

# Temporal Structure of ENSO in 20<sup>th</sup> Century Coupled Simulations

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Simona Masina

## Questions

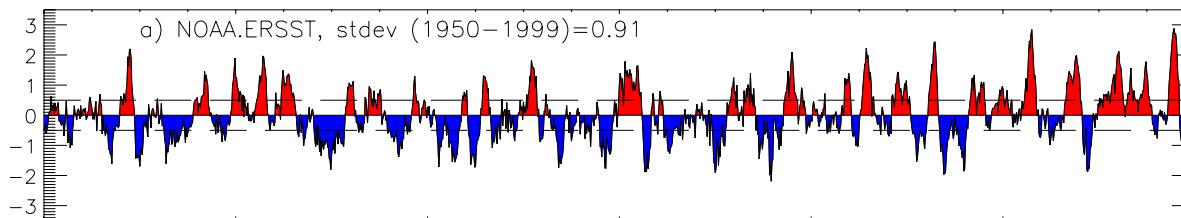
- Is ENSO dynamics consistent with the leading paradigms?
- Can we relate the dominant ENSO timescale to ocean dynamics and atmospheric forcing?
- Is that relationship consistent across different models?

## Models

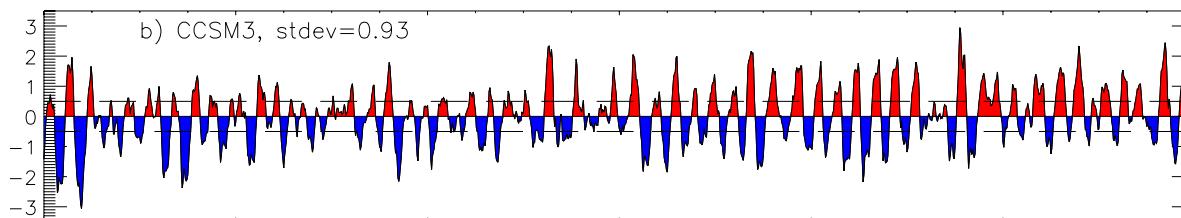
- CCSM3 (USA)
- GFDL-CM2.0 (USA)
- GISS-EH (USA)
- PCM (USA)
- IPSL-CM4 (France)
- CNRM-CM3 (France)
- UKMO-HadCM3 (UK)
- MRI-CGCM2.3.2 (Japan)
- CSIRO-CM3 (Australia)

