

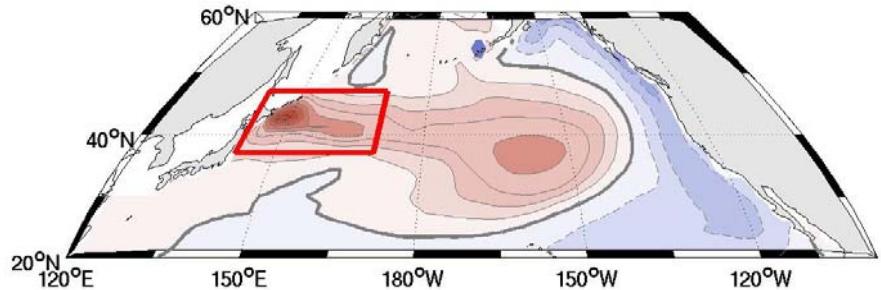
North Pacific Decadal Climate Variability in CCSM3

**Young-Oh Kwon and Clara Deser
NCAR**

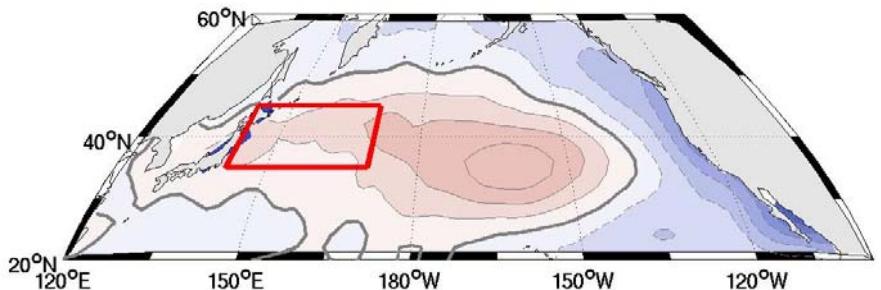
10th Annual CCSM Workshop (June 21-23, 2005 / Breckenridge, CO)

Winter SST Anomalies EOF1

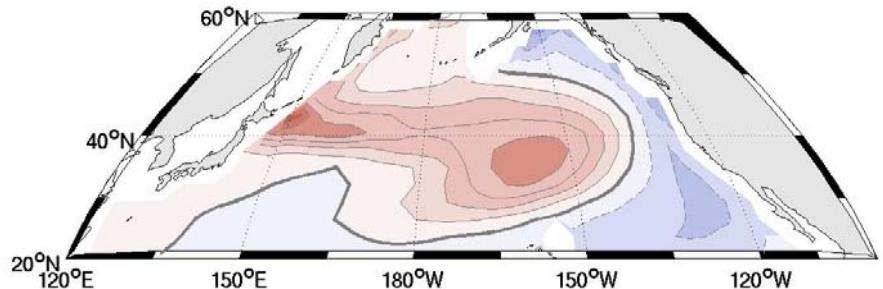
CCSM3 T85 (Yr : 100 – 599; 28 %)



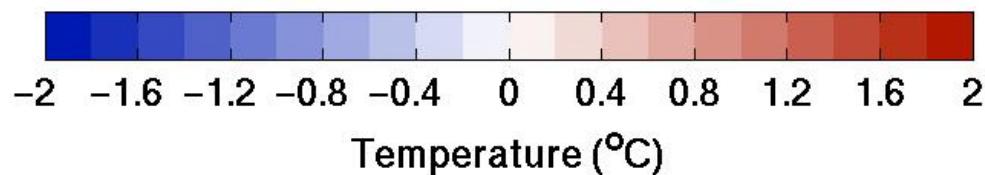
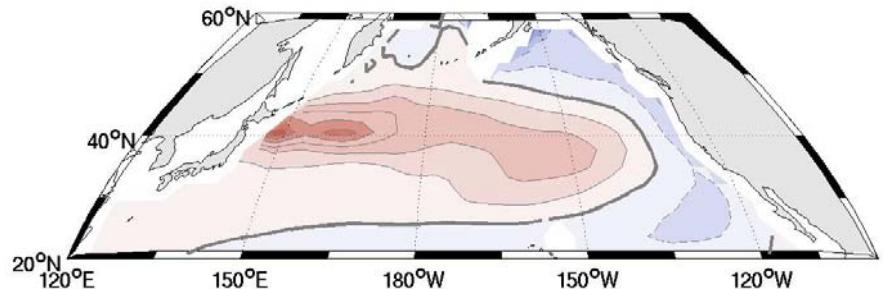
Observation (ERSST : 1901 – 2000; 28 %)



CCSM3 T42 (Yr : 100 – 999; 26 %)



