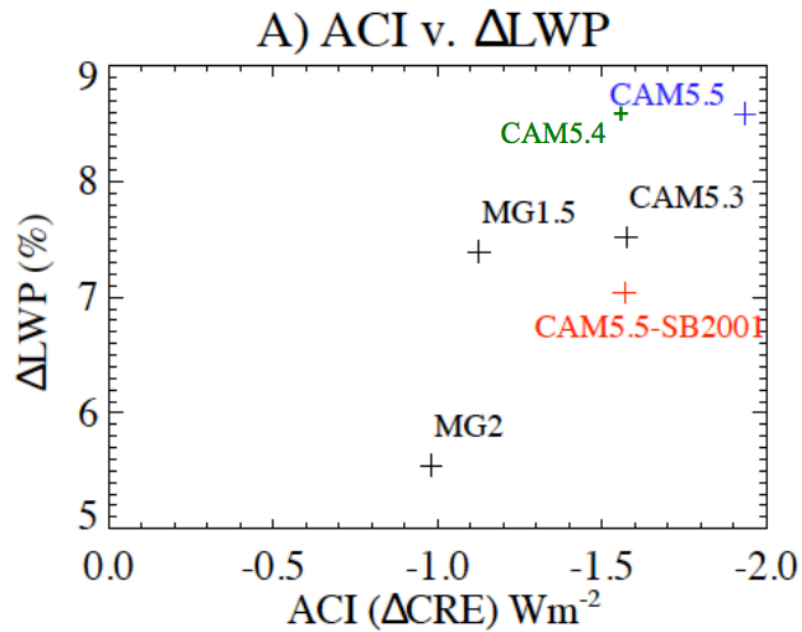


- IPCC AR5: total radiative forcing due to changes in aerosols is 0.7 W/m<sup>2</sup> (compare to 1.68 W/m<sup>2</sup> CO<sub>2</sub> forcing)



# ACI

ACI Definition following Ghan 2013



- Started (CAM5.3) with ACI about  $-1.5 Wm^{-2}$
- Decrease with MG1.5 and MG2
- Increase with CAM5.4 (mixed phase ice nucleation+ MAM4)
- Increase with CAM5.5 (shallow convective regime)
- Can Decrease with new Autoconversion (SB2001)

**Table 2.** Global (land/ocean) annual mean relationship slopes, computed as linear regression between the logarithm of cloud droplet number concentration ( $N_d$ ), liquid water path ( $L$ ), total cloud cover ( $f_{\text{cld}}$ ), cloud-top temperature ( $T_{\text{top}}$ ), planetary albedo ( $\alpha$ ), and outgoing long-wave radiation (OLR) with the logarithm of aerosol optical depth ( $\tau_a$ ). The land/ocean mean annual mean numbers are given as weighted mean of slopes for all seasons and land/ocean regions. Bold numbers show agreement with the Terra data to within  $\pm 25\%$ , gray, underestimation by up to a factor of two, blue, stronger underestimation, green, overestimation by up to a factor of two, red, stronger overestimation compared to the Terra data. The data are plotted in Fig. 2. Gaps indicate that a particular satellite or model did not report all quantities.

Relationship		Terra	Aqua	ORAC	CAM-NCAR	CAM-Oslo	CAM-PNNL	CAM-Umich	ECHAM5	GFDL	GISS	HADLEY	LMDZ-INCA	SPRINTARS
$N_d-\tau_a$	land	<b>0.083</b>	<b>0.078</b>		0.180	0.640	0.531	0.340	0.266	0.375	0.168	0.260	0.175	0.154
	ocean	<b>0.256</b>	<b>0.251</b>		0.408	0.787	0.471	0.348	0.111	0.155	0.162	0.483	0.198	0.213
$L-\tau_a$	land	<b>0.074</b>	0.100	0.148	3.064	0.389	0.218	0.313	0.363	1.557	0.192	0.333	0.896	0.690
	ocean	<b>0.134</b>	0.093	<b>0.136</b>	3.615	0.309	0.466	0.315	0.572	1.422	0.000	1.340	0.339	0.308
$f_{\text{cld}}-\tau_a$	land	<b>0.51</b>	<b>0.48</b>	0.27		0.34	-0.05	0.20	0.11	<b>0.52</b>	-0.04	0.11	0.09	0.13
	ocean	<b>0.31</b>	<b>0.29</b>	0.09		0.59	-0.00	<b>0.26</b>	0.00	1.09	0.15	<b>0.23</b>	0.14	0.21
$T_{\text{top}}-\tau_a$	land	-0.0064	-0.0083	-0.0064	-0.0013	-0.0154	0.0161	-0.0103	-0.0054	-0.0116	0.0083	0.0009	-0.0044	0.0003
	ocean	-0.0150	-0.0141	-0.0082	0.0046	0.0007	0.0195	0.0082	-0.0013	-0.0284	-0.0072	0.0097	-0.0049	0.0200
$\alpha-\tau_a$	land	<b>0.17</b>	<b>0.16</b>			<b>0.14</b>	-0.01	<b>0.13</b>	0.02			0.00	-0.04	0.02
	ocean	<b>0.26</b>	<b>0.25</b>			0.41	0.05	0.27	0.12			0.19	0.00	0.08
OLR $-\tau_a$	land	-0.028	-0.040		-0.070	-0.052	0.053	-0.034	-0.010	-0.060		0.014		0.006
	ocean	-0.050	-0.054		-0.109	-0.084	0.027	-0.042	0.025	-0.140		-0.017		0.034