# Niño3.4 index

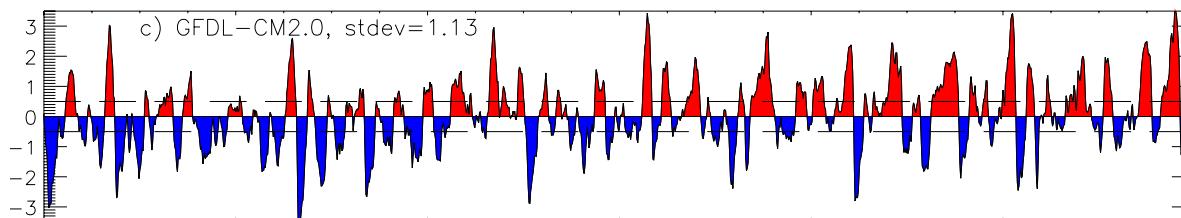
OBS  
NOAA.ERSST



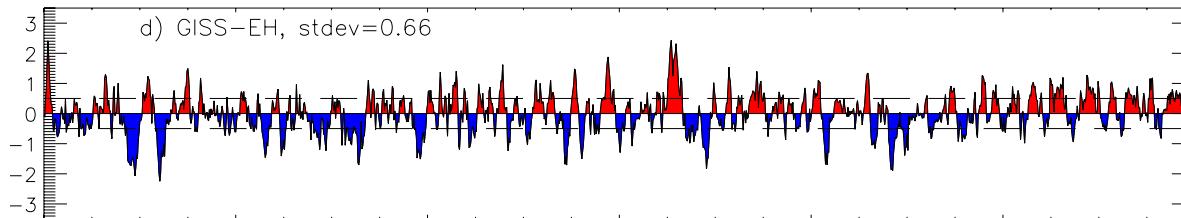
CCSM3



GFDL-CM2.0



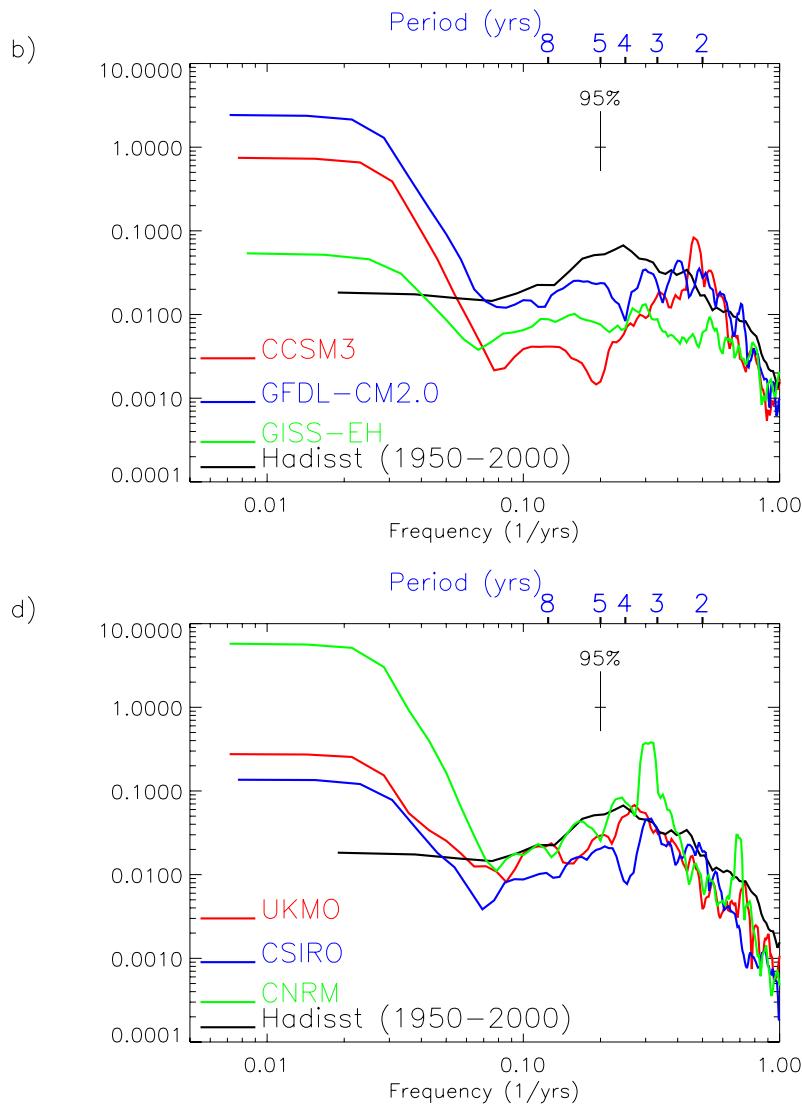
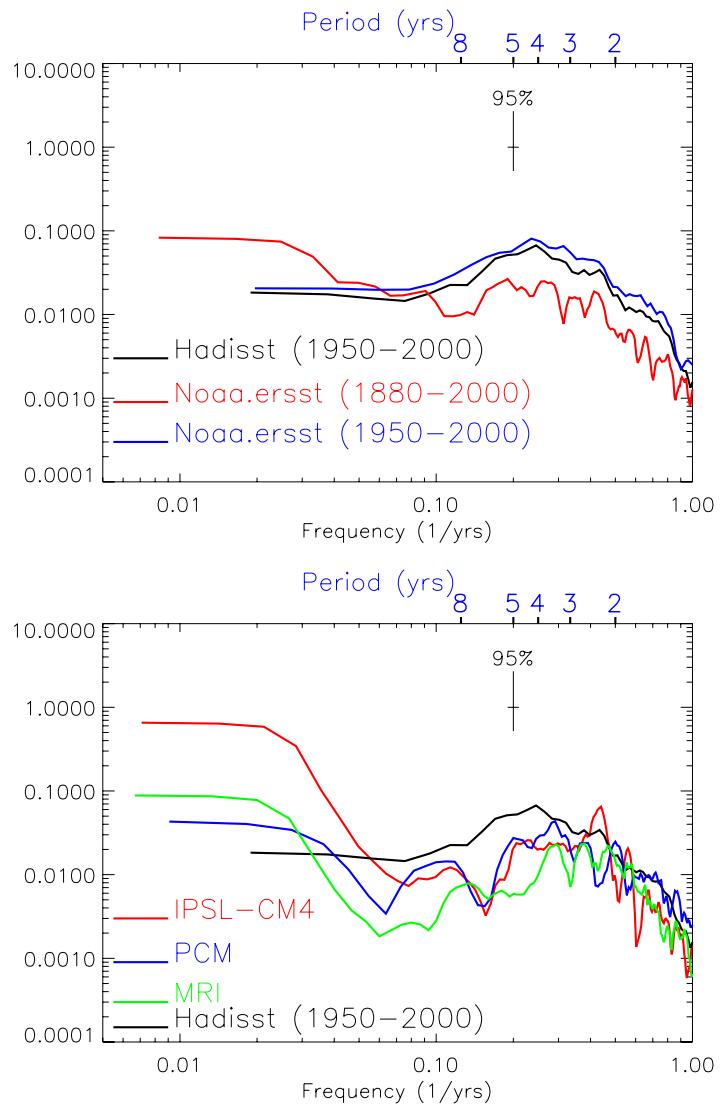
GISS-EH