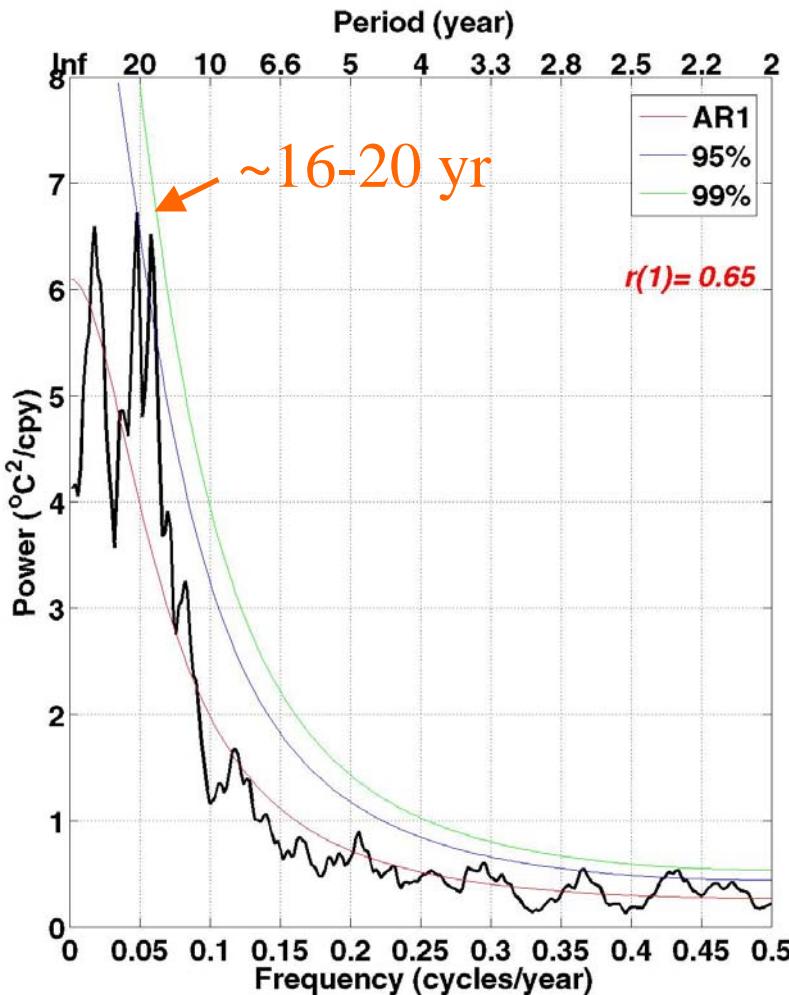
CCSM2 (T42) (Yr : 350 – 999; 20 %)



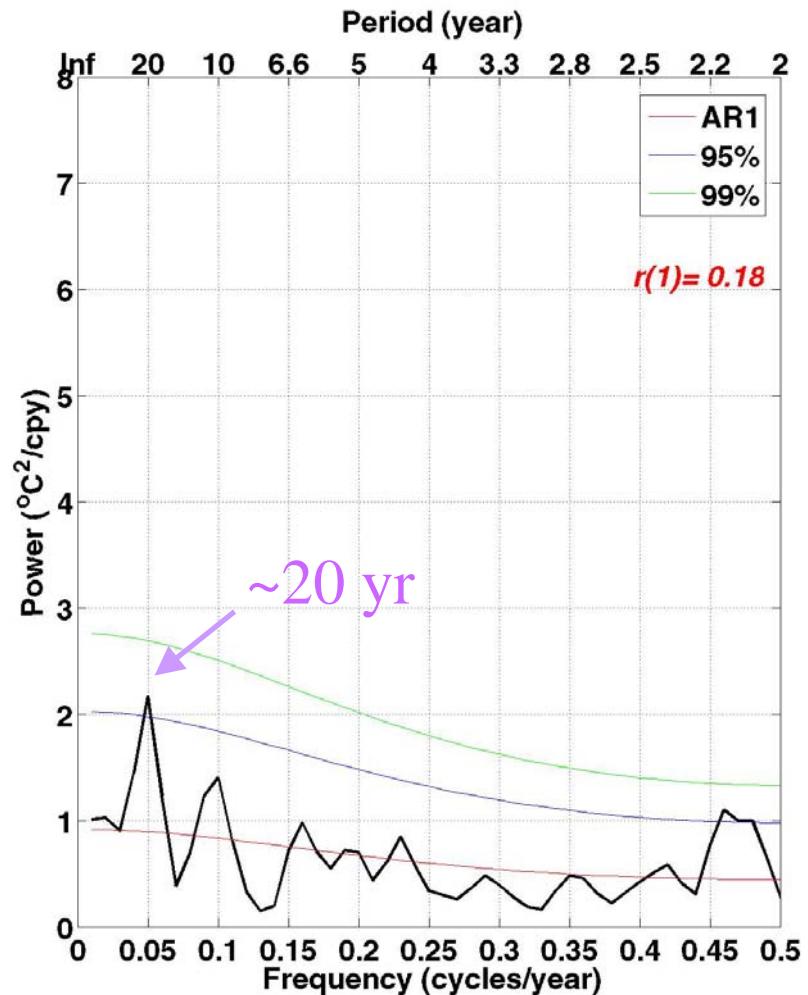
Power Spectrum

Winter SST Kuroshio Extension Index

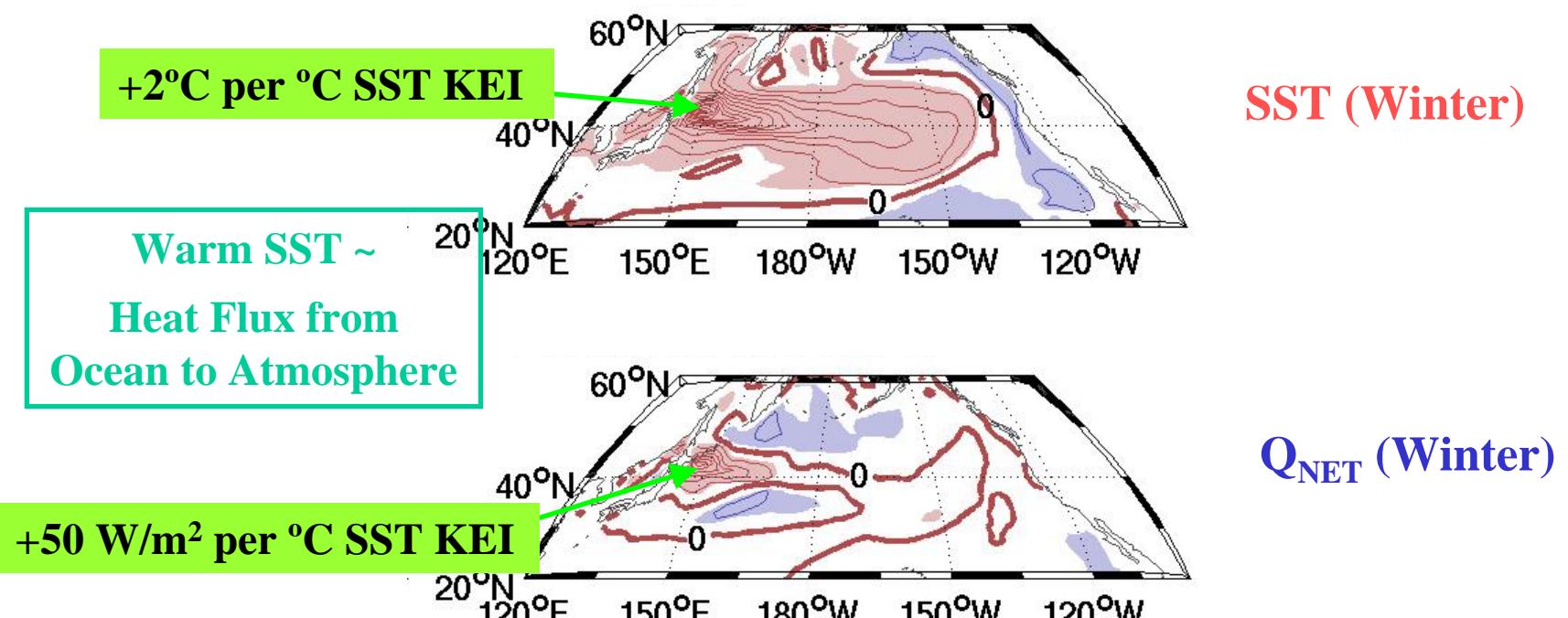
CCSM3 T85 (100-599)



Observation (ERSST: 1901-2000)



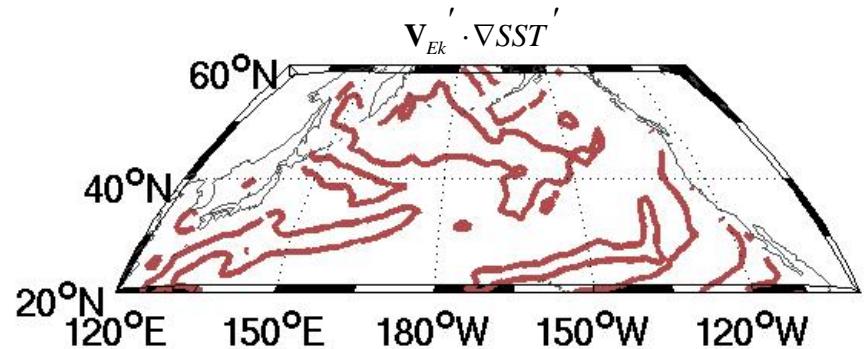
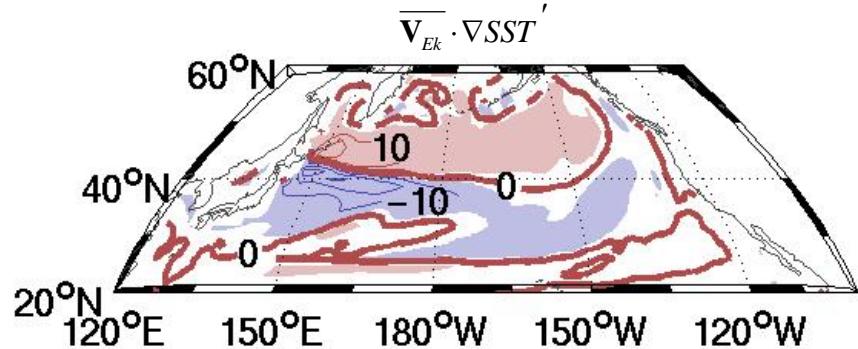
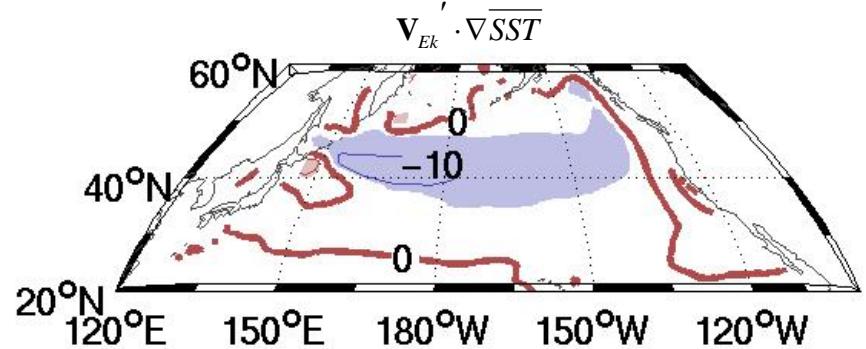
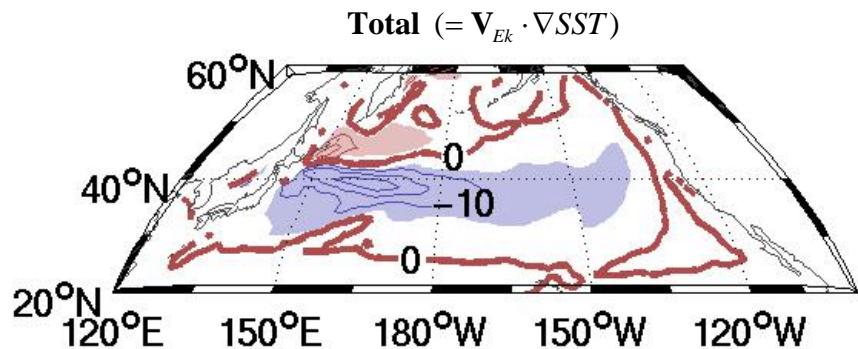
Regression on SST Kuroshio Extension Index