1880

2000

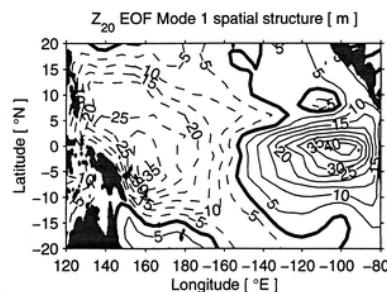
# Niño3.4 spectra



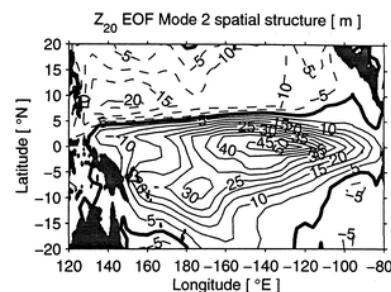
# Thermocline Variability

## Depth of 15°C isotherm (Z15)

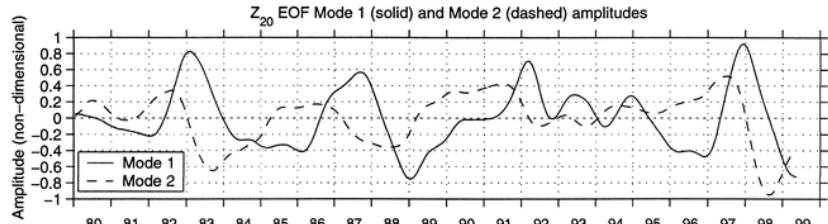
Meinen and McPhaden 2000



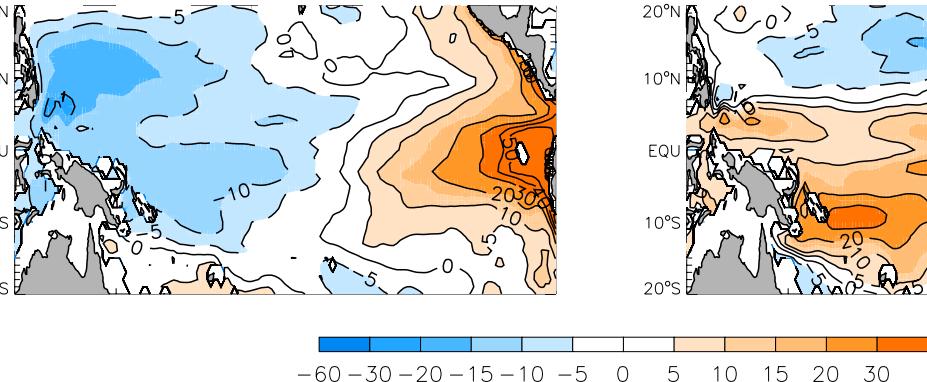
EOF1



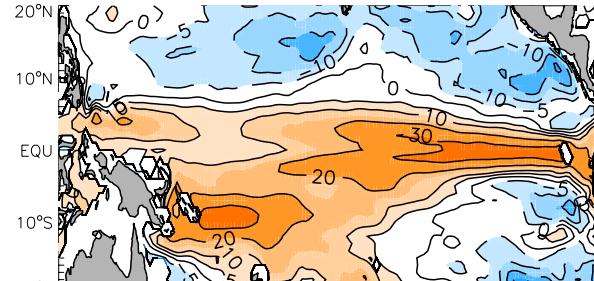
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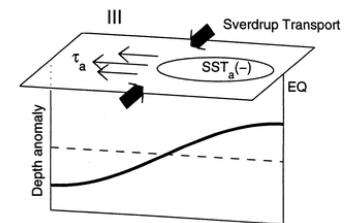
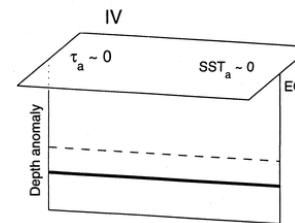
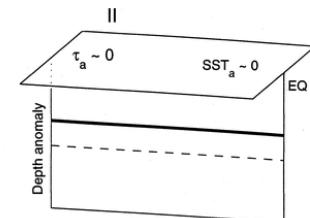
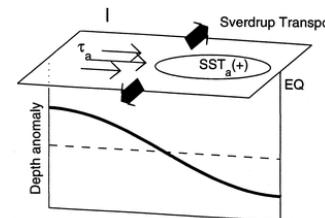
EOF1 45%



EOF2 16%



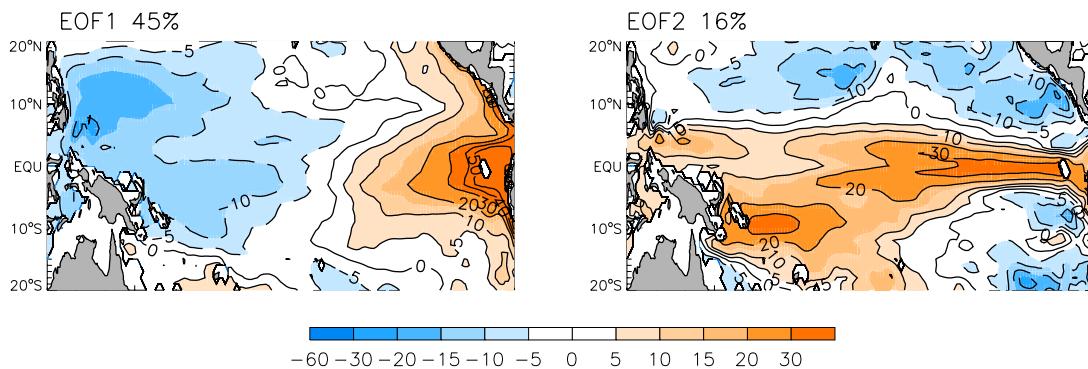
“Recharge oscillator”



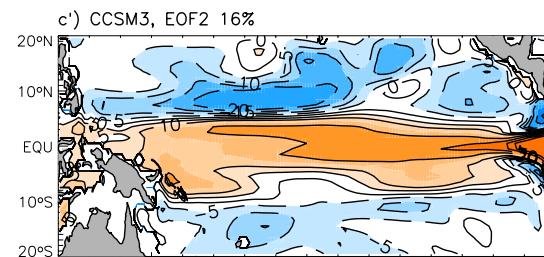
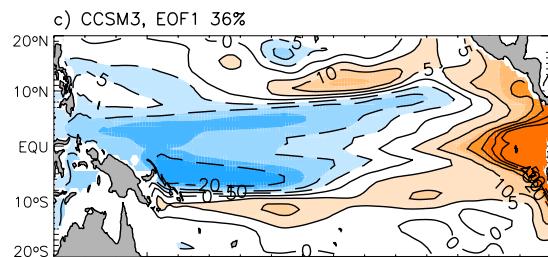
INGV ocean analysis

# EOFs of Z15

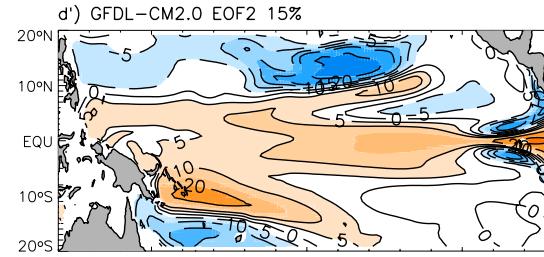
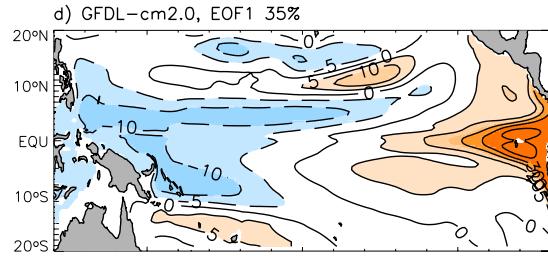
INGV



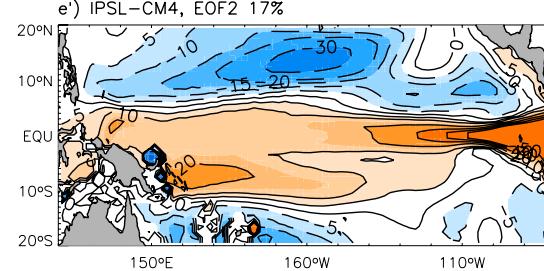
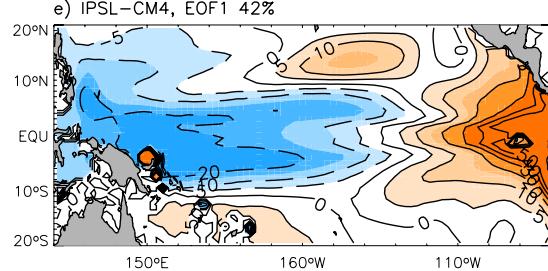
CCSM3



GFDL-CM2.0



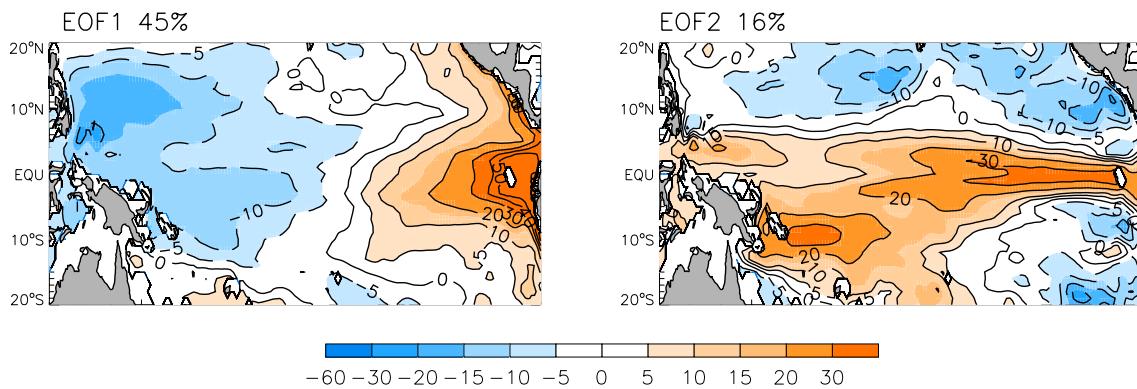
IPSL-CM4



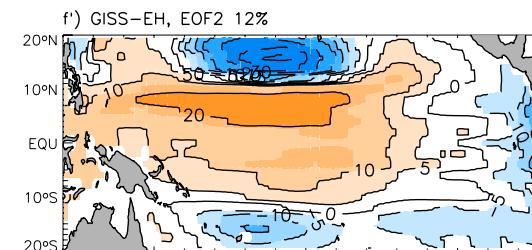
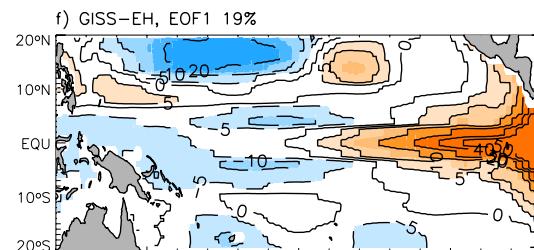
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# EOFs of Z15

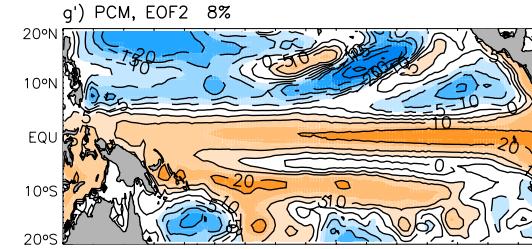
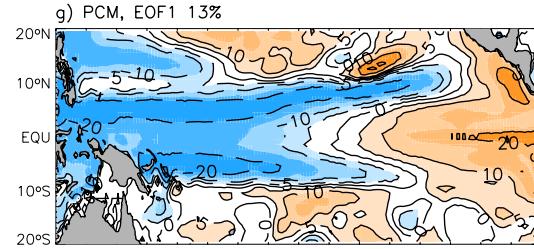
INGV



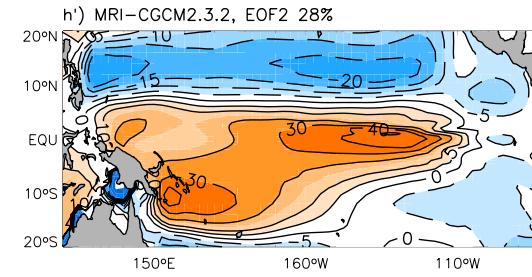
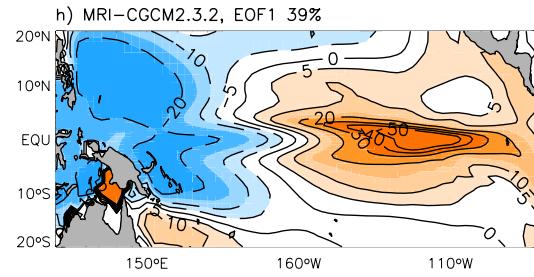
GISS-EH