(Contour Interval: 0.2°C/°C, 10 W/m²/°C ; Shading: significant at 99 %)

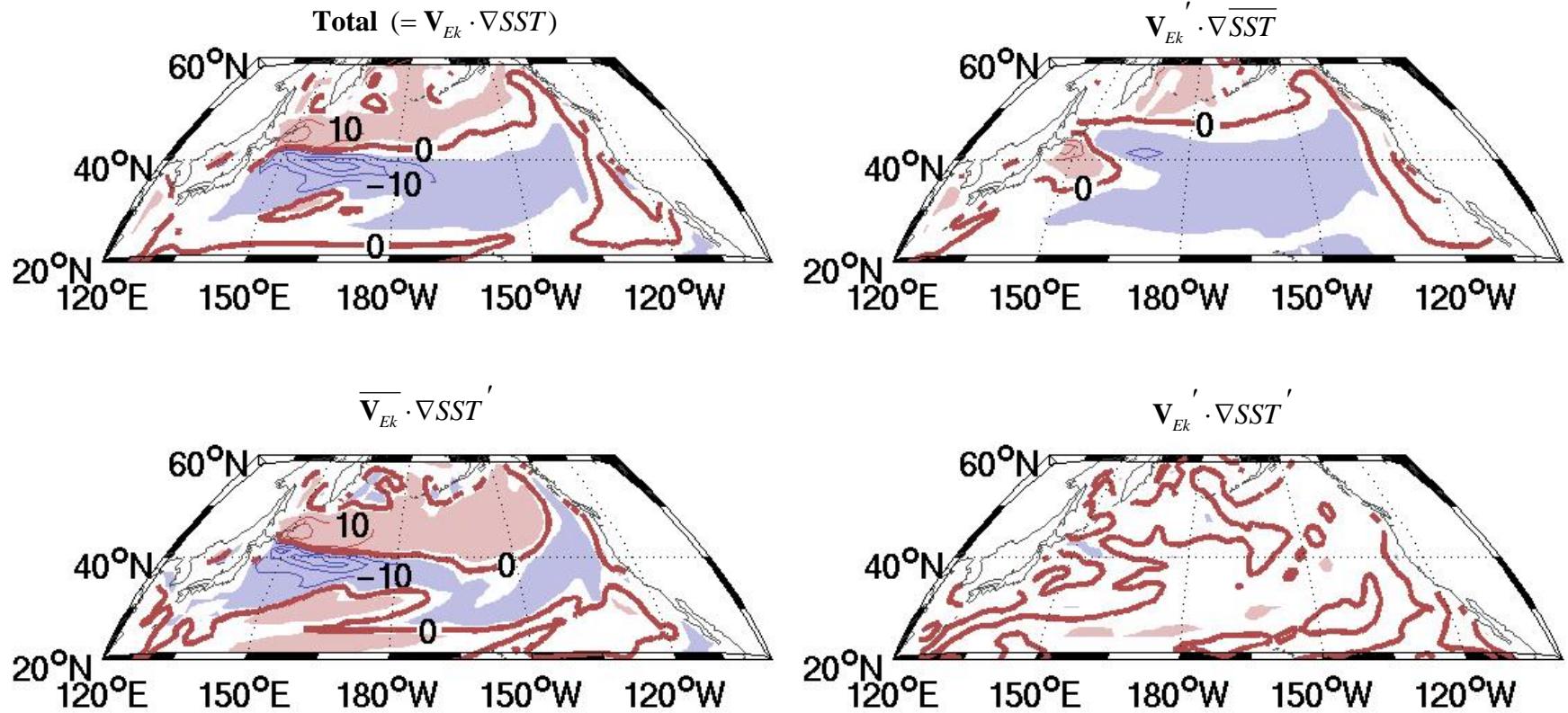
⇒ Local Q_{NET} responds to the SST anomaly, rather than forcing it.

Winter Q_{Ekman} regressed on SST Kuroshio Extension Index



(Contour Interval: 10 W/m²/°C ; Shading: significant at 99 %)

Winter Q_{Ekman} regressed on SST Kuroshio Extension Index (Low-pass filter > 10 yr)



(Contour Interval: 10 $W/m^2/\text{°C}$; Shading: significant at 99 %)

⇒ Anomalous Q_{Ekman} acts as a positive feedback to the SST anomalies.

Upper 200 m Heat Budget in the Kuroshio Extension

$$\frac{\partial \text{HC}}{\partial t} = -Q_{\text{net}} + u \frac{\partial \text{HC}}{\partial x} - v \frac{\partial \text{HC}}{\partial y} - w \frac{\partial \text{HC}}{\partial z} + \text{Residual} \quad [\text{W/m}^2]$$

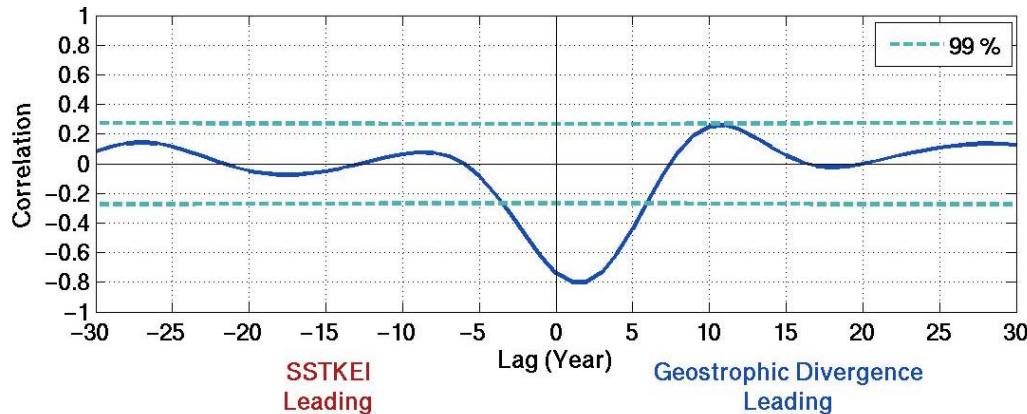
where $\text{HC} = \rho_0 C_p \int_0^z T(z) dz$

Surface Heat Flux

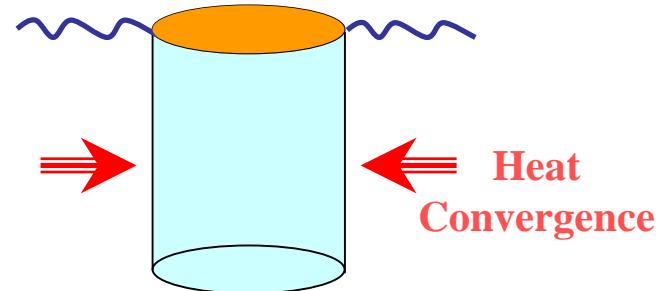
Lateral Ocean Heat Divergence

Correlation in the Kuroshio Extension

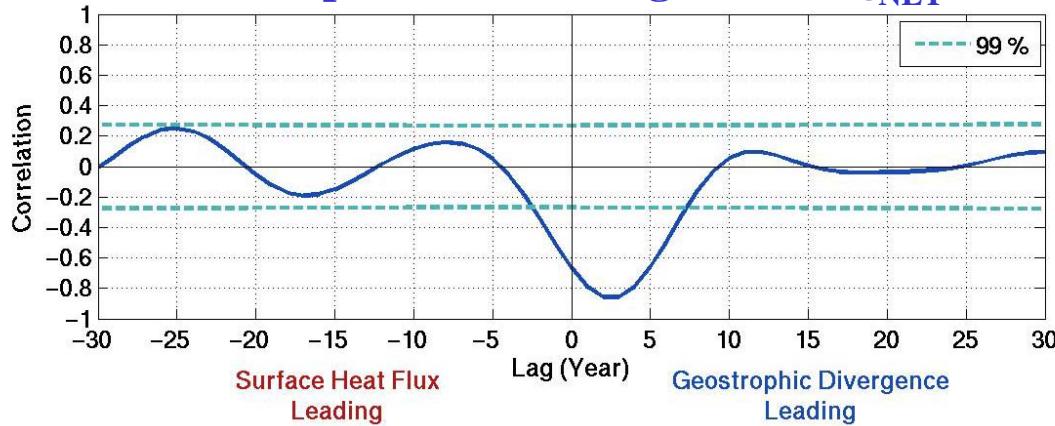
Geostrophic Heat Divergence vs. SST



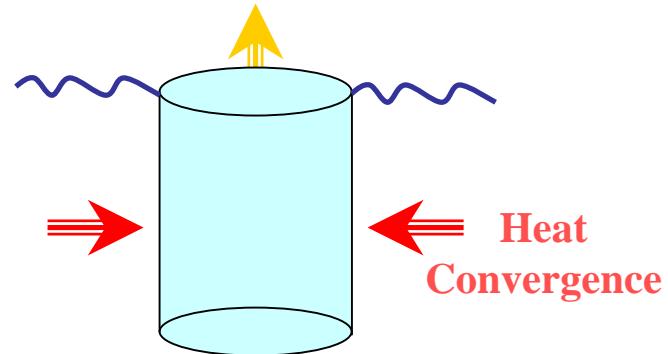
Warmer SST



Geostrophic Heat Divergence vs. Q_{NET}



Heat Flux out of Ocean



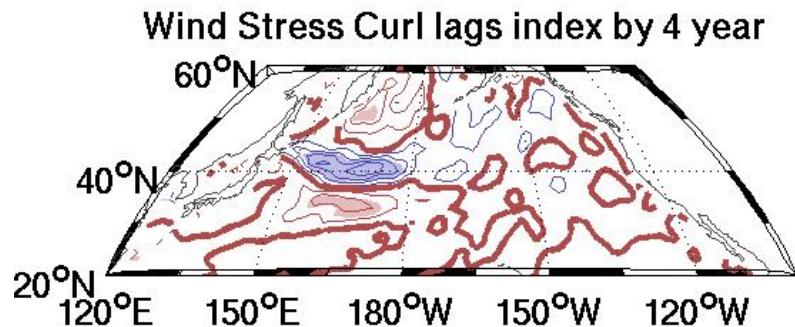
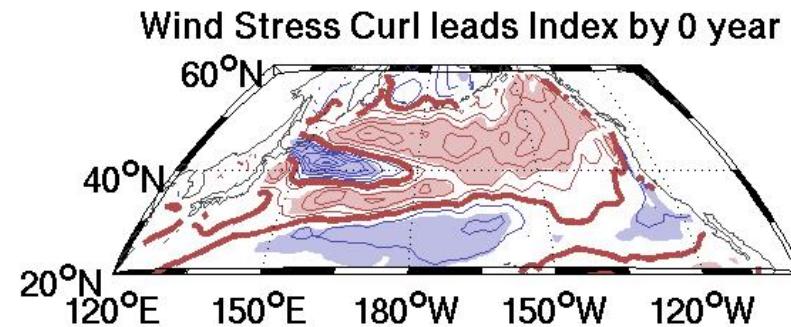
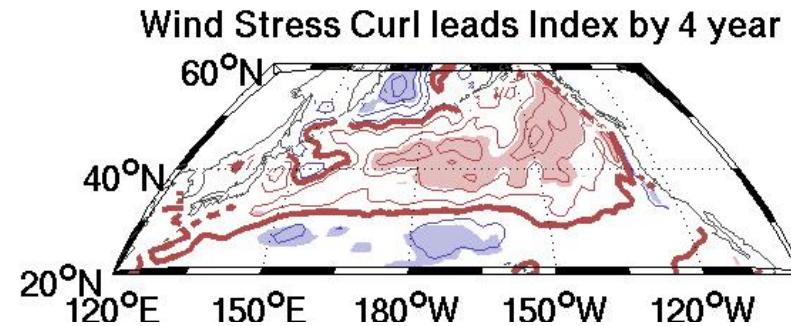
(Low-pass Filter > 10 yr)