PCM

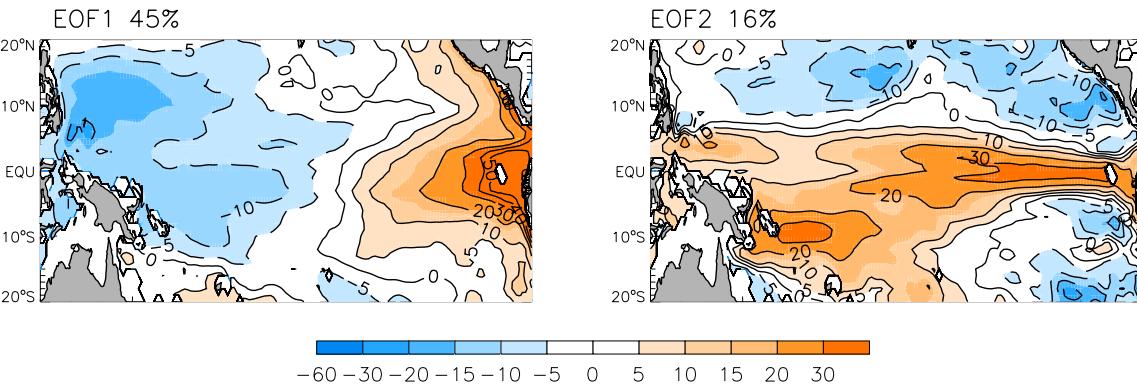


MRI



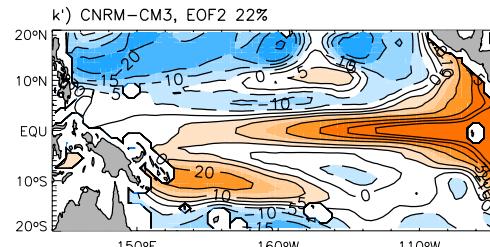
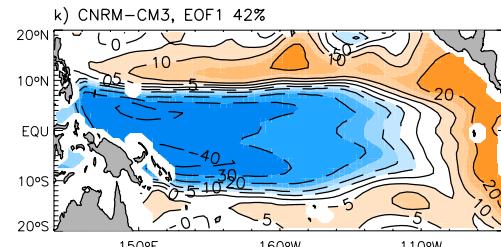
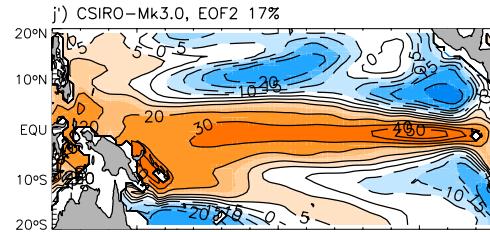
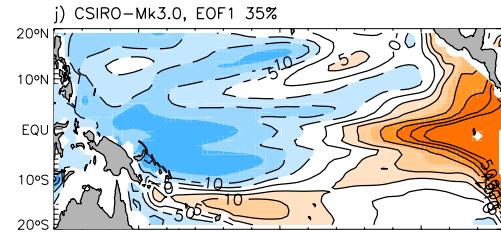
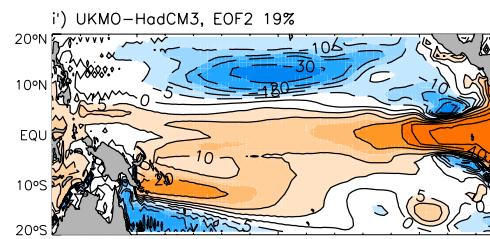
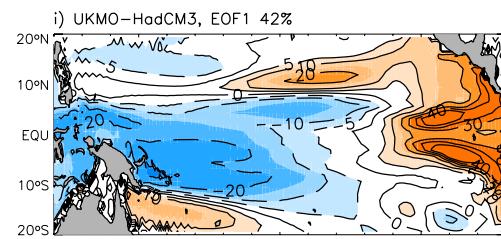
-60 -30 -20 -15 -10 -5 0 5 10 15 20 30

# EOFs of Z15



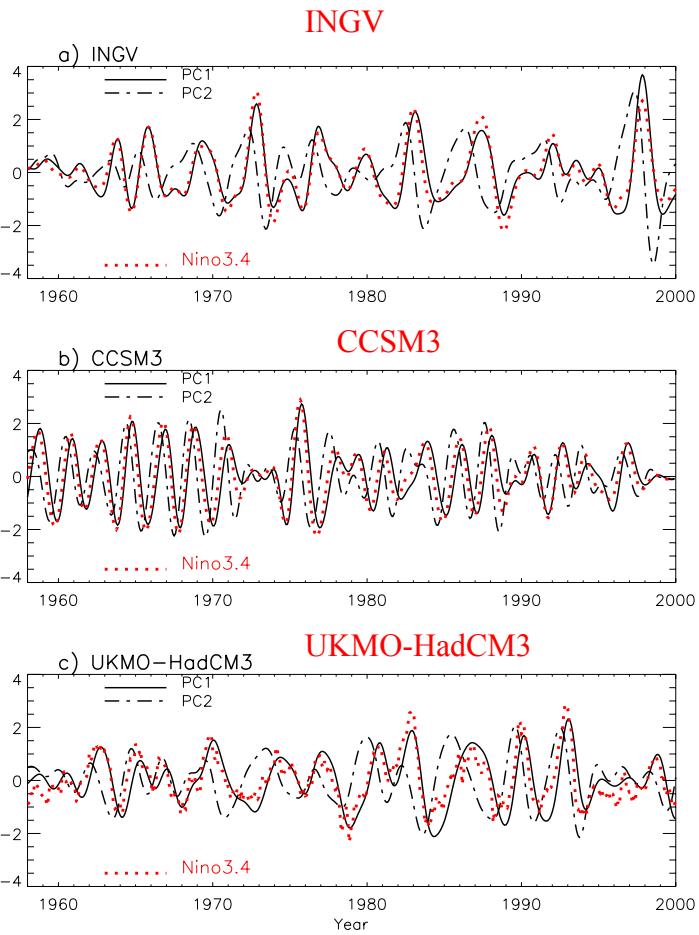
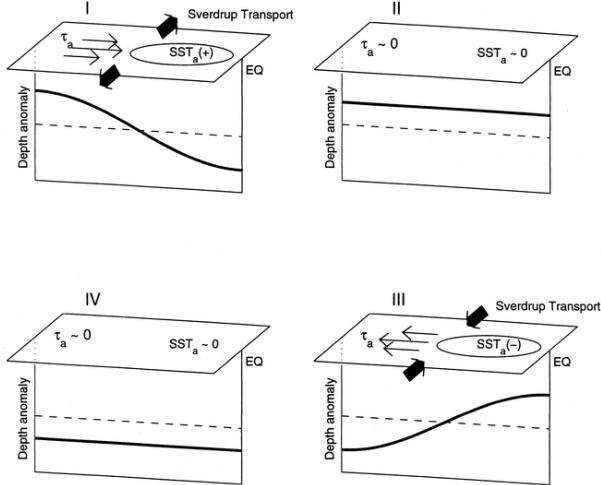
INGV