Ocean heat convergence $\Rightarrow +1\text{-}2$ yr \Rightarrow warmer SST $\Rightarrow +1\text{--}2$ yr \Rightarrow ocean-to-atmosphere Q_{NET}

Summary

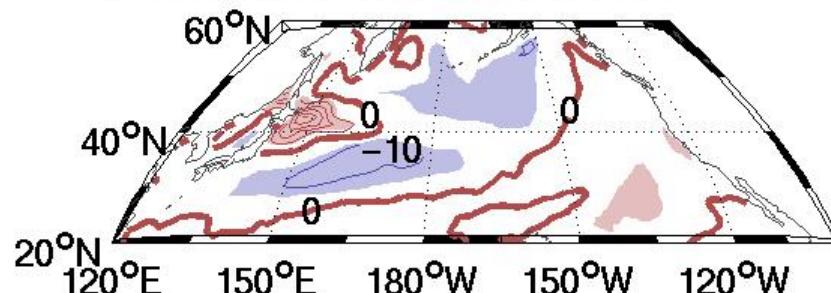
- Leading EOFs of CCSM3 & CCSM2 winter SST have PDO-like horse-shoe pattern.
- Robust decadal (16-20 yr) variability of SST along the Kuroshio Extension.
- Ocean heat divergence driven by stochastic wind stress curl forcing is likely causing the decadal SST anomaly along the Kuroshio Extension.
- Local Q_{net} and Q_{Ekman} are response to the SST anomaly.

Wind Stress Curl regressed on Heat Divergence in the Kuroshio Extension

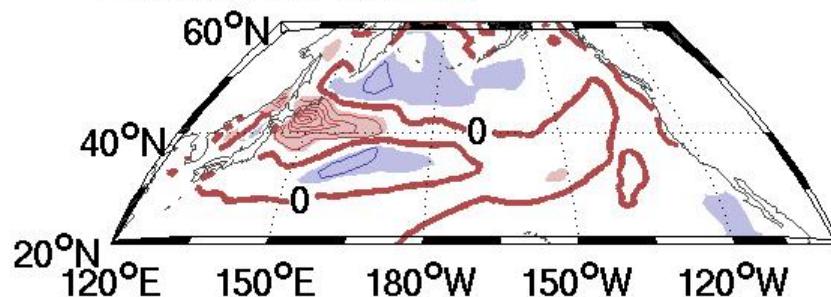


(Contour Interval : $5 \times 10^{-9} \text{ N/m}^3$; Shading : significant at 99 % ; Low-pass filter > 10 yr)

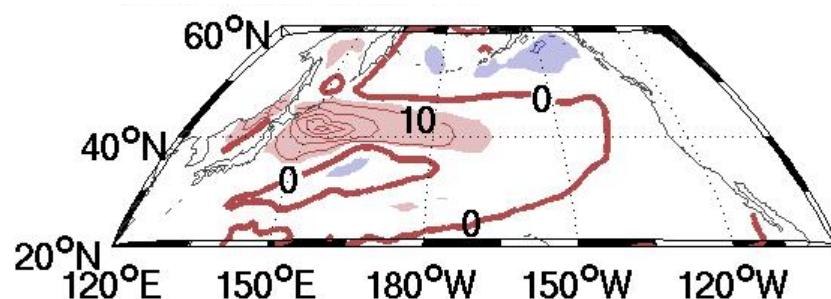
Winter Q_{NET} regression on SST Kuroshio Extension Index



Q_{NET} Leads (+1 Yr)