UKMO-HadCM3

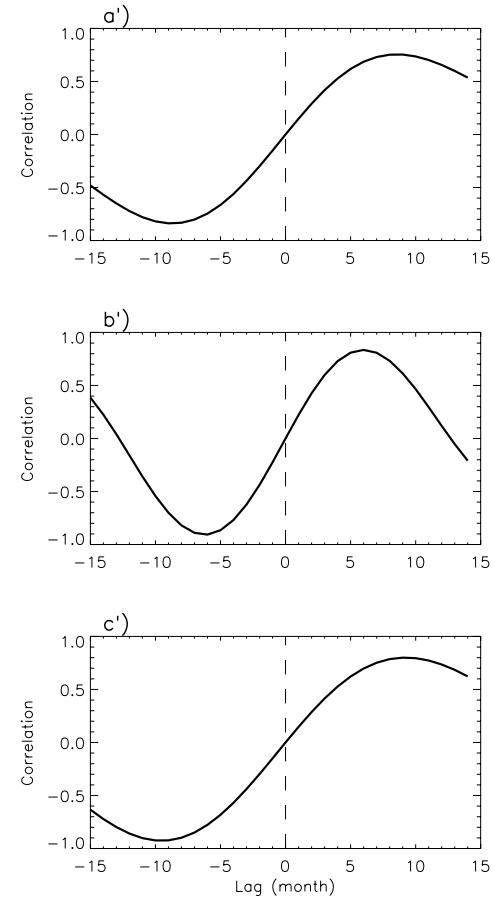


CNRM

# Phase relationship between the Z15 modes



PC2-PC1 lag-correlation

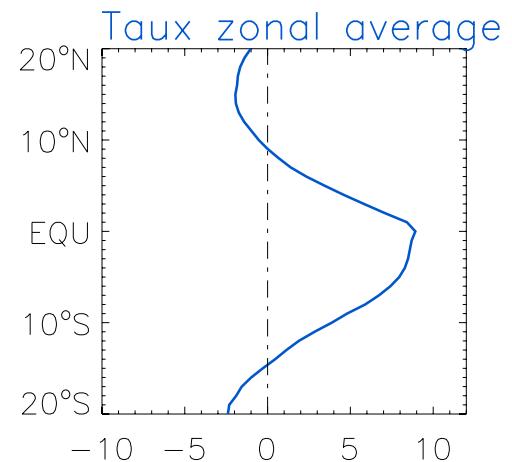
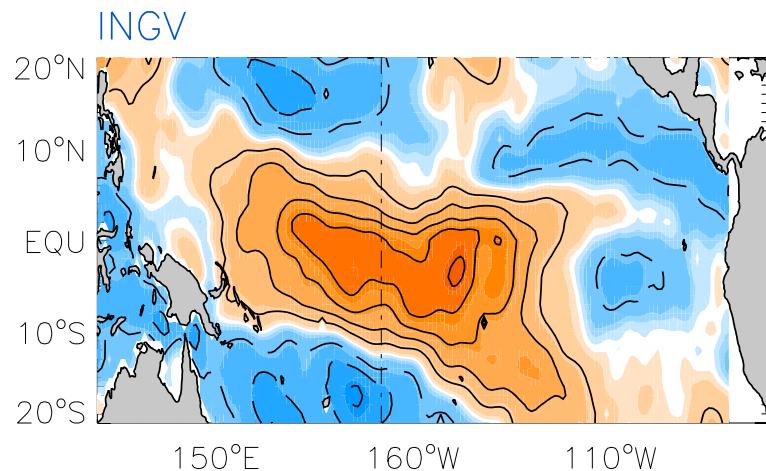


## Wind forcing

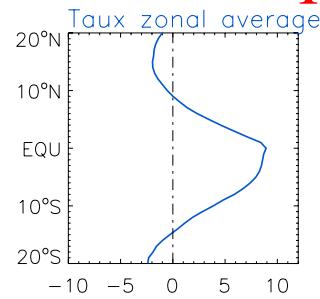
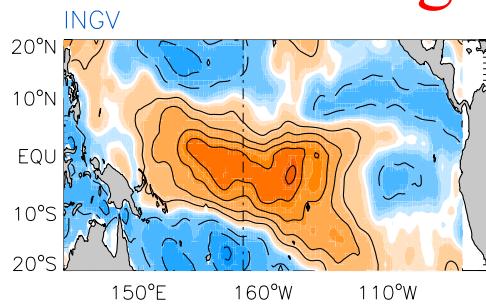
Studies performed with intermediate coupled models have shown that the spatial structure of the wind stress anomalies can influence the ENSO timescale.

- Meridional width of the wind stress anomalies: if the wind stress extends over a broader range of latitudes the adjustment timescale of the equatorial ocean increases.
- Longitudinal position of the wind stress anomalies: it influences the zonal advective feedback. Eastward displacement of the wind stress anomalies favors ENSO growth and longer duration.

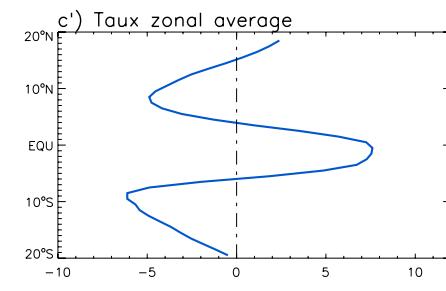
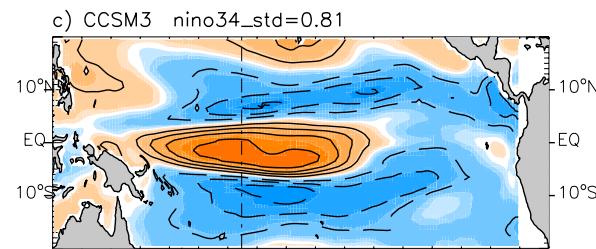
# Regression of $\tau^x$ upon the Niño3.4 index



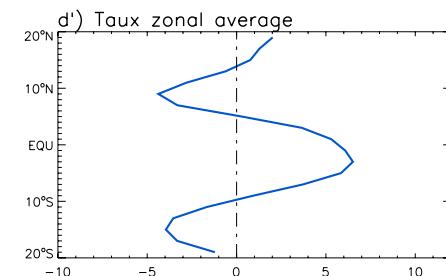
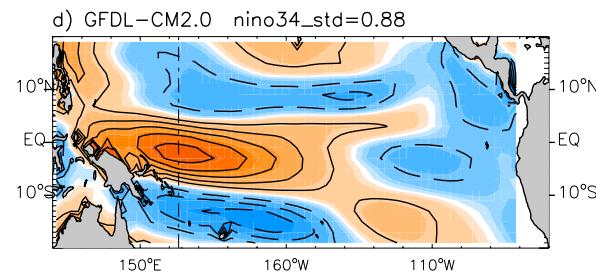
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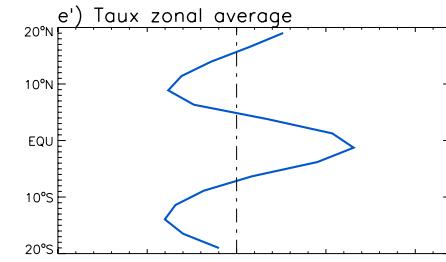
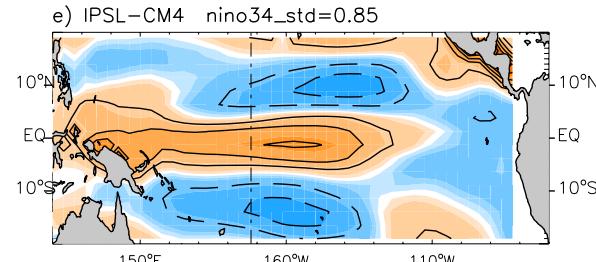
CCSM3



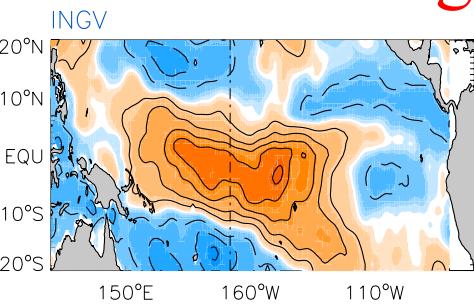
GFDL-CM2.0



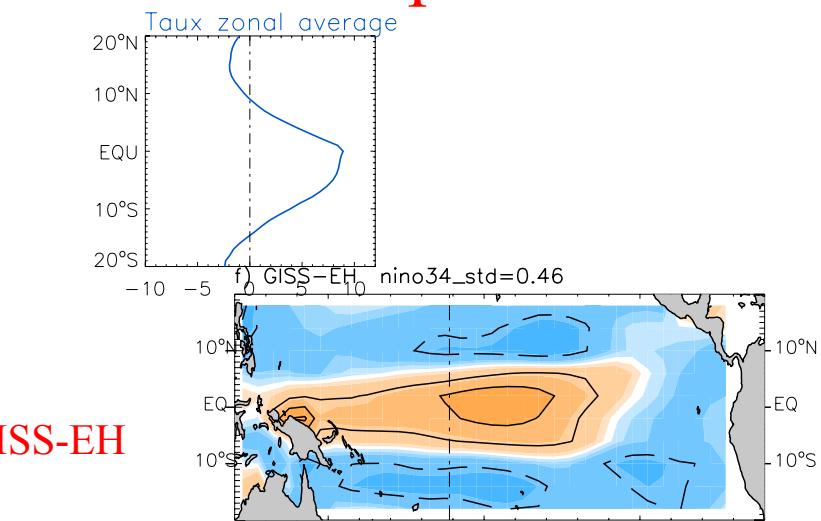
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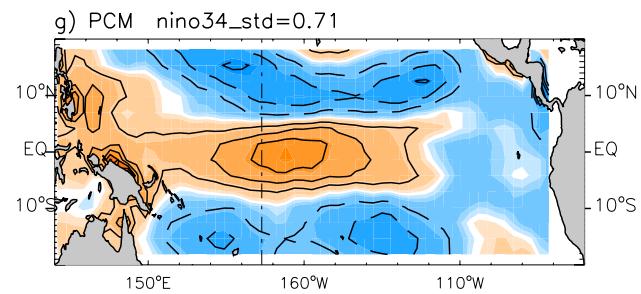
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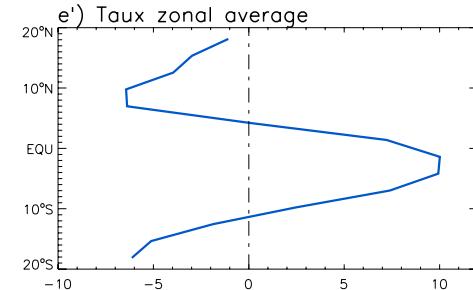
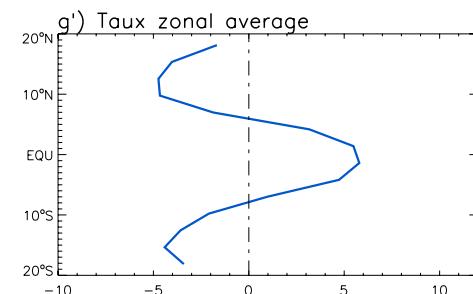
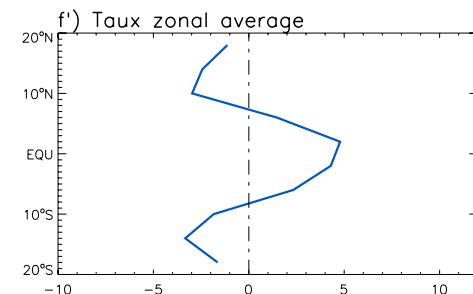
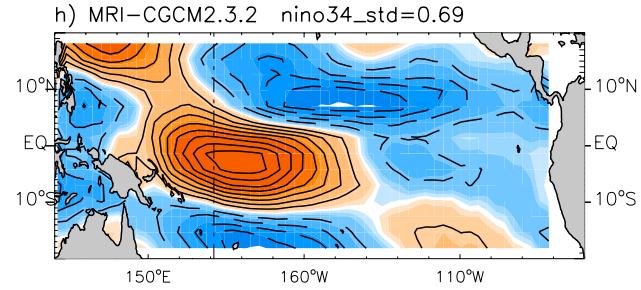
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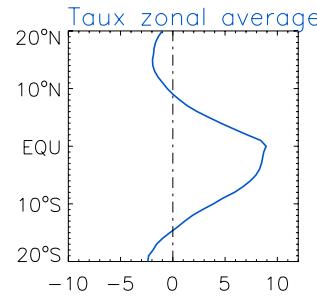
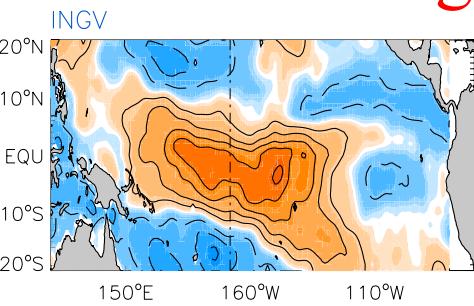
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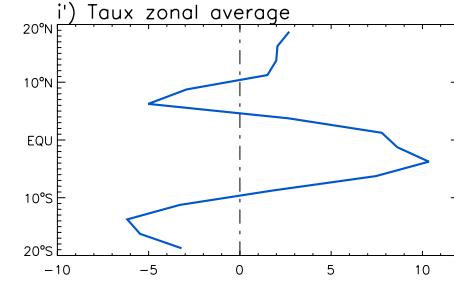
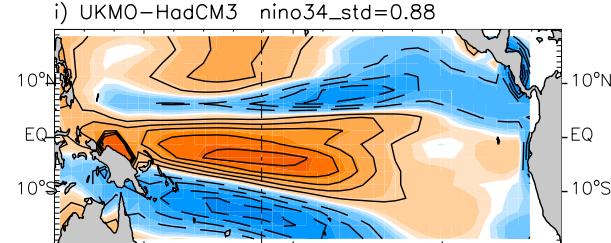
MRI