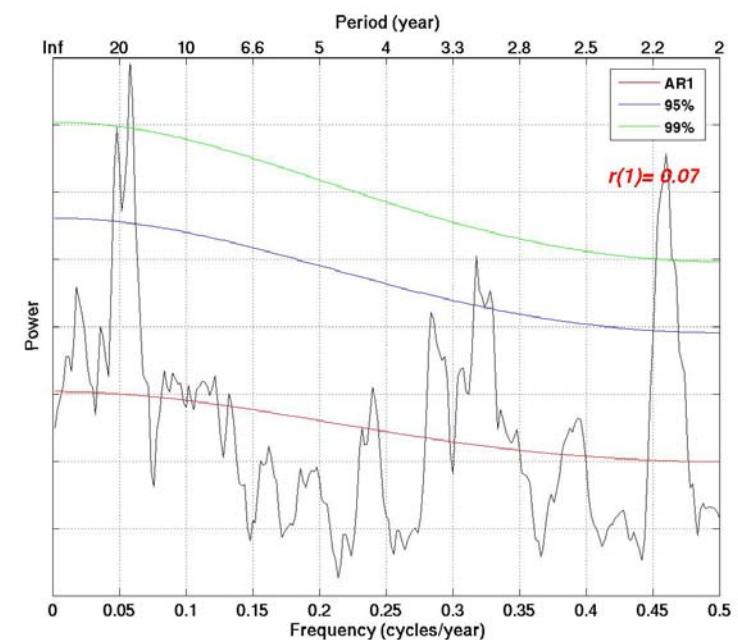
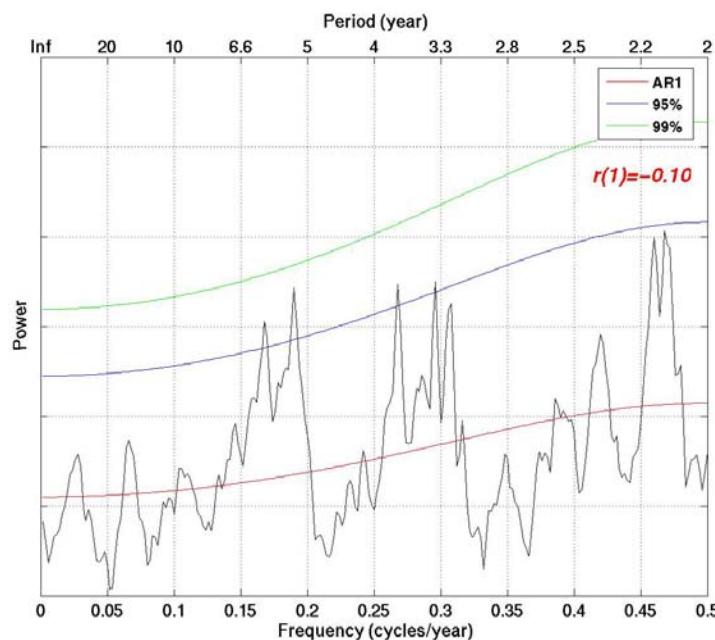
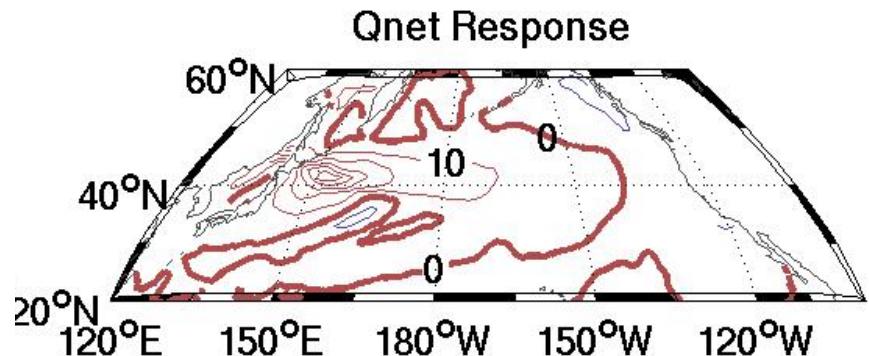
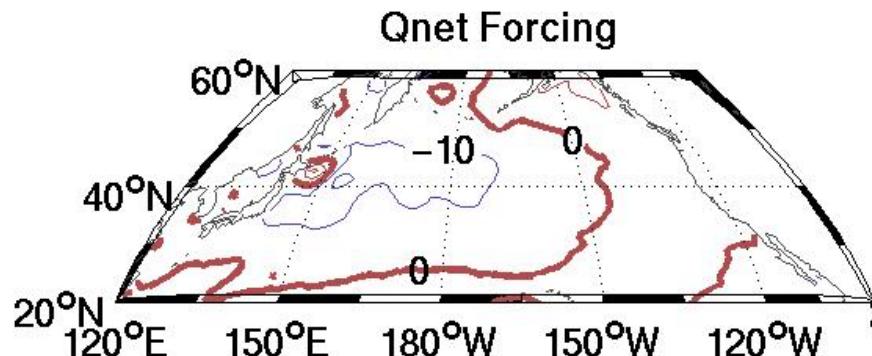
Simultaneous



Q_{NET} Lags (-1 Yr)

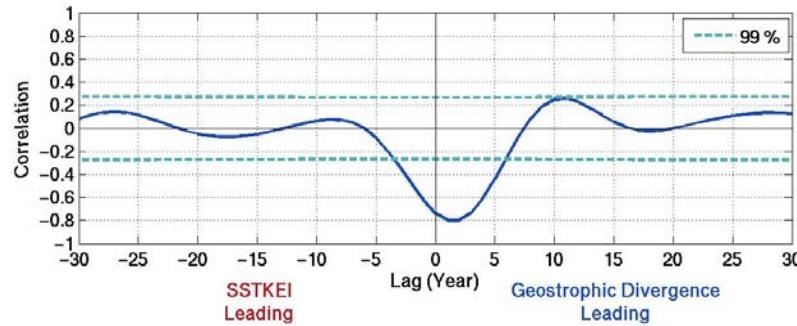
(Contour Interval: 10 W/m²/°C ; Shading: significant in 99 %)

Q_{NET} Forcing / Response to SST Kuroshio Extension Index

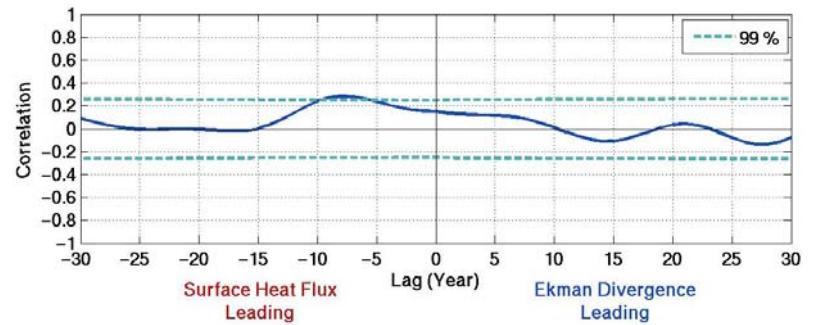
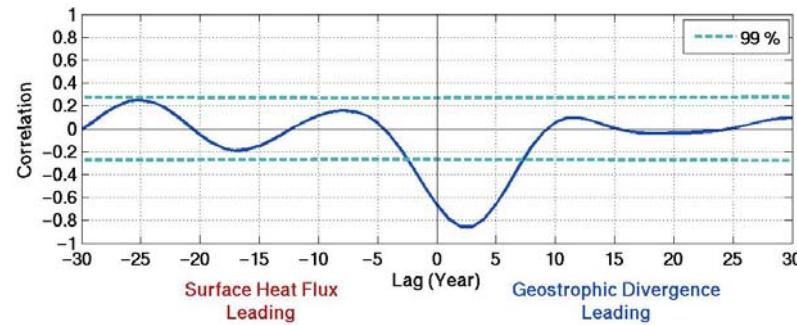
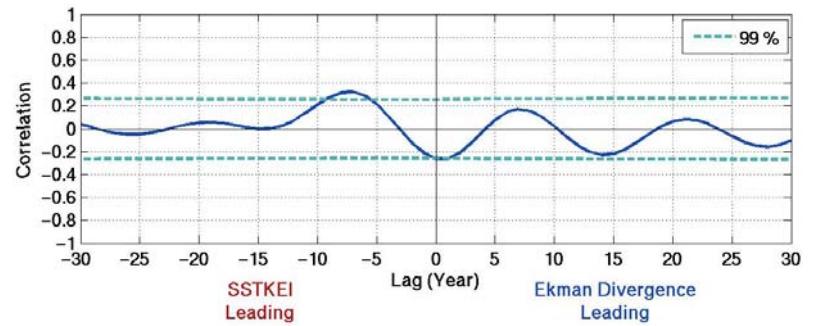


Correlation in the Kuroshio Extension Geostrophic / Ekman Heat Divergence vs. SST / Q_{NET}

Geostrophic Heat Divergence

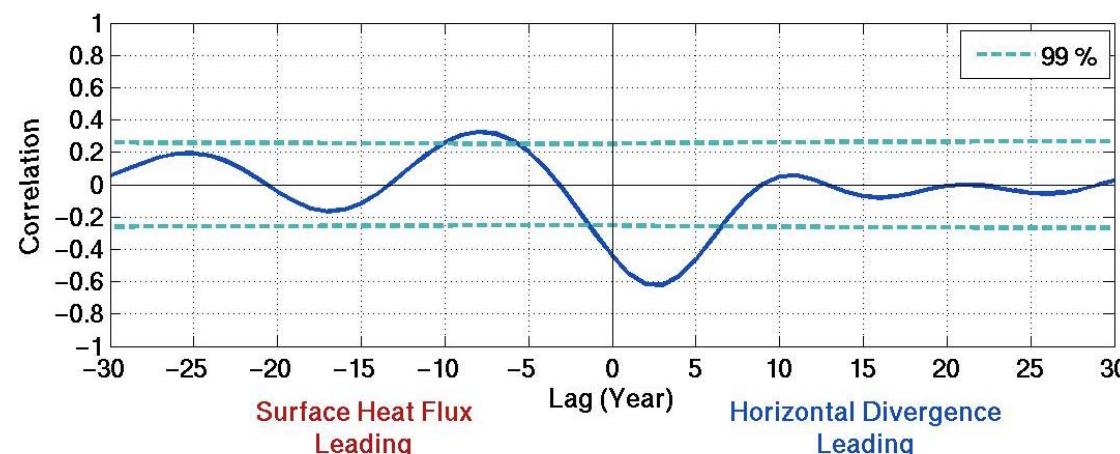
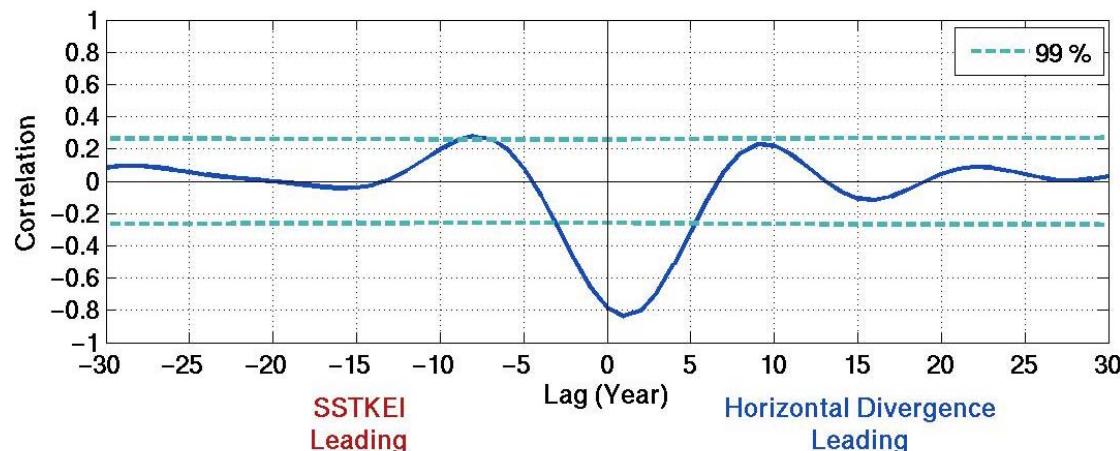


Ekman Heat Divergence



(Low-pass Filter > 10 yr)

Correlation in the Kuroshio Extension Total Horizontal Heat Divergence vs. SST / Q_{NET}



(Low-pass Filter > 10 yr)