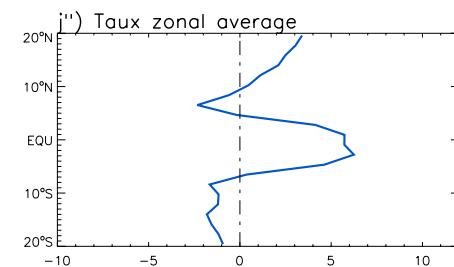
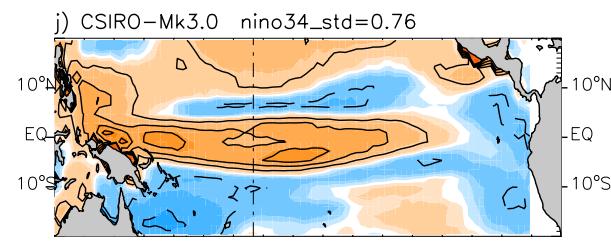
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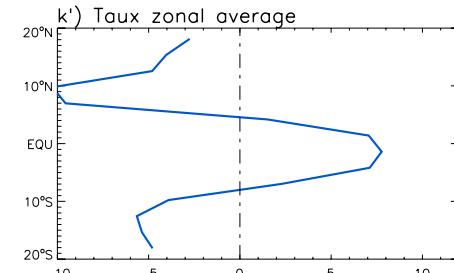
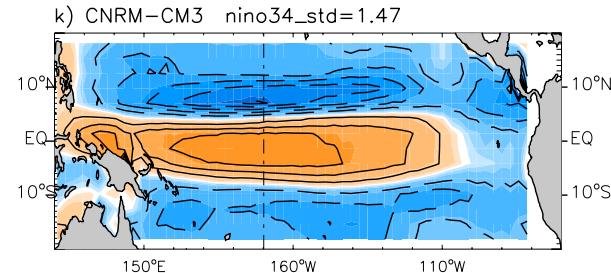
UKMO-HadCM3



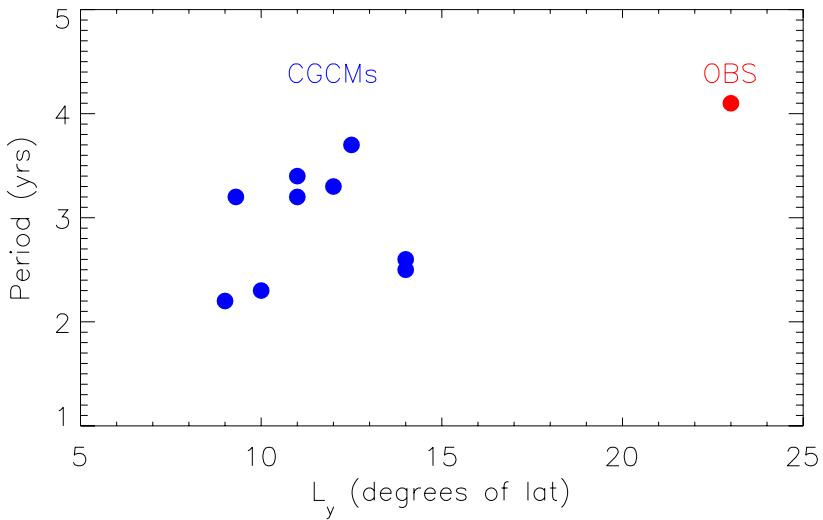
CSIRO



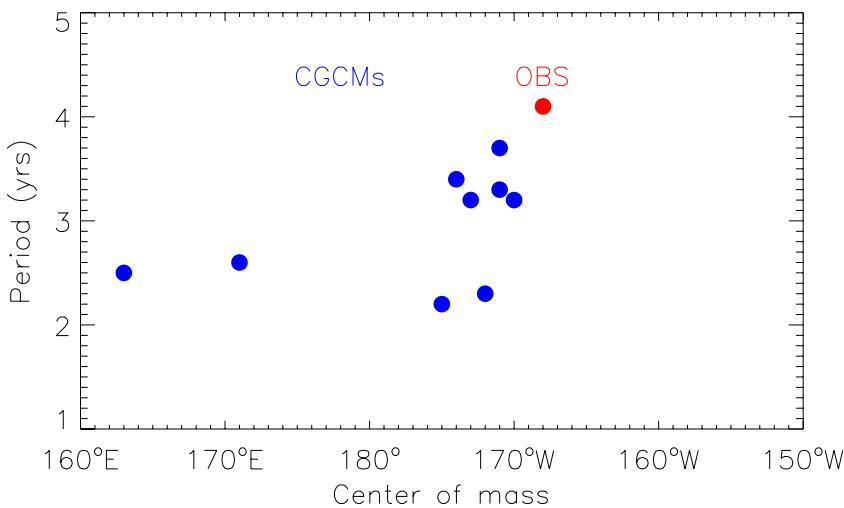
CNRM



# CGCMs vs. OBS.



All the CGCMs have wind stress anomalies narrower than obs.



All the CGCMs have wind stress anomalies displaced westward with respect to the obs.

## Conclusions

- Thermocline variability in the CGCMs is consistent with the “recharge oscillator” paradigm for ENSO.
- Wind stress forcing in the CGCMs has a narrower meridional scale and is displaced westward compared to observations. Both factors can lead to a shorter timescale for ENSO.